

# AMERICAN METEOROLOGICAL SOCIETY

Bulletin of the American Meteorological Society

## EARLY ONLINE RELEASE

This is a preliminary PDF of the author-produced manuscript that has been peer-reviewed and accepted for publication. Since it is being posted so soon after acceptance, it has not yet been copyedited, formatted, or processed by AMS Publications. This preliminary version of the manuscript may be downloaded, distributed, and cited, but please be aware that there will be visual differences and possibly some content differences between this version and the final published version.

The DOI for this manuscript is doi: 10.1175/BAMS-D-15-00031.1

The final published version of this manuscript will replace the preliminary version at the above DOI once it is available.

If you would like to cite this EOR in a separate work, please use the following full citation:

Cheng, L., J. Abraham, G. Goni, T. Boyer, S. Wijffels, R. Cowley, V. Gouretski, F. Reseghetti, S. Kizu, S. Dong, F. Bringas, M. Goes, L. Houpert, J. Sprintall, and J. Zhu, 2015: XBT Science: assessment of instrumental biases and errors. Bull. Amer. Meteor. Soc. doi:10.1175/BAMS-D-15-00031.1, in press.

© 2015 American Meteorological Society



## 1 XBT Science: assessment of instrumental biases

### 2 and errors

- 3 Lijing Cheng<sup>1,\*</sup>, John Abraham<sup>2</sup>, Gustavo Goni<sup>3</sup>, Timothy Boyer<sup>4</sup>, Susan
- 4 Wijffels<sup>5</sup>, Rebecca Cowley<sup>5</sup>, Viktor Gouretski<sup>6</sup>, Franco Reseghetti<sup>7</sup>, Shoichi
- 5 Kizu<sup>8</sup>, Shenfu Dong<sup>9,3</sup>, Francis Bringas<sup>3</sup>, Marlos Goes<sup>9,3</sup>, Loïc Houpert<sup>10</sup>,
- 6 Janet Sprintall<sup>11</sup>, Jiang Zhu<sup>1</sup>
- 7 <sup>1.</sup> International Center for Climate and Environment Science, Institute of
- 8 Atmospheric Physics, Chinese Academy of Sciences, Beijing, China
- 9 <sup>2.</sup> University of St. Thomas, St. Paul, MN, USA
- <sup>3</sup> Atlantic Oceanographic and Meteorological Laboratory, NOAA, Miami, FL, USA
- <sup>4.</sup> National Oceanographic Data Centre, Silver Spring, MD, USA
- 12<sup>5.</sup> CSIRO Oceans & Atmosphere Flagship, Hobart, Tasmania, Australia
- <sup>6.</sup> University of Hamburg, KlimaCampus, Hamburg, Germany
- <sup>7.</sup> ENEA National Agency for New Technologies, Energy and Sustainable Economic
- 15 Development, S.Teresa, Italy 🐂
- 16<sup>8.</sup> Tohoku University, Sendai, Japan
- <sup>9</sup> CIMAS, University of Miami, Miami, FL, USA
- <sup>10.</sup> Scottish Association for Marine Science, Scottish Marine Institute, Oban, UK
- 19 <sup>11.</sup> Scripps Institution of Oceanography, UCSD, La Jolla, CA, USA

20 CORRESPONDING AUTHOR: Lijing Cheng, International Center for Climate and

- 21 Environment Science, Institute of Atmospheric Physics, Chinese Academy of Sciences,
- 22 Beijing, China, 100029
- 23 Email: chenglij@mail.iap.ac.cn

#### 24 Abstract

25 eXpendable BathyThermograph (XBT) data were the major component of the 26 ocean temperature profile observations from the late 1960s through early 2000s, and 27 XBTs still continue to provide critical data to monitor surface and subsurface currents, 28 meridional heat transport, and ocean heat content. Systematic errors have been 29 identified in the XBT data, some of which originate from computing the depth in the 30 profile using a theoretically- and experimentally derived fall rate equation (FRE). 31 After in-depth studies of these biases and discussions held in several workshops 32 dedicated to discuss XBT biases, the XBT science community met at the Fourth XBT 33 Science Workshop and concluded that XBT biases consist of: 1) errors in depth values 34 due to the inadequacy of the probe motion description done by standard FRE, and 2) 35 independent pure temperature biases. The depth error and temperature bias are 36 temperature dependent and may depend on the data acquisition and recording system. 37 In addition, the depth bias also includes an offset term. Some biases affecting the 38 XBT-derived temperature profiles vary with manufacturer/probe type and have been 39 shown to have a time dependence. Best practices for historical XBT data corrections, 40 recommendations for future collection of metadata to accompany XBT data, impact of 41 XBT biases on scientific applications, and challenges encountered are presented in 42 this manuscript. Analysis of XBT data shows that, despite the existence of these 43 biases, historical XBT data without bias corrections are still suitable for many 44 scientific applications, and that bias corrected data can be used for climate research.

### 45 **Capsule Summary**

- 46 Based on in-depth studies, recommendations for correcting biases in XBT data, and
- 47 the impact on applications and ongoing research to improve the quality of future XBT
- 48 data are provided.

#### 50 INTRODUCTION

51 eXpendable BathyThermographs (XBTs) are probes that provided the major 52 portion of ocean subsurface temperature observations during the late 1960s through 53 early 2000s. XBTs were designed for naval use, to enable quick collection of a sound 54 velocity profile and as such do not have a high accuracy or precision. The research 55 community quickly adopted the technology, and many millions of profiles have since 56 been collected. The use of the data has changed over time and now XBT data is a 57 valuable resource for climate studies, despite the simplicity of the probe design. More 58 than 38% (41% for all profiles deeper than 100m depth) of upper ocean temperature 59 profiles in the World Ocean Database 2013 (Boyer et al. 2013) were provided by 60 XBTs from 1970 to 2001 (Figure 1). Currently, approximately 18,000 XBTs are 61 deployed every year mostly along fixed transects and in high-density mode, where 62 each transect is repeated approximately 4 times per year and the deployments are 63 carried out every 20-30 km (Figure 2). Scientific studies to monitor the variability of 64 surface and subsurface currents and of meridional heat transport along fixed transects, 65 ocean and climate modeling, ocean data assimilation, and climate change attributions 66 rely strongly on XBT data (Goni et al. 2010; Abraham et al. 2013; Rhein et al. 2013), 67 and XBTs continue to provide critical data with a spatial and temporal sampling that 68 cannot be currently obtained using any other observational platform (Boyer et al. 69 2013; Abraham et al. 2013).

Biases in XBT data were identified in the 1970s, soon after XBT manufacture
began. The quantity of the data makes it highly valuable and therefore, much effort
has been expended to correct the known biases. Many authors have attempted to
quantify the size of the bias by comparing XBT profiles with co-located high quality
data from Conductivity, Temperature, Depth (CTD) instruments (Anderson, 1980;

75 Flierl and Robinson, 1977; Hallock and Teague, 1992). Hanawa et al. 1995 collected 76 several hundred XBT/CTD comparisons and provided a global correction for XBT 77 data, which was not time-dependent. The community accepted the new coefficients 78 proposed, and data was collected with a mix of both old and new coefficients. In some 79 cases, the coefficients used were not identified in the metadata, compounding the 80 current problem of providing accurate correction schemes. In the 2007 81 Intergovernmental Panel on Climate Change Fourth Assessment Report (IPCC-AR4), 82 (Bindoff et al., 2007) reported a contradiction between observation-based estimate on 83 historical OHC change and models results. For instance, there is a warming decade 84 during late-1970s~early 1980s, which can neither be explained by any existing theory 85 nor be simulated by climate models. Subsequently, Gouretski and Koltermann 2007 86 identified that XBT biases were time variable and identified a large bias during the 87 1970s to 1980s, partly explaining the incorrect warming reported during this decade. 88 This study triggered more scrutiny of XBT biases, with many scientists working to 89 identify, quantify, and remove the biases in XBT data and ultimately provide a 90 reliable climate quality data set of subsurface ocean temperatures. 91 The international XBT science team, formed in 2011, continues efforts to fully 92 understand and correct the XBT bias. A series of workshops have been held: in 93 Miami/United States; 2008, Hamburg/Germany, 2010; and Melbourne/Australia, 2011. 94 Since 2007, several important factors affecting XBT biases have been identified, and 95 more than 10 correction schemes for the historical XBT dataset proposed. 96 It is timely to summarize the recent progresses in XBT science and provide 97 future guidelines for XBT bias corrections and data applications. Understanding and 98 correcting systematic errors in XBT measurements help to enhance their use for 99 broader ocean and climate studies. On this basis, the Fourth XBT Science Workshop

100 was held in Beijing, China, on November 11-13, 2014, with the participation of 34 101 experts from 11 countries and 18 universities, laboratories, and organizations. This 102 workshop focused on discussing recent advances in assessing XBT data biases and 103 their impact on applications, and on reaching a consensus to recommend bias 104 corrections for the global XBT data set. In this manuscript, we present a summary of 105 XBT science and key recommendations agreed upon by members of the XBT 106 scientific community for correcting historical XBT data, and for best practices in the 107 future collection of XBT observations.

108

#### 109 CORRECTING TIME-DEPENDENT XBT BIASES: RECOMMENDED

#### 110 FACTORS

111 It has been found that biases in XBT data consist of both systematic depth error

and independent pure temperature bias (Reseghetti et al. 2007; Gouretski and

113 Resegnetti 2010; Cowley et al. 2013). These biases have been shown to depend on

several parameters, including the probe-type, water temperature, launch height, and

115 data acquisition system, with the total bias being time dependent (Di Nezio and G.

116 Goni, 2011; Gouretski and Reseghetti 2010; Abraham 2012b; Cowley et al. 2013;

117 Cheng et al, 2014; Bringas and Goni, 2015). Variations in manufacturing processes

and changes in recording systems are identified as the primary source of

time-dependent biases. The exact timing of changes to the XBT system (e.g. analog to

120 digital acquisition system, change of probe nose, thermistor, twin-wire, plastic

121 afterbody, wire-coating, etc) is not known. Further, small variations in the probe

122 physical dimensions over time make it difficult to provide a correct description of the

123 problem. As a result, proposed correction schemes should include a time variation in

each of the correction factors.

During the Fourth XBT Science Workshop, the participants agreed that in order
to correct systematic errors in the historical XBT dataset, the following corrections
(which are equally important) should be performed in order to improve the quality of
XBT data:

129

130 1) Fall-rate equation coefficients: These are the coefficients in the fall rate equation 131 (FRE). The FRE models the free falling motion of the XBT probe in the water. The FRE has the form:  $z(t)=at-bt^2$ ; where t is the elapsed time of the descent of the probe 132 133 in the water in seconds, and a and b are fall rate coefficients, representing the initial 134 fall rate and deceleration, respectively (Green 1984). The coefficients a and b in the 135 FRE have been shown to have variability in time (Hanawa and Yoritaka 1987; 136 Hanawa et al. 1995; Gouretski and Reseghetti 2010). Numerous studies show that the 137 depths calculated using the manufacturer coefficients (originally developed by Sippican (main manufacturer of the XBT probes, now acquired by Lockheed Martin)) 138 139 have a systematic bias (e.g. Flierl and Robinson 1977; Hanawa and Yoritaka 1987; 140 Singer 1990). In the mid-1990s, a research group under the coordination of Integrated 141 Global Ocean Services System (IGOSS) (Hanawa et al. 1995) updated the original 142 FRE coefficients for the most commonly used XBT probe types, based on 143 comparisons with the more accurate collocated data obtained by CTD profilers. These 144 new coefficients were expected to fully correct the fall rate biases. However, it was 145 later shown that the FRE coefficients a and b were time-dependent (Wijffels et al. 146 2008; Di Nezio and Goni 2011; Cowley et al. 2013; Cheng et al. 2014) and that 147 variations in the FRE coefficients existed due to probe type and manufacturer, which 148 were previously thought to behave identically (Kizu et al. 2005a; 2005b; Gouretski 149 and Reseghetti 2010; Abraham et al. 2012a; Kizu et al. 2011; Cowley et al. 2013;

150 Cheng et al. 2014). In addition, it has also been shown that the systematic depth errors 151 are a function of water temperature (Thadathil et al. 2002; Kizu et al. 2005b; Cheng et 152 al. 2014). Water viscosity is highly dependent on its temperature, which affects the 153 probe motion. Further studies are needed to correctly quantify this effect.

154

155 2) Pure temperature bias correction: The correction of biases not included in the 156 depth estimates and that are temperature dependent. Studies have shown that XBT 157 recording systems have the largest impact on the pure temperature bias. Analog 158 recording systems were mainly used before 1985 and have been found to produce 159 positive pure temperature biases of approximately 0.15°C (Emery et al. 1986), and 160 0.13°C (Heinmiller 1983; Cowley et al. 2013). Digital systems, which were mainly 161 used after 1989, produced usually smaller biases, in the most cases with positive 162 values ranging from 0.01 to 0.07°C (Emery et al. 1986; Bailey et al. 1989; Wright 163 1991; Kizu and Hanawa 2002b; Cowley et al. 2013). This pure temperature bias is 164 also due to inaccuracies in the data acquisition system (thermistor, copper wire, cables, 165 digitizer, electronics, and computer) (Heinmiller et al. 1983; Green 1984; Reseghetti 166 et al. 2007). It has been also reported that the pure temperature bias is variable with 167 time and probe type (e.g. Gouretski and Reseghetti 2010; Cowley et al. 2013; Hamon 168 et al. 2013; Cheng et al. 2014). This bias has been observed to robustly increase with 169 the temperature of water flowing past the XBT thermistor (Reverdin et al. 2009; 170 Cowley et al. 2013; Cheng et al. 2014). However, the reason for this temperature 171 dependency and recorder dependency is not yet fully understood. 172 173 3) Depth offset correction. Recent studies show that the XBT depth can be better

174 corrected by adding an offset term to the FRE ( $Depth=at-bt^2-offset$ ), which was not

175 allowed conventionally (Cowley et al. 2013; Cheng et al. 2014). This idea was largely 176 introduced for the reduction of subsurface (i.e. 0-50m) depth bias, which could not be 177 achieved only by modifying the two coefficients in the traditional FRE. For the 178 purpose of correcting XBT data, the depth offset has been estimated in previous 179 studies, and considered to originate from various sources (Cowley et al. 2013; Cheng 180 et al. 2014). The original FRE assumes that the XBTs instantly reach their terminal 181 velocity (a) because the coefficient in the quadratic term (b) is much smaller than the 182 coefficient in the linear term. However, recent studies show that this is not true in the 183 early part of the XBT descent (Gouretski and Reseghetti 2010; Cowley et al. 2013; 184 Cheng et al. 2014). The XBT terminal velocity has been assessed and it was 185 determined that the XBTs take up to approximately 1.5s or 20 meters to reach their 186 terminal velocity after they hit the water (e.g. Hallock and Teague 1992; Bringas and 187 Goni 2015). Comparisons between XBT and CTD measurements indicate that there is 188 a depth bias in the initial probe descent period (upper 50 m) (Gouretski and 189 Reseghetti 2010; Cowley et al. 2013; Cheng et al. 2014). Field tests carried out in 190 shallow water (Gouretski and Reseghetti 2010) and in water tanks (Bringas and Goni 191 2015), and numerical simulations of the XBT falling motion (Abraham et al. 2014; 192 Gorman et al. 2014; Shepard et al. 2014) confirmed this finding. Recent studies 193 confirmed that this depth offset is linked to the initial fall velocity of the XBT in the 194 water, which is a function of the XBT deployment height (Bringas and Goni 2015), 195 and, based on numerical simulations, it has been hypothesized to also depend on the 196 conditions of the probe entry into the water (Abraham et al. 2014; Gorman et al. 2014; 197 Shepard et al. 2014). In addition, one study shows that there may be a time offset that 198 translates into a depth offset at the surface caused by timing errors of the data 199 acquisition system (Thresher, 2014) or malfunctioning of the electronics called

200 "pre-mature start". An offset term to the FRE that is a function of the deployment

201 height has been proposed (Bringas and Goni, 2015). This offset term is derived from

an earlier model (Hallock and Teague, 1992), and it is time dependent during the first

203 1.5 s of the XBT descent in the water and constant after that. Research is currently

204 underway to further explore additional sources of the depth offset.

205

#### 206 THE IMPORTANCE OF METADATA

207 The dependency of the biases on time (e.g. manufacture date, system changes)

and probe type have been highlighted in several studies. There are two major

209 manufacturers of XBT probes, Lockheed Martin Sippican, Inc. in the United States;

and the Tsurumi-Seiki Co., Ltd (TSK), in Japan. Each company produces several

types of probes with different maximum depths and for different ship speeds. For

212 Sippican, they are T4 (460 m), T5 (1830 m), T6 (460 m), T7 (760 m), Deep Blue (760

213 m), Fast Deep (1000 m), T10 (200 m), and T11(460 m)

214 (http://www.sippican.com/stuff/contentmgr/files/0dad831400ede7b5f71cf7885fdeb11

215 0/sheet/xbtxsv92005.pdf). For TSK manufactured probes, they are T4 (460 m), T5

216 (1830 m), T6 (460 m), T7 (760 m), and T10 (300 m)

217 (http://www.tsk-jp.com/index.php?page=/product/detail/2/2). Moreover, the

218 maximum depth reached by XBTs is frequently deeper than the nominal values

219 indicated by Sippican and TSK, adding a further contribution to the uncertainty. Their

220 manufacture is fully independent, except for sharing the basic design and using

thermistors of a single brand, and their probes have many differences in their structure

(Kizu et al. 2005b; 2011).

It has been shown that probe types from different manufacturers have distinct

values of bias (Kizu et al. 2005a,b; Ishii et al. 2009; Kizu et al. 2011; Cheng et al.

225 2014). Therefore, bias corrections for each probe type should be assessed separately.

226 Since a large number of available XBT metadata do not include the information of

probe type (Abraham et al. 2013; Cheng et al. 2014), the available metadata, such as

institution and country that carried out the deployment, year of deployment, and

- 229 maximum observed depth, can be used to infer the probe type (Cowley et al. 2013).
- 230

#### 231 MINIMUM REQUIREMENTS FOR XBT METADATA

232 In addition to the standard requirements for metadata for all oceanographic data 233 (e.g., position and time, platform and instrument type information), it is recommended 234 that the following minimum requirements for XBT metadata are included: fall rate 235 coefficients used in the profile, probe type, probe manufacture date and serial number, 236 manufacturer, launch height, type of recording system and software version. It is 237 critically important that no correction scheme is applied to raw XBT data. All 238 archived data should only contain depths calculated from either the manufacturers or 239 the Hanawa et al 1995 coefficients, and temperatures obtained from the collection 240 system.

241

#### 242 CURRENT XBT BIAS CORRECTION SCHEMES

A suite of correction schemes for global historical XBT datasets has been

proposed (Hanawa et al. 1995; Wijffels et al. 2008; Levitus et al. 2009; Ishii and

Kimoto 2009; Gouretski and Reseghetti 2010; Good 2011; Gouretski 2012; Hamon et

al. 2012; Cowley et al. 2013; Cheng et al. 2014). One goal of the Fourth XBT Science

- 247 Workshop was to assess the respective advantages of each of these schemes. The
- 248 workshop participants recommended that correction schemes should correctly account
- for all of the above-discussed parameters, in particular when the XBT data are used

250 for global scale climate research applications. Table 1 lists the factors considered by 251 each scheme. Within these schemes, the above three correction factors (with their two 252 temperature dependent corollaries) along with probe and year variances are explicitly 253 accounted for in only one of them (Cheng et al. 2014), while it is implicitly accounted 254 for in another scheme (Gouretski and Reseghetti 2010). The participants also noted 255 that XBT data sets without all recommended corrections could still be used for several 256 specific applications (discussed below) in which the XBT biases have a minor impact 257 on the results.

258

At the present time, members of the community recommend the Cheng et al. 259 2014 corrections be used because the scheme currently provides the most appropriate 260 bias correction strategy as discussed above, including all of the recommended factors. 261 The performance of all the correction schemes is the subject of a soon to be published 262 study lead by the XBT community members. A preliminary inter-comparison among 263 the 10 correction schemes using metrics to define "goodness" of a scheme, indicates 264 that Cheng et al. 2014, Gouretski and Reseghetti 2010 and Levitus et al. 2009 265 currently outperform other schemes. However, we note that the recommended 266 correction scheme will change as more research leads to improved understanding of 267 the biases and improved quality-control processes. The community will keep 268 scientists informed on updates to the recommended correction scheme. 269 The online NOAA/NODC-WOD13 selection tool provides the ability to 270 automatically apply to the NOAA/NODC-WOD-13 dataset (Boyer et al., 2013) any of 271 the 10 correction schemes currently available for XBT data. 272 273 **BIAS-CORRECTED XBT DATA: APPLICATIONS** 

274 XBT data have been widely used in oceanography and climate studies over the

275 past 40 years (Goni et al. 2010) and continue to be used for a wide range of critical 276 scientific applications. The impact of the XBT bias on applications and products 277 varies. The main scientific applications of XBT data, and the impact of the application 278 of XBT corrections are shown below. We note that these assessments are based on 279 case studies and, therefore, provide only general guidelines for application of 280 corrections. The use of the best data set available is always preferred, however, data 281 sets that do not contain all of the corrections indicated above are still acceptable for a 282 wide range of oceanographic applications.

283

284 1) Global ocean heat content (OHC). OHC is an indicator of the amount of heat

stored over a certain depth range in the ocean. The major source of error for historical

OHC estimates comes from the XBT biases (Lyman et al. 2010; Boyer et al. 2015). It

287 has been shown that the long-term OHC trend calculated by using uncorrected XBT

data was underestimated by a half (Domingues et al. 2008), and created a spurious

decadal variation ("the warming decade", Gouretski and Koltermann 2007; Levitus et

al. 2009; Rhein et al. 2013) (Figure 3). Uncertainties in OHC (0-700m) estimation

induced by different XBT correction schemes range from 8.2 (11.8) to 16.1 (19.6) ZJ

292 (1  $ZJ = 1 \times 10^{21}$  Joules) for 1970-2008 and for years 1993-2008 in parentheses,

depending on the mapping technique used (Boyer et al. 2015).

294 2) Ocean reanalysis/data assimilation. Giese et al. (2010) documented differences in

295 OHC, ocean temperature structure, and velocity of ocean currents due to XBT

296 corrections in the context of global analyses experiments using a Simple Ocean Data

- Assimilation (SODA) system. The quantified impact of the different correction
- schemes on these variables show that: the Levitus et al. (2009) scheme reduced the
- temperature anomalies at 50 m in the eastern equatorial Pacific by 10%–20% and

300 strengthened the zonal currents by ~50% during the 1997-2000 ENSO cycle

301 compared with the Hanawa et al. (1995) correction; while the Wijffels et al. (2008)
302 scheme had little impact on the ENSO representation in the ocean. Therefore, these
303 results indicate that XBT data sets with more accurate correction schemes serve to

304 provide improved estimates of long period ocean signals.

305 3) Trans-basin ocean meridional heat transport (MHT). Results from numerical

- 306 model studies carried out for the South Atlantic Ocean (Goes et al. 2015) show that
- 307 XBT biases need to be corrected in order to detect MHT trends in the South Atlantic.
- 308 The trends in MHT and meridional overturning circulation (MOC) caused by XBT

309 biases are statistically significant after the 1990s, estimated to be 0.02 PW/decade and

310 0.3 Sv/decade, respectively. These trends are higher than the actual trends estimated

from reanalysis data of 0.006 PW/decade and 0.1 Sv/decade for MHT and MOC,

312 respectively. Therefore, appropriate XBT bias corrections with long-term monitoring

313 may reduce the errors for detection of long-term trends. On the other hand, the errors

in the MHT and MOC due to XBT biases are small in comparison to their seasonal

and interannual variability. Therefore, XBT data without corrections can still provide

reliably assessments of MHT and MOC variability on seasonal-to-interannual time

317 scales (Dong et al. 2009).

**4) Geostrophic currents.** Geostrophic current estimates show that errors due to pure

temperature XBT biases, as estimated for tropical Atlantic currents (Goes et al. 2013b)

320 are very small (<0.02 m/s). The same study shows that the maximum geostrophic

321 velocity errors due to XBT depth biases are likely to be < 0.2 m/s, which is

322 comparable to the errors associated with satellite altimetry estimates of current

323 velocities. Therefore, XBT data without corrections can still provide reliably

324 assessments of geostrophic currents.

325 5) Mixed layer depth (MLD). Studies carried out in the Mediterranean Sea showed

that the Cowley et al. (2013) scheme (applied to 45% of the database) did not

327 significantly affect the estimates of the seasonal cycle of the basin-mean of MLD

- 328 (Houpert et al. 2014), as the differences of the MLD with and without XBT
- 329 corrections are two orders of magnitude lower than the amplitude of the seasonal
- 330 variations. Other studies (Gopalakrishna et al. 2010) examined changes in a particular
- isotherm depth from XBT data to investigate features of the Arabian Sea and noted
- that the results were not sensitive to XBT correction.
- 333

#### 334 FUTURE WORK

Extensive progress has been made during the past decades regarding the

336 understanding and assessments of XBT biases and errors. Similarly to corrections

made to data obtained from other observational platforms, continuous efforts will be

338 made to improve the XBT data set. In particular, the following steps are

339 recommended:

340 1) Continue distribution through the main data centers of data with different XBT

341 corrections. At present, the US NOAA/NODC distributes datasets with the 10

342 different correction schemes (Table 1) applied

343 (http://www.nodc.noaa.gov/cgi-bin/OC5/SELECT/builder.pl), the UK Met Office

344 provides datasets with 3 correction schemes (Good et al. 2013), and the China IAP

distributes Cheng et al. (2014) corrected XBT data (http://159.226.119.60/cheng/).

346 Updates to the recommended correction schemes will be posted via the XBT Science

347 Team website (http://www.aoml.noaa.gov/phod/goos/xbtscience/index.php).

348 Future improvements in the data sets will rely on progress made in the following

two areas: (a) Ongoing assessment of XBT biases and errors including a

350 comprehensive intercomparison of the performance of the existing XBT correction 351 schemes, since it is possible that the inclusion of all correction factors does not 352 guarantee providing better data. An intercomparison is currently being undertaken by 353 the XBT community (http://159.226.119.60/cheng/); and (b) Continuous efforts to 354 improve the quality of XBT data with appropriate flags and uncertainties as part of a 355 recently initiated international project: International Quality-controlled Ocean 356 Database (IQuOD) (http://www.iquod.org). The sensitivity of the correction schemes 357 to dataset versions with more data and higher quality control requires more 358 investigation. 359 2) Require that XBT data originators submit the complete metadata to the major data 360 centers (e.g. NODC). Real time data transmitted via the Global Transmission System 361 (GTS) should preferably be submitted using BUFR format to allow the inclusion of 362 all metadata, which is then archived by the data centers. The metadata must include, 363 in addition to existing requirements for all oceanographic data, information on the fall 364 rate coefficients used in the profile, probe type, probe manufacture date, serial number, 365 manufacturer, launch height, type of recording system, and software version. The 366 metadata recommendations will be submitted to Ship Of Opportunity Programme 367 Implementation Panel (SOOPIP) in the Joint WMO/IOC Technical Commission for 368 Oceanography and Marine Meteorology (JCOMM), International Oceanographic 369 Data and Information Exchange (IODE), for approval and then will be disseminated 370 through these organizations. 371 3) Recover historical side-by-side XBT/CTD comparison data. Side-by-side 372 XBT/CTD comparisons enable us to accurately assess XBT bias and assess proposed 373 correction schemes (Cowley et al. 2013). A highly valuable collection of historical 374 datasets with XBT and CTD co-located pairs is currently maintained at

- 375 https://data.csiro.au (DOI: 10.4225/08/543F60A3F1690). All data in the pairs
- 376 database is also present in the World Ocean Database maintained by US

377 NOAA/NODC (http://www.nodc.noaa.gov/OC5/WOD13/). Ongoing addition of

- 378 historical XBT/CTD pairs to the pairs database via submission to US NOAA/NODC
- is strongly encouraged.
- **4)** Assess the cause for the existence of time-varying biases in different probe types. It
- has been hypothesized that slight differences in probe design may result in probe-type
- differences (e.g. Kizu et al. 2005b; Kizu et al. 2011). Recent advances in
- 383 Computational Fluid Dynamics (CFD) may also help to address this question by
- 384 simulating the real characteristics in XBT probe design to examine the differences in
- 385 fall rate (Abraham et al. 2012a, 2012b).
- **5**) Further investigate and assess all parameters that contribute to the depth offset.
- 387 CFD models have identified the launch height (Abraham et al. 2012a, 2012b) and
- 388 water tank experiments (Bringas and Goni 2015) have identified the initial speed of
- descent of XBTs in the water (a function of the launch height) as having the largest
- impact on initial fall rate. Further tests in the ocean have already been conducted to
- 391 confirm these findings and to investigate other parameters that may impact the value
- 392 of the offset. The possibility of clock offsets in digital recording systems adding to the
- depth offset is also being investigated using a precise timing test.
- **6)** Assess the link between water temperature and pure temperature bias, a topic that is
- 395 rarely discussed in the historical XBT literature. Theoretical analysis, bath calibration,
- and more side-by-side XBT/CTD tests in water of different temperatures (or different
- 397 geographical locations) are required to address this.
- 398 7) Evaluate why higher positive temperature bias exists in XBT data collected with
- analog recorders. Intercomparisons of digital devices and still available strip chart
  - 17

recorders may help to evaluate and understand the cause of pure temperature bias and
its temporal variability. Well-designed bath tests may also help to confirm the impact
of the recorder system on the pure temperature bias. Preservation of old acquisition
systems is also desirable for future assessment.

404 8) Improve and continue communications with the XBT manufacturers in order to

405 improve XBT probes. It has been shown that the depth bias could be reduced when

adding one or more pressure switches to XBTs (Goes et al. 2013a). However, the cost

407 of pressure switches precludes their use in a probe that has been designed to provide

408 cost-effective temperature profiles. As a result of this work and after discussions with

409 Sippican, the XBT community recommended that the manufacturers employ tighter

410 controls on probe weight and better calibration of thermistors during the

411 manufacturing process as a more cost-efficient way of reducing biases. In addition,

the community will continue to collect XBT and CTD side-by-side data for the most

413 common probe types to continue assessment of XBT biases.

414

#### 415 SUMMARY

416 XBT data make up a significant amount of the global historical upper-ocean 417 temperature profile database and are still used extensively to study ocean boundary 418 currents, ocean heat content, climate change, and meridional heat transport. Some 419 applications for which XBT data are used require these data to be accurately corrected 420 for depth and temperature biases. Bias corrections have been applied successfully to 421 XBT data in ocean heat content studies (e.g. Domingues et al. 2008; Levitus et al. 422 2012; Boyer et al. 2015). The increasing number of scientific applications for which 423 XBT data are used and the existence of many different bias corrections proposed over 424 the last 30 years highlight the need to propose a corrected historical dataset for

425 climate and oceanographic related studies. This manuscript reports the progress made

426 on XBT bias studies and provides a guide for future data and metadata collection

427 requirements. At the present time, there is one correction scheme (Cheng et al. 2014)

428 that takes account of all the recommended elements. As such, it is currently

429 recommended as the most appropriate correction for XBT data used in calculations of

430 global ocean heat content and ocean reanalysis and data assimilation. Based on

431 previous studies, corrections are not required for calculations of MHT/MOC,

432 geostrophic currents and mixed layer depth calculations.

433 Similar to data obtained from all observational platforms, efforts will continue to

434 be carried out to improve XBT data quality. The XBT Science Team

435 (http://www.aoml.noaa.gov/phod/goos/xbtscience/index.php) and community will

436 continue working on enhancing our understanding of the XBT fall rate biases by

437 addressing the questions posted above, continuously assessing all available data sets

438 and correction schemes to correct historical and future XBT data, improving the

439 quality of XBT profiles for climate research, and providing future recommendations

440 for XBT bias corrections.

441

442 **ACKNOWLEDGMENTS**. We acknowledge the International Center for Climate 443 and Environment Science (ICCES), Institute of Atmospheric Physics (IAP), Chinese 444 Academy of Sciences (CAS), and Third World Academy of Sciences (TWAS) for supporting the 4<sup>th</sup> XBT workshop, which provided the opportunity for scientists in the 445 446 XBT community to meet and discuss XBT biases. L. Cheng and J. Zhu are supported 447 by the project "Structures, Variability and Climatic Impacts of Ocean Circulation and 448 Warm Pool in the Tropical Pacific Ocean" of National Basic Research Program of 449 China (grant no. 41476016), Chinese Academy Sciences' Project "Western Pacific

- 450 Ocean System: Structure, Dynamics and Consequences" (grant no. XDA11010405).
- 451 Work by G. Goni, S. Dong, F. Bringas, and M. Goes was supported by NOAA/AOML
- 452 and by the NOAA Climate Program Office.

#### 454 **REFERENCES**

- 455 Abraham, J. P., J. M. Gorman, F. Reseghetti, E. M. Sparrow, and W. J. Minkowycz,
- 456 2012a: Turbulent and Transitional Modeling of Drag on Oceanographic
- 457 Measurement Devices. *Modelling and Simulation in Engineering*, **2012**, 8.
- 458 http://dx.doi.org/10.1155/2012/567864.
- 459 —, 2012b: Drag Coefficients for Rotating Expendable Bathythermographs and
- the Impact of Launch Parameters on Depth Predictions. *Numer Heat Tr a-Appl*,
- **62,** 25-43.
- 462 Abraham, J. P., and Coauthors, 2013: A Review of Global Ocean Temperature
- 463 Observations: Implications for Ocean Heat Content Estimates and Climate
- 464 Change. *Rev Geophys*, **51**, 450-483, doi: 10.1002/rog.20022.
- 465 Abraham, J. P., J. M. Gorman, F. Reseghetti, E. M. Sparrow, J. Spark, T. Shepard,
- 466 2014: Modeling and Numerical Simulation of the Forces Acting on a Sphere
- 467 During Early-Water Entry. Ocean. Eng., 76, 1-9,
- 468 doi:10.1016/j.oceaneng.2013.11.015.
- 469 Anderson, E. R., 1980: Expendable Bathythermograph(XBT) accuracy studies.
- 470 *Technical Report 550, Naval Ocean Systems Center, San Diego, California*,
  471 201pp.
- 472 Bailey, R., H. Phillips, and G. Meyers, 1989: Relevance to TOGA of systematic
- 473 XBT errors. Proc. Western Pacific International Meeting and Workshop on
- 474 *TOGA COARE*, Noumea, New Caledonia, May 24-30, 1989, 775-784.
- 475 Bindoff, N. L. and Coauthors, 2007: Observations: Oceanic Climate Change and Sea

476	Level, in Climate Change 2007: The Physical Science Basis. Contribution of
477	Working Group I to the Fourth Assessment Report of the Intergovernmental
478	Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M.
479	Marquis, K.B. Averyt, M. Tignor and H.L. Miller (eds.)], Cambridge
480	University Press, Cambridg, UK.
481	Boyer, T., and Coauthors, 2015: Sensitivity of Global Ocean Heat Content Estimates
482	to Mapping Methods, XBT Bias Corrections, and Baseline Climatology. The
483	4th XBT workshop: XBT Science and the Way Forward, Beijing, China,
484	Institute of Atmospheric Physics, Chinese Academy of Sciences. [Available
485	online at http://2014xbtworkshop.csp.escience.cn/dct/page/70005]
486	Boyer, T. P., and Coauthors, 2013: World Ocean Database 2013. Sydney Levitus,
487	Ed.; Alexey Mishonov, Technical Ed.; NOAA Atlas NESDIS 72, 209 pp.
488	Bringas F., and Goni, G., 2015: Early Dynamics of Deep Blue XBT Probes, J. Atmos.
489	Oceanic Technol.,, under review.
490	Cheng, L., J. Zhu, R. Cowley, T. Boyer, and S. Wijffels, 2014: Time, Probe Type and
491	Temperature Variable Bias Corrections to Historical Expendable
492	Bathythermograph Observations. J. Atmos. Oceanic Technol., 31, 8,
493	1793-1825.
494	Cheng, L. J., J. Zhu, F. Reseghetti, and Q. P. Liu, 2011: A New Method to Estimate
495	the Systematical Biases of Expendable Bathythermograph. J. Atmos. Oceanic
496	Technol., <b>28</b> , 244-265.
497	Cowley, R., S. Wijffels, L. Cheng, T. Boyer, and S. Kizu, 2013: Biases in

498	Expendable Bathythermograph Data: A New View Based on Historical
499	Side-by-Side Comparisons. J. Atmos. Oceanic Technol., 30, 1195-1225.
500	Di Nezio, P. N., and G. Goni, 2011: Direct Evidence of Changes in the XBT
501	Fall-rate Bias During 1986-2008. J. Atmos. Oceanic Technol., 28, 1569-1578.
502	Domingues, C. M., J. A. Church, N. J. White, P. J. Gleckler, S. E. Wijffels, P. M.
503	Barker, and J. R. Dunn, 2008: Improved estimates of upper-ocean warming
504	and multi-decadal sea-level rise. Nature, 453, 1090-U1096.
505	Dong, S. S. L. Garzoli, M. O. Baringer, C. S. Meinen, and G. J. Goni, 2009: The
506	Atlantic Meridional Overturning Circulation and its Northward Heat Transport
507	in the South Atlantic. Geophys. Res. Lett., 36, L20606,
508	doi:10.1029/2009GL039356.
509	Emery, W. J., W. Lee, W. Zenk, and M. J., 1986: A low-cost digital XBT system and
510	its application to the real-time computation of dynamic height. J. Atmos.
511	<i>Oceanic Technol.</i> , <b>3</b> , 75-83.
512	Flierl, G. R., and A. R. Robinson, 1977: XBT measurements of thermal gradients in
513	the MODE eddy. J. Phys. Oceanogr., 7, 300-302.
514	Giese, B. S., G. A. Chepurin, J. A. Carton, T. P. Boyer, and H. F. Seidel, 2010:
515	Impact of Bathythermograph Temperature Bias Models on an Ocean
516	Reanalysis. J. Climate, 24.

- 517 Goes, M., G. Goni, and K. Keller, 2013a: Reducing Biases in XBT Measurements
- 518 by Including Discrete Information from Pressure Switches. J. Atmos. Oceanic
- 519 *Technol.*, **30**, 810-824.

520	Goes, M., G. J. Goni, V. Hormann, and R. C. Perez, 2013b: Variability of eastward
521	currents in the equatorial Atlantic during 1993-2010. J. Geophys. Res., 118,
522	3026-3045.
523	Goes, M. M., M. Baringer, and G. Goni, 2015: The impact of historical biases on the
524	XBT-derived meridional overturning circulation estimates at 34°S, Geophys.
525	Res. Let., doi:10.1002/2014GL061802.
526	Goni G.J., and Coauthors, 2010: The Ship Of Opportunity Program. In Proceedings of
527	the "OceanObs'09: Sustained Ocean Observations and Information for Society"
528	Conference (Vol. 2), Venice, Italy, 21-25 September 2009, Hall, J., Harrison
529	D.E. and Stammer, D., Eds., ESA Publication WPP-306.
530	Good, S. A., 2011: Depth Biases in XBT Data Diagnosed Using Bathymetry Data. J.
531	Atmos. Oceanic Technol., 28, 287-300.
532	Good, S. A., M. J. Martin, and N. A. Rayner, 2013: EN4: Quality controlled ocean
533	temperature and salinity profiles and monthly objective analyses with
534	uncertainty estimates, J Geophys Res-Oceans, 118, 6704-6716.
535	Gopalakrishna V., and Coauthors, 2010: Observed intra-seasonal to interannual
536	variability of the upper ocean thermal structure in the southeastern Arabian
537	Sea during 2002-2008, Deep-Sea Res Pt I, 76, 6, 739-754,
538	doi:10.1016/j.dsr.2010.03.010.

- 539 Gorman, J. M., J. P. Abraham, D. B. Schwalbach, T. S. Shepard, J. R. Stark, and F.
- 540 Reseghetti, 2014: Experimental Verification of Drag Forces on Spherical
- 541 Objects Entering Water. J Mar. Ocean Biol. Ocean, **3**, 2,
- 542 http://dx.doi.org/10.4172/2324-8661.1000126.

- 543 Gouretski, V., 2012: Using GEBCO digital bathymetry to infer depth biases in the
- 544 XBT data. *Deep-Sea Res Pt I*, **62**, 40-52.
- 545 Gouretski, V., and K. P. Koltermann, 2007: How much is the ocean really warming?
- 546 *Geophys. Res. Lett.*, **34**, L01610, doi:10.1029/2006GL027834.
- 547 Gouretski, V., and F. Reseghetti, 2010: On depth and temperature biases in
- 548 bathythermograph data: Development of a new correction scheme based on
- analysis of a global ocean database. *Deep-Sea Res Pt I*, **57**, 812-833.
- 550 Green, A. W., 1984: Bulk Dynamics of the Expendable Bathythermograph (Xbt).
- 551 *Deep-Sea Res Pt I* **31**, 415-426.
- Hallock, Z. R., and W. J. Teague, 1992: The fall-rate of the T7 XBT. J. Atmos.
- 553 *Oceanic Technol.*, **9**, 470–483.
- Hamon, M., G. Reverdin, and P. Y. Le Traon, 2012: Empirical correction of XBT
- 555 data. J. Atmos. Oceanic Technol., **29**, 960-973.
- Hanawa, K., and H. Yoritaka, 1987: Detection of systematic errors in XBT data and
  their correction. J. Oceanogr. Soc. Japan, 43, 68-76.
- Hanawa, K., P. Rual, R. Bailey, A. Sy, and M. Szabados, 1995: A New Depth Time
- Equation for Sippican or Tsk T-7, T-6 and T-4 Expendable Bathythermographs
  (Xbt). *Deep-Sea Res. Pt I*, **42**, 1423-1451.
- 561 Heinmiller, R., C. Ebbesmeyer, B. Taft, T. Olson, Nikitin, O., 1983: Systematic
- 562 errors in expendable bathythermograph (XBT) profiles. *J. Oceanogr.*, 65,
  563 287-299.
- Houpert, L., and Coauthors, 2014: Seasonal cycle of the mixed layer, the seasonal
  thermocline and the upper-ocean heat storage rate in the Mediterranean Sea

- derived from observations. *Prog. Oceanogr.*, **132**, 333-352,
- 567 doi: 10.1016/j.pocean.2014.11.004.
- 568 Ishii, M., and M. Kimoto, 2009: Reevaluation of historical ocean heat content
- 569 variations with time-varying XBT and MBT depth bias corrections. J.
- 570 *Oceanogr.*, **65**, 287-299.
- 571 Kizu, S., and K. Hanawa, 2002a: Start-up transient of XBT measurement. *Deep-Sea*572 *Res. Pt I*, 49, 935-940.
- 573 —, 2002b: Recorder-dependent temperature error of expendable
- bathythermograph. J. Oceanogr., **58**, 469-476.
- 575 Kizu, S., H. Yoritaka and K. Hanawa, 2005a: A new fall-rate equation for T-5
- 576 expendable bathythermograph (XBT) by TSK. J. Oceanogr., **61**, 115-121.
- 577 Kizu, S., S. Ito, and T. Watanabe, 2005b: Inter-manufacturer difference and
- 578 temperature dependency of the fall-rate of T-5 expendable bathythermograph.
- 579 *J. Oceanogr.*, **61**, 905-912.
- 580 Kizu, S., C. Sukigara, and K. Hanawa, 2011: Comparison of the fall rate and
- structure of recent T-7 XBT manufactured by Sippican and TSK. *Ocean Sci.*, 7,
  231-244.
- 583 Levitus, S., J. I. Antonov, T. P. Boyer, R. A. Locarnini, H. E. Garcia, and A. V.
- 584 Mishonov, 2009: Global ocean heat content 1955-2008 in light of recently
- revealed instrumentation problems. *Geophys. Res. Lett.*, **36**, L07608,
- 586 doi:10.1029/2008GL037155.
- 587 Levitus, S., and Coauthors, 2012: World ocean heat content and thermosteric sea
- 588 level change (0-2000 m), 1955-2010. *Geophys. Res. Lett.*, **39**, L10603, doi:
- 589 10.1029/2012GL051106
- 590 Lyman, J. M., S. Good, V. Gouretski, M. Ishii, G. Johnson, M. Palmer, D. Smith,

- and J. Willis, 2010: Robust warming of the global upper ocean. *Nature*, 465,334-337.
- 593 Reseghetti, F., M. Borghini, and G. M. R. Manzella, 2007: Factors affecting the 594 quality of XBT data - results of analyses on profiles from the Western 595 Mediterranean Sea. Ocean Sci., 3, 59-75. 596 Reverdin, G., F. Marin, B. Bourles, and P. L'Herminier, 2009: XBT Temperature 597 Errors during French Research Cruises (1999-2007). J. Atmos. Oceanic 598 Technol., 26, 2462-2473. 599 Rhein, M., and Coauthors, 2013: Observations: Ocean. In: Climate Change 2013: 600 The Physical Science Basis. Contribution of Working Group I to the Fifth 601 Assessment Report of the Intergovernmental Panel on Climate Change 602 [Stocker, T.F., D. Qin, G.-K. Plattner, M. Tignor, S.K. Allen, J. Boschung, A. 603 Nauels, Y. Xia, V. Bex and P.M. Midgley (eds.)]. Cambridge University Press, 604 Cambridge, United Kingdom and New York, NY, USA. 605 Roemmich, D., B. Cornuelle, 1987: Digitization and calibration of the expendable 606 bathy thermograph. Deep-Sea Research, 34, 229-307.
  - 607 Shepard, T., J. Abraham, D. Schwalbach, S. Kane, D. Siglin, and T. Harrington,
  - 608 2014: Velocity and Density Effect on Impact Force during Water Entry of
    609 Sphere. *Geophysics & Remote Sensing*, 3, 2169-0049.
  - 610 Singer, J. J., 1990: On the error observed in electronically digitized T-7 XBT data. J.
    611 *Atmos. Oceanic Technol.*, 7, 603-611.
  - 612 Thadathil, P., A. K. Saran, V. V. Gopalakrishna, P. Vethamony, N. Araligidad, and R.
  - Bailey, 2002: XBT fall rate in waters of extreme temperature: A case study in
  - 614 the Antarctic Ocean. J. Atmos. Oceanic Technol., **19**, 391-396.
  - 615 Thresher, A, 2014: Mk21 and Devil/Quoll timing test. *The 4th XBT workshop: XBT*

616	Science and the Way Forward, Beijing, China, Institute of Atmospheric
617	Physics, Chinese Academy of Sciences. [Available online at
618	http://2014xbtworkshop.csp.escience.cn/dct/page/70005]
619	Wijffels, S. E., J. Willis, C. Domingues, P. Barker, N. White, A. Gronell, K. Ridgway,
620	J. Church, 2008: Changing Expendable Bathythermograph Fall Rates and
621	Their Impact on Estimates of Thermosteric Sea Level Rise. J. Climate, 21,
622	5657-5672.
623	Wright, D. M., 1991: Field evaluation of the XBT bowing phenomenon. OOD Data
624	Report 91-2, Ocean Observations Division/Office of Ocean Services, National
625	Ocean Service, NOAA, 12 pp.

**Table 1.** Summary of the ten of the available correction schemes that analyzed historical and global XBT datasets. The check mark denotes
whether a factor has been considered by a given scheme. And the check mark inside a square indicates a scheme partly or implicitly considering
the specific factor.

630 Gouretski & Wijffels et Levitus et Gouretski Hamon et Cowley et Cowley et Ishii and Cheng et al. 2014 Good 2011 Factors Reseghetti 2010 al. 2009 2012 al. 2012 al. 2013 al. 2013 CH Kimio, 2009 al. 2008 Pure temperature  $\checkmark$  $\checkmark$  $\checkmark$  $\checkmark$  $\checkmark$  $\checkmark$  $\checkmark$ bias correction Pure  $\checkmark$  $\checkmark$  $\checkmark$  $\checkmark$  $\checkmark$ Time-variable  $\checkmark$ temperature ☑Cold & bias Temperature  $\checkmark$  $\checkmark$ warm dependency water Depth bias  $\checkmark$  $\checkmark$  $\checkmark$  $\checkmark$  $\checkmark$  $\checkmark$  $\checkmark$  $\checkmark$  $\checkmark$ correction  $\checkmark$ Time-variable < V <  $\checkmark$  $\checkmark$ < < **V** Depth bias Surface depth bias < V V **\** V J Temperature  $\mathbf{\nabla}$  $\checkmark$ dependency T7/DB; T4/T6; T10; T4/T6; T4; T7; T4/T6; T4/T6; T7; T4; T6; Sippican T4/T6; T7/DB T5; T10 T7/DB T10 T7/DB T7/DB; T5; T10; FD; TSK T4/T6; T5; T7 TSK TSK T4; T6; T7 Deep; Deep; Probe Type Deep Deep Unknown; Shallow Shallow Unknown; Shallow Unknown Unknown Shallow Unknown Unknown

#### 631 Figures Caption Lists

632 Fig. 1. Percentage of different instruments in the ocean subsurface temperature

- 633 observation system from 1966 to 2013 (based on WOD2013). Data include
- 634 Expendable Bathythermograph (XBT); Profiling Floats (PFL); Autonomous Pinniped
- 635 Bathythermographs (APB); Mechanical Bathythermographs (MBT); Ocean Station
- 636 Data (OSD); High Resolution CTD/XCTD (CTD); Drifting Buoys (DRB); Undulating
- 637 Oceanographic Recorder (UOR); Glider data (GLD); Moored Buoys (MRB).
- 638 Fig.2. Location of approximately 35,000 global XBT deployments during 2013-2014.
- 639 Colors indicate the length of time the currently operated transect has been maintained.
- 640 Fig. 3. Upper (0-700m) Ocean Heat Content (OHC) calculated using corrected and
- 641 uncorrected XBT data (black curve). The XBT data is corrected using ten of the
- schemes including CH14 (Cheng et al. 2014), L09 (Levitus et al. 2009), GR10
- 643 (Gouretski and Reseghetti, 2010), W08 (Wijffels et al. 2008), IK09 (Ishii and Kimoto,
- 644 2009); GD11 (Good 2011); G12 (Gouretski 2012); H12 (Hamon et al. 2012); CW13
- 645 (Cowley et al. 2013) and CWCH (CH method in Cowley et al. 2013). The annual
- 646 mean of global OHC anomaly (OHCA) is calculated by simply averaging the 1° by 1°
- 647 grid means of OHCA over the global ocean.







Fig. 1. Percentage of different instruments in the ocean subsurface temperature

observation system from 1966 to 2013 (based on WOD2013). Data include

653 Expendable Bathythermograph (XBT); Profiling Floats (PFL); Autonomous Pinniped

Bathythermographs (APB); Mechanical Bathythermographs (MBT); Ocean Station

655 Data (OSD); High Resolution CTD/XCTD (CTD); Drifting Buoys (DRB);

656 Undulating Oceanographic Recorder (UOR); Glider data (GLD); Moored Buoys

657 (MRB).



660 Fig.2. Location of approximately 35,000 global XBT deployments during 2013-2014.

661 Colors indicate the length of time the currently operated transect has been maintained.



Fig. 3. Upper (0-700m) Ocean Heat Content (OHC) calculated using corrected and

uncorrected XBT data (black curve). The XBT data is corrected using ten of the

schemes including CH14 (Cheng et al. 2014), L09 (Levitus et al. 2009), GR10

- 667 (Gouretski and Reseghetti, 2010), W08 (Wijffels et al. 2008), IK09 (Ishii and Kimoto,
- 668 2009); GD11 (Good 2011); G12 (Gouretski 2012); H12 (Hamon et al. 2012); CW13
- 669 (Cowley et al. 2013) and CWCH (CH method in Cowley et al. 2013). The annual
- 670 mean of global OHC anomaly (OHCA) is calculated by simply averaging the 1° by 1°
- 671 grid means of OHCA over the global ocean.