# A Reanalysis of the 1911 to 1920 Atlantic Hurricane Database

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### ABSTRACT

A reanalysis of the Atlantic basin tropical storm and hurricane database ("best track") for the period of 1911 to 1920 has been completed. This reassessment of the main archive for tropical cyclones of the North Atlantic Ocean, Caribbean Sea and Gulf of Mexico was necessary to correct systematic biases and random errors in the data as well as to search for previously unrecognized systems. Methodology for the reanalysis process for revising the track and intensity of tropical cyclone data is provided in detail. The dataset now includes several new tropical cyclones, excludes one system previously considered a tropical storm, makes generally large alterations in the intensity estimates of most tropical cyclones (both toward stronger and weaker intensities), and typically adjusts existing tracks with minor corrections. Average errors in intensity and track values are estimated for both open ocean conditions as well as for landfalling systems. Finally, highlights are given for changes to the more significant hurricanes to impact the United States, Central America and the Caribbean for this decade.

# **1. Introduction**

This paper details efforts to re-analyze the National Hurricane Center's (NHC's) North Atlantic hurricane database (or HURDAT, also called the "best track" since they are the "best" post-season determination of tropical cyclone (TC) tracks and intensities for the period of 1911 to 1920. The original database of six-hourly TC (including tropical storms and hurricanes, but not non-developing tropical depressions) positions and intensities was assembled in the 1960s in support of the Apollo space program to help provide statistical TC track forecasting guidance (Jarvinen et al. 1984). Since its inception, this database<sup>2</sup> has been utilized for a wide variety of additional projects: setting of appropriate building codes for coastal zones (ASCE 1998), risk assessment for emergency managers (Jarrell et al. 1992), analysis of potential losses for insurance and business interests (Malmquist and Michaels 2000), intensity forecasting techniques (DeMaria et al. 1999), verification of official and model predictions of track and intensity (McAdie and Lawrence 2000), seasonal forecasting (Gray 1984), and climatic change studies (Landsea et al. 1999). Unfortunately, HURDAT was not designed with all of these uses in mind when it was first put together and not all of them may be appropriate, given its original motivation and limitations.

There are many reasons why a re-analysis of the HURDAT dataset was both needed and timely. HURDAT contained many systematic biases and random errors that needed correction (Neumann 1994). For example, in the early part of the 20<sup>th</sup> Century, a TC's intensity and position were only estimated once per day, which was later interpolated to six hour intervals for HURDAT. Such a linear interpolation scheme is problematic for systems that make landfall because of the tendency for TCs to retain their intensity until the time that the center crosses the coast followed by a period of exponential decay (Kaplan and DeMaria, 1995). Cases where the TC's winds were

artificially weakened before landfall in HURDAT occurred in a majority of landfalling hurricanes in the first half of the 20<sup>th</sup> Century. Other systematic errors included unrealistic translational velocities at the beginning and/or end of the TC track because of the digitization process in the 1960s and a lack of realistic wind speed decay when a TC traversed substantial peninsulas and islands (such as the Yucatan of Mexico and Hispaniola).

Additionally, as our understanding of TCs developed over the years, analysis techniques at NHC have changed and led to biases in the historical database that have not been addressed. For example, Landsea (1993) documented an artificial change to the central pressure-maximum wind relationship, where the HURDAT winds in the 1940s to the 1960s were systematically stronger than those in the 1970s and 1980s for the same central pressure. Another methodological concern is that the winds in HURDAT just before a hurricane landfall in the United States often do not match the assigned Saffir-Simpson Hurricane Scale. C. Neumann and J. Hope developed the first digital HURDAT records with six hourly position and maximum wind estimates in the late 1960s (Jarvinen et al. 1984), before the Saffir-Simpson Scale was devised (Saffir 1973, Simpson 1974). The U.S. Saffir-Simpson Scale categorizations for the 20<sup>th</sup> Century were first assigned by Hebert and Taylor (1975), based primarily upon central pressure observations or estimates at landfall. It was not until the late 1980s that the use of the Saffir-Simpson Scale categorization was based upon the winds exclusively, which is the current standard at NHC (OFCM 2005). Thus, reanalysis efforts in Landsea et al. (2004a, b) and in the work presented here have utilized the estimated maximum sustained winds for assignment of Saffir-Simpson category in order to be consistent with today's analysis techniques. Finally, new understanding of the wind structure in hurricanes from GPS-based dropwindsondes launched in the eyewalls of hurricanes since 1997 have provided

<sup>&</sup>lt;sup>2</sup> Available at NHC's webpage: http://www.nhc.noaa.gov/pastall.shtml.

a systematic way to adjust aircraft flight-level winds to the surface (Dunion et al. 2003, Franklin et al. 2003). This new methodology has already been applied to 1992 Hurricane Andrew (Landsea et al. 2004b) and resulted in numerous revisions to that TC's wind speed records. Such standardization will be crucial for reanalysis efforts during the post-1943 reconnaissance era, as aircraft data have provided a substantial portion of HURDAT wind speed estimates during last several decades.

The first phase of the reanalysis efforts for the period of 1851 through 1910 was reported on in Landsea et al. (2004a). That earlier work covered the era that was first fully investigated by Partagas and Diaz (1996) and resulted in the introduction of 240 TCs during a period of 35 years (1851 to 1885) in HURDAT, detailed 22 new TCs from 1886 to 1910, and made alterations to about 200 other tropical storms and hurricanes in that latter time period. The current paper moves forward sequentially in time to the second decade of the 20<sup>th</sup> Century.

Data sources will be described in the next section followed by a discussion of the methodologies used to estimate TC track and intensity, their likely errors, and criteria utilized to either add new TCs or to remove systems from HURDAT. The Results section goes through the overall changes implemented for the 1911 through 1920 timeframe and highlights changes in some of the more noteworthy hurricanes that have impacted the United States and other countries in the North Atlantic basin. The Summary and Future Work section revisits the larger points within the paper and mentions the directions to be taken to move forward with the project. Finally, an appendix describes in full the reanalysis of a single TC that occurred during this period – the 1919 Key West hurricane.

### 2. Data Sources:

The Atlantic HURDAT contains six-hourly intensity (maximum sustained 1-minute winds at the surface [10 m] and, when available, central pressures) and position (to the nearest  $0.1^{\circ}$ latitude and longitude) estimates of all known tropical storms and hurricanes from 1851 to today (Jarvinen et al. 1984, Landsea et al. 2004a). Tropical storms and hurricanes that remained out over the Atlantic Ocean waters during the second half of the 19<sup>th</sup> Century and first half of the 20<sup>th</sup> Century had relatively few chances to be observed and thus included into HURDAT. This is because, unlike today, the wide array of observing systems such as geostationary/polar orbiting satellites, aircraft reconnaissance, radars, and moored/drifting buoys were not available. Landsea (2007) provides an example of the typical distribution of marine observations available in the early 20<sup>th</sup> Century versus those that are taken today. Detection of tropical storms and hurricanes up until the mid-1940s was limited to those tropical storms and hurricanes that affected ships and those that impacted land. Until the utilization of two-way radio in the first decade of the 20<sup>th</sup> Century, the only way to obtain ship reports of hurricanes at sea was after the ships made their way back to port. Observations from these late ship reports were not of use to the fledgling weather services in the United States and Cuba operationally, though some of them were available for post-analyses of that season's TC activity. 1909 marked the first time that a ship reported a hurricane by radio in the Atlantic basin (Neumann et al. 1999). Despite the substantial increase in shipping traffic during the first few decades of the 20<sup>th</sup> Century, more widespread utilization of on-board barometers and the use of radio to both send and receive reports about these storms led to modest decreases in ship-based observations of TCs because of better knowledge of where the systems were occurring and where they would likely track. It is estimated that more than three tropical cyclones a year were likely missed in the pre-geostationary satellite era between 1900 and 1965 (Landsea 2007).

The bulk of the data utilized for the reanalysis efforts for the period 1911 to 1920 are ship observations from the *Historical Weather Map* (HWM) series, the Comprehensive Ocean-Atmosphere Data Set (COADS - Woodruff et al. 1987), the *Monthly Weather Review* (MWR), and miscellaneous ship reports obtained from the National Climatic Data Center. The HWM series, a reconstruction of daily surface northern hemispheric synoptic maps begun by the U.S. Navy and U.S. Weather Bureau in the 1920s, was conducted for the years of 1899 through 1969. While COADS is one of the most comprehensive observational ship database available and often contains most ship observations found in HWM, there are some data in HWM not available in COADS. The *Monthly Weather Review* regularly published an "Ocean Gales and Storms" section that had significant (gale force winds [34 kt] or greater) ship observations, which also were occasionally not found in COADS. Overall, for TCs over the open ocean, COADS provided the majority of relevant ship observations for the reanalyses. It is to be noted that COADS was not generally utilized in the reanalysis efforts for the period for 1851 to 1910 conducted by Partagas and Diaz (1996) and quality controlled/digitized by Landsea et al. (2004a).

Once a TC impacted land in the early 20<sup>th</sup> Century, then both station-based meteorological observations and more anecdotal reports become readily available. Station data are available from HWM, the U.S. Weather Bureau Original Monthly Records<sup>3</sup> (OMR), MWR, the Cuban meteorological journal *Reseña*, and original sources from the Mexican Weather Service. The MWR, in particular for the era of the 1910s, was quite detailed in providing many raw observations as well as providing descriptions of the impacts of the landfalling systems both in the United States and elsewhere in the Atlantic basin. MWR also routinely provided a graphic called Tracks of the Centers of Cyclones that was the first depiction of TC (and extratropical storm)

positions twice a day in the United States, northern Mexico, southern Canada, the Gulf of Mexico and the northwest Atlantic Ocean. Although this was a useful product, it was still often necessary to consult the original observations of the U.S. Weather Bureau found in the OMR reports to best estimate exact landfall position and intensity.

Other miscellaneous data sources that helped provide information on the track and intensity of existing TCs and helped identify previously overlooked systems included the following for the period of 1911 to 1920: Barnes (1998a, 1998b), Boose et al. (2001, 2002), Cline (1926), Connor (1956), Dunn and Miller (1960), Ellis (1988), Hall (1913), Ho et al. (1987), Hudgins (2000), Jarrell et al. (1992), Jarvinen et al. (1985), Kasper et al. (1998), Mitchell (1924), Neumann et al. (1999), Perez (1971), Perez et al. (2000), Rappaport and Partagas (1995), Roth (1997a, 1997b), Roth and Cobb (2001), Schwerdt et al. (1979), Tannehill (1938), Tucker (1982), Wiggert and Jarvinen (1986) and various newspaper accounts.

All available oceanic and coastal observations were then analyzed once daily (more frequently if the TC was over heavily trafficked shipping lanes or over land with more data being available) and the resulting estimated TC positions and intensities compared with the HWM, MWR, and original HURDAT tracks. Changes to the original HURDAT were made only if observations supported making substantial alterations to the track (generally at least  $0.3^{\circ}$  latitude-longitude) and intensity (generally at least 10 kt,  $1 \text{ kt} = 0.5144 \text{ ms}^{-1}$ ). Figure A1 in the Appendix provides an example of the synoptic analysis conducted for one day during Storm #2, 1919 (the Key West hurricane). Possible alterations considered for each storm were for genesis, duration of the system, intensity, and decay and/or transformation into an extratropical cyclone. (Subtropical storms, which are included into HURDAT beginning in 1968, are not a category explicitly used in

<sup>&</sup>lt;sup>3</sup> These data are available on-line through the National Climate Data Center's Climate Database Modernization

the reanalysis during the 1910s due to lack of information about thermodynamical structure in the vertical and convective organization. Some TCs of the 1910s, however, do appear because of their large size to have some subtropical cyclone characteristics and a few of these might have been subtropical storms. Such systems are noted as such in their metadata writeup.) All official revisions to HURDAT have been examined, commented upon, and approved by the NHC Best Track Change Committee.

# 3. Track Estimation and Errors:

TC positions were determined in this study primarily by wind direction observations from ships and coastal stations and secondarily by sea level pressure measurements, reports of damages from winds, storm tides, and fresh-water flooding. With these observations and the knowledge that the surface flow in a TC is relatively symmetric (i.e. circular flow with an inflow angle of 10-20°, Houston et al. 1999), a relatively reliable estimate of the center of the storm can be obtained from a few peripheral wind direction measurements (see Figure 2 from Landsea et al. 2004a). While geographical positions of TCs in HURDAT were estimated to the nearest 0.1° degrees latitude-longitude (~6 nmi, 1 nmi = 1.852 km), the average errors were typically much larger in the early 20th Century than this precision might imply. Holland (1981) demonstrated that even with the presence of numerous ships and buoys in the vicinity of a strong TC that was also being monitored by aircraft reconnaissance, there were substantial errors in estimating its exact center position from the ship and buoy data alone. Another complicating issue in utilizing ship observations from COADS is that most ships of the era provided position estimates to a resolution of 0.5 to1.0° latitude-longitude because of the imprecision in navigation at the time (Figure 1). Based upon

Program: http://www.ncdc.noaa.gov/oa/climate/cdmp/cdmp.html .

these considerations, storms documented over the open ocean during the period of 1911 to 1920 were estimated to have position errors that averaged 100 nmi, with ranges of 150 to 240 nmi errors being quite possible in data sparse regions of the Caribbean Sea and central North Atlantic Ocean (Table 1). This position error estimate is the same as the preceding 25 years despite increased shipping traffic, because of the increasing ability of ships at sea to steer clear of an encounter with a TC.

At landfall, knowledge of the location of the TC was generally more accurate, as long as the storm came ashore in a relatively populated region (Table 1). By the early part of the 20<sup>th</sup> Century most coastal locations along the Gulf of Mexico, Caribbean Sea and western North Atlantic were settled and thus impacts of TCs facilitated more accurate estimates of landfall positions. The main exception to this was along the Mexican coastline, where - because of the ongoing conflict later named the Mexican Revolution – there was substantially decreased meteorological monitoring from 1910 until 1920. Average errors for position at and after landfall from 1911 to 1920 were on the order of 60 nmi (110 km) with somewhat smaller values occurring over densely populated and meteorologically monitored locations like Puerto Rico and the U.S. mainland coast between Georgia and Maine.

### 4. Intensity Estimation and Errors:

In comparison with TC position and track, analysis of TC intensity is much less straightforward when analyzing cyclones from the first half of the 20<sup>th</sup> Century. Intensity, described as the maximum sustained 1-min surface (10 m) winds, is recorded at a resolution of 10 kt from 1851 to 1885 and 5 kt for the period of 1886 to date. The reanalysis of peak winds for the Atlantic basin TCs that occurred from 1911 to 1920 was based upon (in decreasing order of weighting) central pressure observations, in situ wind observations from anemometers, Beaufort wind estimates, peripheral pressure measurements, wind-caused damages along the coast, and storm tide. These various observations are similar to what were available for the first reanalyses conducted for the years 1851 to 1910, though the measurements from instruments become relatively more common during 1911 to 1920.

Sea level central pressure (eye) measurements can provide relatively reliable estimates of the maximum wind speeds in a TC in the absence of in situ observations of the peak wind strength. If central pressure is not available, it can be estimated from peripheral (eyewall or rainband) pressure measurements if accurate values of the radius of maximum wind (RMW) and environmental (or surrounding) sea level pressure can also be obtained. Typically, this was possible at landfall when the RMW was estimated by measuring the mean distance from the hurricane's track to the location of the peak storm surge and/or peak wind-caused damages. Central pressure can then be estimated from an empirical formula found in Schloemer (1954) and Ho (1989).

Once a central pressure has been estimated, maximum wind speeds can be obtained from a pressure-wind relationship. The current standard pressure-wind relationship for use in the Atlantic basin by NHC (OFCM 2005) is that developed by Dvorak (1984) [modified from earlier work by Kraft (1961)]. The earlier reanalysis work (Landsea et al. 2004a) developed new pressure-pressure relationships that were latitude dependent. The resultant pressure-wind relationships for the four regions of the Gulf of Mexico, southern latitudes (south of 25°N), subtropical latitudes (25-35°N), and northern latitudes (35-45°N) gave similar results to Dvorak (1984) for weaker TCs with relatively high pressures (> 980 mb), but differed significantly for stronger hurricanes. For example, for a central pressure of 960 mb, both the Gulf of Mexico and southern latitude

relationships would suggest a maximum wind of 100 kt, while the subtropical latitude relationship gives 94 kt and the northern latitudes only 90 kt. Compared to Dvorak (1984), the Gulf of Mexico and southern latitude relationships are most similar, while the subtropical and northern latitude relationships indicate significantly weaker winds than Dvorak. These latitudinally-based pressure-wind relationships from Landsea et al. (2004a) were utilized exclusively in the reanalysis for 1851 to 1910 and were the primary tool for 1911 to 1920.

A new set of pressure-wind relationships based upon data since 1998 were developed by Brown et al. (2006). While similar to Landsea et al. (2004a) for the southern and subtropical latitudes, Brown et al.'s association for the Gulf of Mexico suggest weaker winds for given pressures in the hurricane intensity range. They found no significant difference in the pressurewind relationship between those TCs in the Gulf of Mexico versus those over the Atlantic within the same latitude belt, which was in contrast to Landsea et al. Moreover, Brown et al. were also able to stratify by those TCs that are deepening and those that are filling. They did not have enough cases north of 35°N to evaluate the northern latitudes relationship. The Brown et al. revised relationships were utilized for Gulf of Mexico hurricanes for the period 1911 to 1920.

The use of pressure-wind relationships to estimate winds in TCs has a few associated caveats. First, for a given central pressure, a smaller-sized TC (measured either by RMW or radius of hurricane winds) will produce stronger winds than a large TC (Knaff and Zehr 2007). Vickery et al. (2001), building from earlier work by Ho et al. (1987), developed a statistical relationship between RMW and central pressure, environmental pressure, and latitude from hurricanes that made landfall in the continental United States. Tropical storms and hurricanes with observed/estimated RMWs that were smaller (larger) by 25-50% from the these climatological RMW values for their given central pressure, environmental pressure and latitude had wind speeds

increased (decreased) in the reanalysis work by about 5 kt above that suggested by the latitudinally-based pressure-wind relationships. TCs with RMW dramatically (more than 50%) different from climatology had winds adjusted by about 10 kt, accordingly. It is acknowledged that this is a somewhat arbitrary adjustment process, though there is not a straightforward alternative available.

Another caveat concerns the translational speed of the TC. In general, the translational speed is an additive factor on the right side of the storm and a negative factor on the left for northern hemisphere TCs (Callaghan and Smith 1998, Knaff and Zehr 2007). At low to medium translational speeds (less than around 20 kt), the variation in storm winds on opposite sides of the storm track is approximately twice the translational velocity, although there is substantial uncertainty and non-uniformity regarding this impact on TC winds. At faster translational speeds, this factor is somewhat less than two (Boose et al. 2001). Storms that move at least 50% faster than the regionally-dependent climatological translational speeds (Neumann 1993, Vickery et al. 2001) have been chosen in the reanalysis to have higher (5 kt) maximum wind speeds than slower storms with the same central pressure at that location. Similarly, storms with significantly slower than usual rates of translational velocity (> 50%) are given slightly reduced winds (5 kt) for a given central pressure.

One final adjustment to maximum winds provided by central pressure is based upon the environmental pressure. TC's embedded in higher (lower) than climatological environmental pressures will have stronger (weaker) pressure gradients and thus increased (decreased) winds, if all other factors are equal (Knaff and Zehr 2007). While the climatological pressures vary by month and location, in general, when environmental pressures are higher than about 1016 mb, 5 kt

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additional wind would be indicated over that suggested by the pressure-wind relationship, while pressures lower than about 1010 mb would suggest lowering the winds by about 5 kt.

For many early 20th Century storms, the central pressure could not be estimated from peripheral pressure measurements with the Schloemer equation because values for the RMW were unknown. However, one can get a wind from the peripheral pressure based upon the latitudinallybased pressure-wind relationships. This wind would represent a minimum estimate of what the strongest winds were at the time, given that the central pressure would be lower – perhaps by a few mb, perhaps by substantially more. In most of these cases, the best track winds chosen for the reanalysis were 5-10 kt higher than that suggested by the pressure-wind relationship itself.

For land based observations of wind speed, the period of the 1910s was just before a time of transition regarding the type of anemometer generally being utilized. The original four-cup anemometer, first developed by Robinson in the 1840s (Kinsman 1969), was still widely used in the United States and other countries until the 1920s. Its primary limitations were in calibrating the instrument and its mechanical failure in hurricane-force wind conditions. Even as late as the 1890s, the highest wind that could be reliably calibrated with this instrument was only about 30 kt (from a whirling machine – similar in structure to a record player), due to a lack of reliable comparisons with a known quantity of faster motion. By the early 1920s, wind tunnels allowed for calibration against much stronger winds. These showed that the winds from the early cup anemometers had a strong overestimation bias, which was most pronounced at hurricane-force wind speeds (Fergusson and Covert 1924). For example, when these instruments indicated winds of minimal hurricane-force of 64 kt, the true wind was only 50 kt. Moreover, most of these early four-cup anemometers were disabled or destroyed by the TC before sampling the highest winds.

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min sustained wind measurement of 100 kt in Storm #11, 1916 at Mobile, Alabama (Kadel 1926). (A standard of 5-min was typically utilized in U.S. Weather Bureau reports of "maximum winds" in the era, due to instrumental uncertainties in obtaining shorter time period winds.) With the availability of reliable calibrations beginning in the 1920s, the true velocity of this observation was determined to be only about 77 kt. Current understanding of gustiness in hurricane conditions suggest a boost of 1.06 to convert from a 5-min to a 1-min maximum sustained wind (Powell et al. 1996), giving a best estimate of the maximum 1-min sustained wind of about 82 kt. These older style anemometers were replaced by the more reliably calibrated three-cup anemometers during the mid and late 1920s (Fergusson and Covert 1924), though these new instruments still suffered from mechanical failure in extreme winds. These corrections were thus applied for the 1910s and had been previously incorporated into the 1851 to 1910 time period reanalysis efforts (Landsea et al. 2004a).

However, the bulk of wind speed observations in Atlantic basin TCs during 1911 to 1920 were those subjective determinations of oceanic winds using the Beaufort Scale. This scale was refined and promoted as a wind force scale for sailing ships by Admiral Francis Beaufort and required in all British Royal Navy log entries by 1838 (Kinsman 1969). Subsequently, the scale evolved into one associated with specific wind speed ranges as specified by interpretations of the sea state, rather than the wind's impact on a ship's sails as sailing ships were replaced by those with engines later in the 19<sup>th</sup> Century. Due to limitations at the top end of the Beaufort Scale, the COADS, HWM and other ship data sources of the time generally list reports of "hurricane" force winds as 70 kt winds. The listed wind speeds were boosted to 90 kt only when ship reports included terms such as "severe", "violent", "terrific", or "great hurricane". Hurricanes at sea were not reanalyzed with a best track intensity value of major hurricane (Saffir-Simpson Scale Category)

3, 4 or 5; 96 kt or greater maximum surface wind speeds) unless corresponding central/peripheral pressure data was able to confirm such an intensity. Caution was warranted in the direct use of these Beaufort Scale wind estimates for tropical storm and hurricane intensity assignments due to a lack of consistency and standardization in the scale during the early 20th Century (Cardone et al. 1990). However, in many cases, these Beaufort Scale measurements by mariners were the primary tools available for estimating the intensity of TCs in this era.

In the absence of instrumental observations of winds and pressure, one can utilize windcaused destruction and storm surge measurements to make estimates of intensity of TCs at landfall. Indeed, the work of Boose et al. (2001, 2004), which utilized wind-caused destruction in New England and Puerto Rico to assess hurricane impacts, favorably matched instrument-based assessments in Ho et al. (1987) and Ho (1989) and in the reanalysis work reported in Landsea et al. (2004a) for the period of 1851 to 1910. Such damage assessments can narrow down the uncertainty of intensity estimates for landfalling hurricanes in settled areas within about one category on the Saffir-Simpson Scale. However, wind-caused destruction alone is too complex to reliably estimate an exact maximum wind speed. In addition to maximum winds encountered, hurricane wind damage is also dependent on the duration of destructive winds, the wind steadiness (change of wind direction), the exposure, and the building materials, workmanship and building codes employed in the construction of the structures (Cochran 2000, Dunion et al. 2003). Thus wind-caused damage from hurricanes is only given a small weight in determining intensity at landfall.

Storm surge measurements can also assist in the determination of TC intensity at landfall, such as that listed in the Saffir-Simpson Hurricane Scale (Simpson 1974). However, such categorizations are only a rough estimate and are extremely variable because of several factors

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other than intensity: RMW, coastline shape, local offshore bathymetry/inland topography, astronomical tides, wave setup, and inflow angle (i.e., Jelesnianski et al. 1992). However, one can utilize several reliable storm surge measurements along with an accurate track of the landfalling hurricane in sensitivity tests using the Sea, Lake, and Overland Surges from Hurricanes (SLOSH – Jelesnianski et al. 1992) model to obtain a central pressure and RMW that produces the best fit of the simulated storm surge values to the observations. This has been done for several landfalling hurricanes, such as the 1915 Galveston hurricane (Storm #2, Wiggert and Jarvinen 1986) and the 1898 Brunswick hurricane (Storm #7, Sandrik and Jarvinen 1999). With these derived central pressure and RMW values, the maximum winds can then be straightforwardly estimated, but an isolated maximum storm surge value without the assistance of SLOSH modeling runs is of limited use in estimating landfall intensity.

Once the landfall intensity of a U.S. continental hurricane strike is determined, the spatial variations (what U.S. states or portions of states) are analyzed and compared with the existing classification in HURDAT. In addition to the previously mentioned factors that are utilized for determining maximum wind, a simple parametric wind model (Schwerdt et al. 1979) is employed to assist in the delineation of states impacted. This model, given inputs of TC position, translational speed and direction, maximum wind, RMW and location of interest, provides the approximate winds (marine exposure) for that location. A series of runs with the model can provide estimated peak sustained winds experienced at that location, which allows for an objective determination of Saffir-Simpson Scale categorization for places not directly impacted by the right front quadrant RMW where the peak wind typically resides. For example, the Key West major hurricane of 1919 (Storm #2, 1919 – see Appendix) was originally assessed to be a Category 4 for both southwest Florida (BFL4) at landfall in the Florida Keys and again when it reached south

coastal Texas (ATX4). After the reanalysis of meteorological data and applications of Schwerdt et al.'s model, it was determined that the conditions at the Keys landfall were unchanged (Category 4 for southwest Florida). However, the peripheral impacts were increased to include a Category 2 impact for southeast Florida (BFL2). Additionally, the peak impact in Texas was downgraded to Category 3 for south coastal Texas (ATX3), but central coastal Texas was also added as Category 3 (BTX3).

After landfall, existing HURDAT TC intensity estimates are problematic as mentioned earlier because of errors introduced by interpolation and the often unrealistic, complete lack of weakening when the systems were over peninsulas and large islands. Analyses of intensity in the decaying phase over land are primarily based upon observations of pressures and winds as well as models of pressure and wind decay for TCs described below. An observation of central pressure after landfall can be easily converted to an equivalent maximum wind with the appropriate pressure-wind relationship. However, these algorithms were derived assuming over-water conditions. The use of the associations for TCs overland must consider the increased roughness length of most land surfaces and the dampening of the maximum sustained wind speeds that result. In general, maximum sustained wind speeds over open terrain exposures (roughness lengths of 0.03 m) are about 5-10% slower than over-water wind speeds (Powell and Houston 1996), though for rougher terrain the wind speed decrease is substantially greater. Ho et al. (1987) developed several relationships for the decay of TC central pressure after landfall, which were stratified by geographic location and value of the pressure deficit (environmental pressure minus central pressure) at landfall. This pressure decay model can be utilized to estimate central pressure for a weakening system after landfall or to analyze the pressure at landfall given a central pressure reading well inland.

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Because of the mesoscale nature of TCs, even for systems that made landfall in a relatively data rich region like the United States, only rarely are there enough direct winds observations to reasonably insure that an actual measurement of maximum winds were made. The Kaplan and DeMaria (1995, 2001) inland wind decay model provided guidance for determining wind speeds after landfall of a TC. This model utilizes the maximum wind at landfall and provides decayed maximum wind speed values out to about two days after landfall. The decay of winds by the model over higher terrain areas such as Hispaniola and much of Mexico is inadequate (e.g., Bender et al. 1985). For these cases, a faster rate of decay than that given from the model (on the order of 30% accelerated rate of decay) was utilized in the re-analysis. The results from the Kaplan and DeMaria inland wind decay model were compared with available observations and only utilized when actual pressure and wind data were too sparse to adequately estimate the maximum wind from direct observations.

Original and reanalyzed best track intensity estimates for the 1910s were based mainly upon observations by ships at sea, which more often than not, would not sample the most intense part of the storm (typically only 30-60 nmi in diameter). Holland (1981) demonstrated that even in a relatively data-rich region of ship and buoy observations within the circulation of a TC, the actual intensity was likely to be substantially underestimated. Figures 3 and 4 from Landsea et al. (2004a) provided a graphic demonstration of this for Major Hurricane Erin of 2001 that made a close by-pass of Bermuda. Aircraft winds extrapolated to the ocean surface indicated maximum surface winds of about 100 kt in Erin at 1930 UTC on the 9<sup>th</sup> of September 2001. However, despite transiting within 85 nmi of Bermuda, the highest observed surface winds from ships and coastal stations were only around 40 kt. Such an underestimation of TC intensities was likely common in the pre-satellite and pre-aircraft reconnaissance era. It is estimated that the intensity measurements

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for 1911 to 1920 were in error an average of 20 kt over the open ocean, with a bias toward underestimating the true intensity (Table 1). These values are the same as the period of 1886 to 1910, but smaller than 1851 to 1885. For TCs landfalls during the 1910s, intensity estimates were improved and show a negligible bias as most coastlines around the western North Atlantic, Gulf of Mexico and Caribbean Sea were substantially settled by that time (Table 1). Again, these values are the same as the period of 1886 to 1910, but smaller than 1851 to 1885. A notable exception to this for landfalling TCs is for Mexico, due to the lack of meteorological monitoring during the Mexican Revolution of 1910 to 1920.

## 5. Criteria for Adding New or Removing Existing Tropical Cyclones:

Based upon examination of the *Historical Weather Maps*, monthly synoptic assessments contained in the *Monthly Weather Review*, the COADS ship database and other sources, potentially new TCs were considered for inclusion into the Atlantic hurricane database. The current definition of "tropical cyclone" utilized at the National Hurricane Center today is the following: "A warm-core non-frontal synoptic-scale cyclone, originating over tropical or subtropical waters, with organized deep convection and a closed surface wind circulation about a well-defined center." Given that only TCs of tropical storm intensity or greater are included into HURDAT, the definition of "tropical storm" is also relevant: "A tropical cyclone in which the maximum sustained surface wind speed (using the U.S. 1-minute average) ranges from 34 kt (39 mph or 63 km/hr) to 63 kt (73 mph or 118 km/hr)."

Systems were thus considered for inclusion into HURDAT during the era of the 1910s based upon the following criteria:

1. Non-frontal (not an extratropical cyclone);

2. Closed surface wind circulation;

3. At least two separate observations of sustained tropical storm force winds (at least 34 kt) or the equivalent in sea level pressure (roughly 1005 mb or lower). The two separate observations could come from the same ship/station or two different platforms.

Systems that could not unambiguously meet all of these criteria were not included into HURDAT, but were described in the metadata file as a possible.

On occasion, there were systems listed in HURDAT that appeared to not be TCs based upon today's definitions. However, only if it could be reasonable certain through sufficient observations that no tropical storm force winds were present at any point throughout the time that the system maintained a non-frontal, closed circulation structure, would a listed TC be considered for removal.

As with other changes in HURDAT, additions of new TCs and removal of existing TCs were officially decided by the NHC Best Track Change Committee.

### 6. Results:

#### a. Overall activity:

A summary of the yearly changes to HURDAT is provided in Figure 2 and Table 2. Figure 2 shows the revised and original track maps for the individual seasons from 1911 to 1920. It is apparent that most of the track changes introduced for these years are fairly minor (less than a 120 nmi alteration in position at anytime during the TC's lifetime) as easily seen in the case of the five original TCs in 1915, though examples can be seen of more dramatic alterations on occasion (e.g., old Storm #4/new Storm #6 in 1911, old Storm #6/new Storm #7 in 1912, Storm #3 in 1913).

Despite making relatively minor changes overall, nearly every existing TC was adjusted for at least some portion of its track.

In addition to track alterations of existing systems, new TCs were discovered and added into HURDAT for the first time and one existing system in HURDAT was reanalyzed to not be a tropical storm and thus removed from the database. In total, 13 new TCs had sufficient observational evidence to document their existence and were added into HURDAT: 2 in 1911, 1913, 1916 and 1919; 1 in 1912, 1915, 1917, 1918, and 1920; and no new systems in 1914 and 1916. Of these 13, four of the new TCs were landfalling systems: new Storm #6, 1913 in Cuba (as a hurricane); new Storm #1, 1916 in Florida; new Storm #5, 1916 in Mexico (possibly a hurricane); and new Storm #4, 1919 in Georgia. Additionally, one system during the 1910s in HURDAT was removed because of a lack of gale force winds (old Storm #8 in 1916). In other years in the reanalysis work (e.g., 1891), two separate TCs were found to be actually one continuous system and thus so changed to reflect this. There also has been a TC removed from HURDAT because the system was shown to be an extratropical cyclone throughout its lifetime (e.g., old Storm #5,1855). However, for the period of 1911 through 1920, no such TCs were identified.

Table 2 lists the original and revised tallies of tropical storms and hurricanes, hurricanes, major hurricanes, and Accumulated Cyclone Energy (ACE – an index for overall TC activity that takes into account the total frequency, intensity and duration of TCs, Bell et al. 2000). ACE is calculated by summing the squares of the estimated 6-hourly maximum wind speed in knots to be found in HURDAT for all periods while the system is either a tropical storm or hurricane.

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The average number of tropical storms and hurricanes increased from 4.9 per year in the original HURDAT to 6.1 after the reanalysis (Table 2). This net increase includes new systems that we added into the database as well as the one that was originally in HURDAT but was discarded. Both values are substantially below the long-term average of 11.1 per year recorded in the satellite era of 1966-2006 (Blake et al. 2007). The tropical storms and hurricanes that stayed out at sea for their duration and did not encounter ships will at this point remain undocumented for the time period of the 1910s. It is estimated that the number of undetected tropical storms and hurricanes for the 1911 to 1920 era is on the order of 3-4 per year (Table 1). While this is an improvement over the number missed in the first three and a half decades in HURDAT (4-6 per year during 1851 to 1885), it is the same estimate as the previous 25 years. This is the case despite the increased shipping traffic from 1911 to 1920, because of the better ability of mariners to avoid TC encounters with the more widespread employment of on-board barometers. (The use of twoway radios likely also contributed toward fewer encounters of ships with TCs, but presumably there had to be at least one encounter with the TC by a ship in order for other ships to avoid a known storm.) By no means should the TC record for the Atlantic basin as a whole be considered complete for either the frequency or intensity of tropical storms and hurricanes for the years 1851 to 1920. (These estimates of "missed" tropical storms and hurricanes are narrowed from that originally shown in Landsea et al. 2004, based upon the new work presented in Landsea 2007.)

In contrast, the hurricane, major hurricane and ACE averages (Table 2) show smaller changes in recorded values. Hurricane frequency had a small increase from 3.5 to 3.8 per year (6.2 per year in the modern era), major hurricanes remained unchanged at 1.3 per year (2.3 per year recently), and ACE dropped slightly from 61.1 to 58.7 per year (91.0 per year recently). The decrease in ACE is likely due to a systematic tendency for the original HURDAT to somewhat

overestimate the intensity of hurricanes from 1911 to 1920, especially over the open ocean (e.g., Storm #3, 1913, old Storm#9/new Storm #10, 1916). With regards to ACE, one year recorded a substantial increase in activity (ACE higher by at least 10.0 - 1918), two years saw a substantial decrease in activity (ACE lower by at least 10.0 – 1912 and 1916), and the remaining seven years had minor alterations in overall intensity, duration and frequency. Despite a moderate increase in the number of tropical storms and hurricanes because of the use of more data than were available to meteorologists of the era, the overall activity is slightly reduced by a modest amount because of the correction of an overestimation in intensity in the original HURDAT. In general, large changes to intensity (at least a 20 kt alteration at some point in the TC's lifetime) were recorded – both upward and downward – for the majority of individual TCs, typically with more significant changes than those introduced for track.

# b. Continental United States Hurricanes:

Table 3 summarizes the continental U.S. hurricanes for the period of 1911 to 1920 and the states impacted by these systems. U. S. hurricanes are defined as those hurricanes that are analyzed to cause maximum sustained (1 min) surface (10 m) winds of at least 64 kt for an open exposure on the coast or inland in the continental United States. Hurricanes that make a direct landfall with the circulation center (eye) of the system crossing the coast as well as those that make a close bypass are considered. In addition to the parameters also common to HURDAT (e.g. latitude, longitude, maximum winds and central pressure), the U.S. hurricane compilation also includes the outer closed isobar, the mean size of the outer closed isobar, and - when available - the RMW. These parameters provide information regarding the size of the hurricanes, which can vary

considerably from system to system. For these TCs, winds listed in HURDAT in the last six hourly period before landfall are now consistent with the assigned Saffir-Simpson Hurricane Scale category, which was not the case in the original HURDAT database before the reanalysis efforts. For most U.S. hurricanes of this era, a central pressure observation or estimate was obtained from original sources, Ho et al. (1987) or other references, which was then used to determine maximum wind speeds through the application of one of the new pressure-wind relationships. In the cases where there was no central pressure value directly available, the estimated winds at landfall were then used via the pressure-wind relationship to back out a reasonable central pressure. In either case, the objective was to provide both an estimate of the maximum wind and a central pressure at landfall for all U.S. hurricanes.

There were 19 U.S. hurricanes (seven that were major hurricanes) during the 1911 to 1920 period after the reanalysis. This represents one less hurricane than the original HURDAT database contained, with two new U.S. hurricane added (Storm #5, 1913 as a Category 1 in South Carolina and Storm #1, 1915 as a Category 1 in northeast Florida) and three U.S. hurricanes removed. Two of the three removed former U.S. hurricanes were analyzed instead to only be of tropical storm intensity at landfall (old Storm #2/new Storm #3, 1916 in Massachusetts and Storm #3, 1920 in North Carolina) and the other one was analyzed instead to be an extratropical cyclone by landfall (old Storm #14/new Storm #15, 1916 in southwest Florida). No major hurricanes were either added or removed from the U.S. hurricane list.

Notable hurricanes that affected the continental United States for 1911 through 1920 (Blake et al. 2007) after reanalysis include Storm #2, 1915 Category 4 in north Texas; old Storm #5/new Storm #6, 1915 Category 3 in Louisiana; old Storm #4/new Storm #6, 1916 Category 4 in south Texas; and Storm #2, 1919 Category 4 in south Florida and Category 3 in south Texas.

During the period of 1911 to 1920, the first very destructive hurricane to strike the continental United States was Storm #2, 1915, which hit the north Texas coast near Galveston, killed about 275 people, and would cause on the order of \$71 billion in total damages if the same system made landfall today (Blake et al. 2007). This TC was originally listed as a Category 4 for the north Texas coast with a 945 mb central pressure at landfall. The revised central pressure of a deeper 940 mb along with a large RMW of 25 nmi suggests winds of about 115 kt, which supports a Category 4 status. Also in 1915, old Storm #5/new Storm #6 struck Louisiana south of New Orleans, killed about 275 people and was originally listed as a 931 mb Category 4 hurricane at landfall. The reanalysis raised the central pressure upward to 944 mb, which along with a large RMW of 26 nmi suggests winds of about 110 kt necessitating a reduction to a Category 3 at landfall. In the following year of 1916, old Storm #4/new Storm #6 made landfall in along the south Texas coast as a Category 3 hurricane with a central pressure of 948 mb in HURDAT originally. The reanalysis of this system gave a deeper central pressure of 932 mb and a very large RMW of 40 nmi, suggesting winds at landfall of 105 kt, retaining this hurricane as a Category 3 in south Texas. In 1919, Storm #2 hit the Florida Keys and south Texas as a Category 4 hurricane in both locations originally, killing 287 people, and causing about \$14 billion in damages if the same system were to hit today. The Florida Keys landfall retained the 927 mb central pressure in HURDAT and along with the moderately sized 15 nmi RMW gave winds of 130 kt, keeping the system as a Category 4 at that location. However, in south Texas, the hurricane is reanalyzed to have had a central pressure of 950 mb, a large 35 nmi RMW and a low environmental pressure of 1006 mb giving winds of about 100 kt, and it was downgraded to a Category 3 for this second U.S. landfall.

Summarizing, there were only two sizable alterations for U.S. major hurricanes in the reanalysis (Table 3): old Storm #5/new Storm #6, 1915 was revised downward from a Category 4 to a Category 3 in Louisiana; and Storm #2, 1919 was decreased from a Category 4 to a Category 3 in southern Texas (though Category 4 was retained for the Florida Keys).

### c. Major Hurricanes outside of the Continental U.S.:

Outside of the continental United States, major hurricanes impacted only a few locations from 1911 to 1920. Three separate major hurricanes made landfall either in the Lesser Antilles, Greater Antilles or Bermuda. Of note was that all of Central America including all of the east coast of Mexico was spared from any direct strikes by major hurricanes during this time period. However, the 1910s also corresponded with the Mexican Revolution, so monitoring of the weather and particularly of hurricanes in Mexico was incomplete during this time and it is possible that a major hurricane may have been misclassified as a minor hurricane along Mexico's Gulf coast.

Two of the more noteworthy major hurricane impacts for 1911 to 1920 were the following (Rappaport and Partagas 1995, Pielke et al. 2003, Blake et al. 2007): old Storm #6/new Storm #7, 1912 in Jamaica that killed 200 people; and old Storm #3/new Storm #4, 1917 in Cuba (known as "Nueva Gerona"). Both of these had substantial changes to their intensity, though only the track of the 1912 hurricane had major alterations. The 1912 hurricane in Jamaica was originally assessed to be a 130 kt at landfall, but was downgraded to a 100 kt Category 3 based upon a 965 mb central pressure at landfall and small RMW. (Most damage from this slow moving hurricane was rainfall-produced flash flooding, which has a weak relationship to intensity of the system.) The 1917 Nueva Gerona hurricane in Cuba was revised upward from 100 kt up to a 120 kt Category 4

hurricane with a 928 mb central pressure at landfall. Overall of the three major hurricane strikes listed in Table 4, one (1917's Nueva Gerona in Cuba) had a substantial increase in listed intensity and one (1912's Jamaican hurricane) had a sizable reduction in intensity at landfall.

## 7. Summary and Future Work:

Historical TC reconstructions are inevitably subject to revisions whenever new archived information is uncovered or when new analysis techniques are devised. Thus, while a couple thousand alterations and additions to HURDAT have been completed for the years 1911 to 1920, this does not insure that there may not be further changes once new information or revised physical understanding is made available. Such an archive of historical data – especially one based upon quasi-objective interpretations of limited observations of a mesocale feature like a TCs intensity – should always be one that can be revised when more data or better interpretations of existing information becomes available. A key to the analyses conducted here is that all of the raw meteorological observations – in addition to the smoothed "best track" revisions – are made available for the first time<sup>4</sup>. This allows users to inspect the changes made to TCs of interest, see the observations that the changes are based upon, and come to differing conclusions if warranted.

Highlights of accomplishments attained for this stage of the Atlantic hurricane database reanalysis project for 1911 to 1920:

A) Track alterations were implemented for most TCs in the existing HURDAT, though the majority were for minor changes;

<sup>&</sup>lt;sup>4</sup> All raw observations, revised HURDAT, annual track maps, metadata regarding changes for individual tropical cyclones, and comments from/replies to the National Hurricane Center's Best Track Change Committee can be found at : http://www.aoml.noaa.gov/hrd/data\_sub/re\_anal.html .

- B) Intensity changes were incorporated into nearly all TCs with a much larger proportion with major alterations in their intensity, either toward stronger or weaker winds;
- C) 13 new TCs were discovered and added into HURDAT, while 1 system was removed from the database because it was not of tropical storm intensity;
- D) While the frequency of tropical storms during the era was increased from 4.9 to
  6.1 annually because of these net changes, the overall effect of track and
  intensity alterations was to produce slightly less activity during the era than
  existed originally because of a small overestimation bias in the intensity of some
  existing TCs;
- E) 19 continental U.S. hurricanes were identified, one less than originally listed in HURDAT because of an addition of two new U.S. hurricanes and the removal of three hurricanes during the time period. No changes were made to the number of major continental U.S. hurricanes, though two Category 4 U.S. hurricanes were reclassified as a Category 3 strike;
- F) Only three major hurricanes struck other places in the Atlantic basin Cayman Islands, Cuba and Jamaica. Of these, one had a substantial increase in intensity and one was sizably reduced in intensity at landfall;
- G) Despite the reanalysis changes, there exists significant uncertainty in TC tracks, significant undercounts in TC frequency, and significant underestimation of TC intensity, especially for those systems over the open ocean.

However, much more work still needs to be accomplished for the Atlantic hurricane database. One essential project is a Partagas and Diaz (1996) style reanalysis for the years before

1851. This may lead to a complete dataset of U.S. landfalling hurricanes for the Atlantic coast from Georgia to New England back to at least 1800, given the relatively high density of population extending that far into the past. While the reanalysis efforts thus far have extended HURDAT back to 1851 and revised it through 1920, these did not make extensive use of COADS until the decade of the 1910s (Landsea et al. 2004a). Further improvements in HURDAT could be achieved by utilizing this massive ship database for the years of 1851 to 1910. An ongoing project is to complete the current reanalysis efforts through the remainder of the 20<sup>th</sup> Century. Beginning in 1944, the Atlantic TC database incorporates aircraft reconnaissance data. Already, methodologies have been established on how to objectively reanalyze TCs with highly detailed aircraft reconnaissance observations (Dunion et al. 2003, Landsea et al. 2004b). Additionally, new techniques for utilizing pressure-wind relationships in the context of global reanalysis datasets are also emerging (e.g., Knaff and Zehr 2007). Work to complete the Atlantic hurricane basin database reanalysis is crucial because of current important questions that are being raised about anthropogenic climate change on TC activity (International Workshop on Tropical Cyclones 2006).

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# Appendix

All Atlantic basin tropical storms and hurricanes in the new best track database are accompanied by a "metadata file". This file consists of a day-by-day listing of peak meteorological observations and previous estimates of the storm's position and intensity. The metadata also contains a descriptive paragraph about the particular methodology employed for making changes in the genesis, track, intensity and decay of that TC including what sources were crucial for revising the best track, whether or not a wind-pressure relationship was utilized, if wind decay models were used for inland wind estimates, and any other pertinent information. All of the tropical storms and hurricanes for the period of 1911 to 1920 are considered "UNNAMED". However, many of these storms have been recognized by various informal names. These are included in the metadata file when at all possible. The following is an example of a single metadata entry for Storm #2, 1919 – the Key West Hurricane. Table A1 provides significant (near hurricane force and greater) reports collected for this system and made available in the raw

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database. Figure A1 provides a single, daily analysis of the synoptic observations available at 12 UTC, 10 September, 1919.

### Storm #2, 1919 (The Key West Hurricane)

Major changes to the track and intensity shown in Neumann et al. (1999). Evidence for these alterations comes from the Historical Weather Map series, the COADS ship database, *Monthly Weather Review*, the Original Monthly Records from the National Climatic Data Center, Connor (1956), Dunn and Miller (1960), Schwerdt et al. (1979), Jarvinen et al. (1985), Ho et al. (1987), Jarrell et al. (1992), and Perez et al. (2000).

September 1: HWM and COADS observations possibly indicate a wave approaching the Lesser Antilles without any indication of a closed low (though data are sparse east of the islands). No gale force winds (or equivalent in pressure) were observed.

September 2: HWM indicates a closed low of at most 1010 mb at 13.5N, 64W. HURDAT lists this system as a tropical storm at 15.4N, 63.5W at 12 UTC. The MWR "Summary of the Hurricanes of 1919, 1920, and 1921" does not begin the system until either late on the 2nd or early on the 3rd. Available observations suggest that the cyclone was substantially east-northeast of HURDAT's postion. No gale force winds (or equivalent in pressure) were observed. "The minor disturbance ... was first noted on the evening of September 2 ... a little west of the island of Antigua" (MWR).

September 3: HWM indicates a closed low of at most 1010 mb at 16N, 66W. HURDAT lists this system as a tropical storm at 17N, 67W at 12 UTC. The MWR "Summary of the Hurricanes of 1919, 1920, and 1921" shows a center near 18N, 65W. Available observations suggest a center east-northeast of HURDAT's estimate. No gale force winds (or equivalent in pressure) were observed. "This ... minor disturbance moved west-northwestward at about a normal rate, passing near the southern portion of the island of Porto Rico" (MWR).

September 4: HWM indicates a closed low of at most 1010 mb at 20N, 70W. HURDAT lists this system as a tropical storm at 19.2N, 69W at 12 UTC. The MWR "Summary of the Hurricanes of 1919, 1920, and 1921" shows a center at 19N, 69.5W. Available observations suggest a center between all three estimates. No gale force winds (or equivalent in pressure) were observed.

September 5: HWM indicates a closed low of at most 1010 mb at 21N, 73W. HURDAT lists this system as a Category 1 hurricane at 21N, 71.8W at 12 UTC. The MWR "Summary of the Hurricanes of 1919, 1920, and 1921" shows a center near 20.5N, 72W. Available observations suggest a center west of the MWR Summary estimate. No gale force winds (or equivalent in pressure) were observed. "By the evening of the 4th it had reached the north coast of the island of Santo Domingo with a barometer reading of about 29.80 inches. On the morning of the 5th the center of the disturbance was approximately 100 miles southwest of Turks Island with about the same barometric pressure" (MWR).

September 6: HWM indicates a closed low of at most 1010 mb at 21.5N, 72.5W. HURDAT lists this system as a Category 1 hurricane at 22.2N, 72.4W at 12 UTC. The MWR "Summary of the Hurricanes of 1919, 1920, and 1921" shows a center near 21.5N, 72.5W. Available observations suggest a center west of the HURDAT estimate. No gale force winds (or equivalent in pressure) were observed. "By the evening of the 5th the winds at Turks Island had changed from east to west, and were southerly over Santo Domingo and Haiti, still light in character, apparent evidence that the disturbance had recurved to the northeastward during the day, and that it was moving in that direction in very moderate form" (MWR).

September 7: HWM indicates a closed low of at most 1010 mb at 22N, 73.5W. HURDAT lists this system as a Category 2 hurricane at 23.4N, 74.1W at 12 UTC. The MWR "Summary of the Hurricanes of 1919, 1920, and 1921" shows a center near 23N, 73.5W. The MWR Tracks of Lows shows a center near 23N, 74.5W with 1003 mb (a.m). Available observations suggest a center just southwest of HURDAT's estimate. Ship highlight: 35 kt SE and 1009 mb at 26N, 74.4W at 23 UTC (COA). "On the evening of the 6th pressure and wind conditions over Santo Domingo and the Bahamas indicated the possible presence of a disturbance over the eastern Bahamas. Conditions were slightly more pronounced on the morning of the 7th ... there were slight indications of a disturbance over the central Bahamas" (MWR).

September 8: HWM indicates a closed low of at most 1005 mb at 21.5N, 76W. HURDAT lists this system as a Category 3 major hurricane at 23.9N, 77W at 12 UTC. The MWR "Summary of the Hurricanes of 1919, 1920, and 1921" shows a center near 23.5N, 76W. The MWR Tracks of Lows shows a center near 23.5N, 77W with 998 mb (a.m.). Available observations suggest a position just southwest of HURDAT's estimate. Ship highlight: 35 kt NNE and 1006 mb at 25.5N, 80.5W at 21 UTC (COA). Station highlight: 51 kt NE and 998 mb at Nassau at 01 UTC (MWR) "A belated report on September 8 that a severe storm could be located south of and near the Andros Islands" (MWR).

September 9: HWM indicates a closed low of at most 1000 mb at 23.5N, 81.5W. HURDAT lists this system as a Category 3 hurricane at 24N, 79.8W at 12 UTC. The MWR "Summary of the Hurricanes of 1919, 1920, and 1921" shows a center near 24N, 79.5W. The MWR Tracks of Lows shows a center near 24N, 79.5W. Available observations suggest that the center was between the HWM and HURDAT estimates. Ship highlights: 70 kt N and 938 mb at 24.6N, 82.9W at 21 UTC (MWR). Station highlights: 50 kt NE and 986 mb at Key West at 23 UTC (MWR); 57 kt NE at Sand Key at 1748 UTC (MWR). "Considerable local damage was done in Miami and vicinity, although nothing very serious resulted. Tides were unusually high and many small boats suffered. The greatest loss was probably in the fruit crop ... Press reports indicated that considerable damage was also done along the northwest coast of Cuba" (MWR) ... The greatest [shipping loss was] the Spanish steamship Valbanera, off Rebecca Shoals Light, about 40 miles west of Key West. The vessel arrived off Morro Castle, Habana, on September 9, but owing to the hurricane, was unable to enter the harbor, and nothing further was heard from her until a diver discovered her beneath the waters off Rebecca Shoals. The Valbanera was from Spanish ports for New Orleans, via Habana, and her 400 passengers and crew of 88 must have perished" (MWR). "El Huracan del Valvanera - Category 1 in Cuba - September 9 and 10" (Perez et al.).

September 10: HWM indicates a closed low of at most 995 mb at 24N, 82W. HURDAT lists this system as a Category 4 hurricane at 24.6N, 82.7W at 12 UTC. The MWR "Summary of the Hurricanes of 1919, 1920, and 1921" shows a center near 24.5N, 83W. The MWR Tracks of Lows shows a center near 24.5N, 83W. Available observations suggest that the center is west-northwest of HURDAT's estimate. Ship highlights: 927 mb (eye?) at 24.6N, 82.9W (MWR); 930 mb (eye?) at 24.6N, 82.9W at 05 UTC (MWR). Station highlights: 937 mb at Rebecca Shoals Light; 932 mb (eye) at Dry Tortugas; 82 kt NE at 0148 UTC and 960 mb at 0510 UTC at Sand Key. "The storm center passed about 30 or 40 miles south of Key West about midnight of September 9. At this time the barometer at Key West read 28.83 inches with an east wind of an estimated velocity of 105 miles an hour, which increased slightly during the next hour. At Sand Key, the lowest barometer at about the same time was 28.35 inches, a difference of 0.48 inch within a distance of 8 miles ... The following report on the storm at Key West and vicinity was prepared by Mr. H. B. Boyer, official in charge of the Weather Bureau office at that place: `The storm that passed over Key West on September 9 and 10 was, without question, the most violent experienced since records at this station began. While the minimum barometric reading, 28.81 inches, was not as low as that recorded in 1909 (28.52) and in 1910 (28.47), the violence of the wind was undoubtedly greater. It is to be regretted that owing to the vibrations of the tower supporting the wind instruments the anemometer cups were shaken loose and blown away at 7:30 pm on the 9th in gusts ranging between 75 and 80 miles an hour, and thereafter until 3:35 p.m. of the 10th the wind-velocity record was lost. The wind-vane was blown away at 12:45 a.m. of the 10th during the winds of greatest intensity ... In the terrific gusts that prevailed during the height of the storm stanch brick structures had walls blown out and large vessels, firmly secured, were torn from their fastenings or moorings and blown on the bank ... the great loss, estimated at \$2,000,000 ... Owing to the very slow progressive movement of the storm in this vicinity, winds of gale force and over lasted continuously from about 7 a.m. of the 9th to about 9:30 p.m. of the 10th ... From the

forenoon of the 9th squalls of wind and rain progressively increased in force and frequency, culminating in terrific gusts of great violence between midnight of the 9th and 2 a.m. of the 10th ... Probably not a structure on the island escaped being damaged more or less ... three lives were lost by drowning' ... The report of the storm experiences at Sand Key, Fla., was prepared by Mr. Eugene M. Barto, observer, and is as follows: `The record showed that the anemometer cups blew away at 9:35 p.m. with a wind velocity of 84 miles an hour. The wind vane was probably blown away shortly after midnight. This was also the time of the lowest barograph record, which was 28.35 ... The highest [wind] recorded was 94 miles an hour from the northeast at 8:39 p.m.' ... The center of the storm passed directly over Dry Tortugas, 65 miles west of Key West, with a reported barometer reading of 27.51 inches, while at Rebecca Shoals Light, about 40 miles west of Key West, the lowest reading was 27.66 inches ... The steamship Winona went ashore at 10 a.m., September 10, on a reef on the northeast portion of the Tortugas group, near Pulaski Shoals ... the barometer [fell at midnight on the 9th] to 27.45 inches ... A later report from the tank steamer, Fred W. Weller, showed a barometer reading of 27.36 inches in the vicinity of Dry Tortugas on September 9 ... These [close readings] within a very limited area, make it safe to assume that they were substantially correct" (MWR). "September 10, 1919, 929 mb Central Pressure, 24.6N, 82.9W Landfall Point, 15 nmi Radius of Maximum Wind" (Ho et al.) "1008 mb environmental pressure, 115 kt maximum 1 min surface wind" (Schwerdt et al.) "Tropical

Cyclones in Florida, September 9-10, Key West, Major, Marine casualties 300 plus" (Dunn and Miller). "Saffir-Simpson Category 4 for FL Keys/S TX with 927 mb central pressure" (presumably for FL landfall) (Jarrell et al.)

September 11: HWM indicates a closed low of at most 995 mb at 25.5N, 87W. HURDAT lists this system as a Category 4 hurricane at 25.6N, 84.7W at 12 UTC. The MWR "Summary of the Hurricanes of 1919, 1920, and 1921" shows a center near 25.5N, 86W. The MWR Tracks of Lows shows a center near 26N, 85.5W. Available observations suggest that the center is southwest of the HURDAT estimate. Ship highlights: 45 kt SSE and 998 mb at 26.6N, 85.8W at 23 UTC (COA). "[One the 11th], the tide reached a crest of 5.55 feet above low-water mark, 2 feet higher than ever before recorded in the annals of the United States Engineers. The tide did some little damage along that section of the coast, but none of consequence" (MWR).

September 12: HWM indicates a closed low of at most 995 mb at 27N, 89W. HURDAT lists this system as a Category 4 hurricane at 26.7N, 88W at 12 UTC. The MWR "Summary of the Hurricanes of 1919, 1920, and 1921" shows a center near 26.5N, 88W. The MWR Tracks of Lows shows a center near 26.5N, 88W. Available observations suggest that the center is just south of the MWR Summary and Tracks estimates. Ship highlights: 944 mb (eye?) at 26.2N, 87.8W at 14 UTC (MWR); 948 mb at 27N, 89W at 22 UTC (MWR); 942 mb (eye?) at 27N, 88.5W at 23 UTC (MWR). "After the morning of the 10th, at which time the storm center was apparently very near Dry Tortugas, Fla., its path could only be approximated. It happened, however, that a report received by mail from the steamship Lake Deval nearly two weeks after the storm located the center with a fair degree of definiteness on the morning of the 12th [about 150 miles south-southeast of the mouth of the Mississippi River]" (MWR).

September 13: HWM indicates a closed low of at most 995 mb at 27N, 92.5W. HURDAT lists this system as a Category 4 hurricane at 26.5N, 91.6W at 12 UTC. The MWR "Summary of the Hurricanes of 1919, 1920, and 1921" shows a center near 26N, 91W. The MWR Tracks of Lows shows a center near 26.5N, 91W. Available observations suggest that the center is west of the HURDAT estimate. Station highlights: 36 kt SE and 1002 mb at Burrwood at 12 UTC (MWR). Ship highlights: 931 mb (eye) at 26.5N, 90.5W at 05 UTC (MWR). "The tide was 6 feet above normal on Lake Borgne and on Grand Isle, and 5 to 6 feet above normal on Lake Ponchartrain, on the afternoon of the 13th ... By a little after sunset the tide [at Port Aransas] had reached 5 feet above mean sea level" (MWR).

September 14: HWM indicates a closed low of at most 995 mb at 27.5N, 96.5W. HURDAT lists this system as a Category 2 hurricane offshore Texas at 27N, 95.7W at 12 UTC. The MWR "Summary of the Hurricanes of 1919, 1920, and 1921" shows a center near 27N, 95.5W. The MWR Tracks of Lows shows a center near 27N, 96W. Available observations suggest that HURDAT's estimate is most accurate. Station highlights: 61 kt at 18 UTC and 970 mb at 21 UTC at Corpus Christi (OMR). Ship highlights: 950 mb (eye?) at 27N, 95W at 14 UTC (MWR). "On the morning of September 14 the storm center was not far from the coast of Texas, between Corpus Christi and Brownsville, and during the day in passed inland, with marked although with steadily diminishing intensity ... The tide ... reached its highest point of 8.8 feet [at Galveston] at 7 a.m. of the 14th. Two men lost their lives in the storm in this immediate vicinity [Galveston] ... both men were apparently overtaken by the rising tide and drowned. ... From reports received the height of the tide accompanying the storm ranged in this district from about 4 feet at Orange, Tex., to approximately 13 feet at Port O'Connor, Tex. With this tide and the high wind accompanying it, some damage resulted at many points, especially along the water front. At Seabrook, Tex., there were a few buildings, mostly light structures, destroyed ... At points to the south of Galveston, however, there was more damage done ... At Matagorda, Palacios, and Port Lavaca, Tex., there was considerable damage to wharves, fish houses, and small boats. Similar damage resulted at Port O'Connor, Tex. ... Stretching along the beach [of Corpus Christi] for 23 blocks homes were crushed and hurled away or wrecked by the tidal wave, which reached a depth of 15 feet in some places. Over much of the beach section not an indication of former homes now remains, except here and there a bathtub or part of a brick chimney ... In the

downtown [Corpus Christi] district utter demolition of some of the city's most important industrial and public plants marked an area extending for six blocks along the water front and more than a block in width, while beyond that block, extending back toward the bluff section, every commercial establishment's first floor was wrecked, and in some cases the entire building rendered useless, over a corresponding area two blocks wide. The tremendous property damage is becoming daily more apparent and prominent business men and other trained observers predicted to-night [Sep. 18] that \$20,000,000 would be a conservative estimate of the monetary loss in Corpus Christi. 284 bodies, almost entirely those of Corpus Christi victims, have been found ... Details of conditions at Port Aransas and other parts of the islands between Corpus Christi Bay and the Gulf were ascertained ... The docks and buildings in Port Aransas have been wiped out with the exception of a school building ... The large oil tanks there also were destroyed. The five who lost their lives [at Port Aransas] were drowned while attempting to leave the island in a lifeboat ... The Gulf storm caused a 6-foot tide here [Anahuac, Mexico], but Anahuac is situated on a 25-foot bank of Trinity Bay, hence no damage was done. The wind reached a velocity of perhaps 30 miles. ... The storm was only the second September storm of this character of any consequence that reached the south Texas coast during the last 45 years, the other having occurred in 1910. The storm of 1919 was by far the more violent of the two, and was probably the greatest of all Gulf storms ... The full force of the storm was experienced between

Aransas Bay and the mouth of the Rio Grande, where the high tides resulted in a toll of 183 dead and 174 missing" (MWR). "Sep. 14, Estimated Lowest Pressure 27.36" [for Dry Tortugas on the 10th], Tide Info - Corpus Christi 16', Galveston 8.8', Aransas Pass 11.5', Brownsville 3.6', Port Isabel 8', Sabine 8', Anahuac 10', La Porte 8.5', Carancahua 13', Ingleside 12', Velasco 10', Port O'Connor 13' " (Connor). "Sep. 14, Landfall point of 27.2N, 97.3W, 950 mb Central Pressure, 35 nmi Radius of Maximum Wind" (Ho et al.). "1007 mb environmental pressure" (Schwerdt et al.) "Tropical Cyclones in Texas, Sep. 14, Corpus Christi, Extreme, 300-600 killed, damage \$20,270,000" (Dunn and Miller). "Saffir-Simpson Category 4 in FL Keys/S TX with 927 mb central pressure" (presumably for FL, not TX) (Jarrell et al.) "Landfall around 18 UTC on the 14th, 950 mb central pressure, 35 nmi radius of maximum wind, 1010 mb ambient pressure, assumed that central pressure filled from 931 mb to 950 mb the six hours before landfall, after landfall analyzed 977 mb around 00 UTC on the 15th (Jarvinen et al.)

September 15: HWM indicates a closed low of at most 990 mb at 28N, 100.5W. The MWR Tracks of Lows shows a center near 28N, 100.5W. HURDAT lists this system as a tropical depression at 28.2N, 100.2W at 12 UTC. Available observations suggest a center west of the HURDAT estimate. Station highlights: 49 kt E at 14 UTC and 993 mb at 1140 UTC at Del Rio (OMR). September 16: HWM does not analyze a closed low, though a weak center is near 31.5N 106W. No gale force winds (or equivalent in pressure) were observed.

Genesis for this tropical cyclone was delayed by 12 hours consistent with the poorly organized circulation exhibited by numerous observations on the 2nd at 12 UTC. Minor changes to the track were made on most days in accordance with available observations. The exception was the 2nd where a major shift to the east-northeast was introduced. Decay of the tropical cyclone was delayed a day to account for a more intense system still in existence on the 15th as well as a weak vortex apparent from observations on the 16th. Intensity from the 2nd to the 6th reduced significantly based upon available observations, which also agrees with the *Monthly Weather Review* analyses of a weak tropical cyclone during these dates. Hurricane intensity is analyzed to have been attained on the 7th (two days later than originally shown in HURDAT). A 998 mb peripheral pressure with 51 kt winds from Nassau at 01 UTC on the 8th suggests winds of at least 51 kt from the southern pressure-wind relationship - 95 kt retained in HURDAT, as it appears that Nassau was on the outskirts of a large hurricane. Winds are also retained from 00 to 12 UTC on the 8th as the cyclone became a major hurricane. A 938 mb peripheral pressure (not eye) at 21 UTC on the 9th suggests winds of at least 120 kt from the southern pressure-wind relationship - 125 kt chosen for HURDAT (up from 110 kt originally). Three eye pressure measurements were

observed near Dry Tortugas, Florida early on the 10th: 927, 930 and 932 mb. 927 mb was selected by Jarrell et al. and is retained here for HURDAT, which suggests 129 kt from the southern pressure-wind relationship. Ho's estimate of an RMW of 15 nmi is quite close to the 14 nmi for climatology for this central pressure and latitude (Vickery et al.). Thus 130 kt is chosen for HURDAT at 06 UTC on the 10th, up from 115 kt originally. This retains the Category 4 assessment for the Florida Keys. Because of the revised definitions of the boundary between southwest and southeast Florida (BFL and CFL, accordingly) and through an application of the simplified wind model in Schwerdt et al., Category 2 conditions are estimated to have occurred in the Upper Keys and thus southeast Florida (CFL2). As is typical, anemometers at Key West and Sand Key were rendered inoperable before the passage of peak winds and these only recorded at most Category 1 conditions.

Three low pressure readings were observed from ships on the 12th - 944 mb at 14 UTC, 948 mb at 21 UTC, and 942 mb at 22 UTC. It is likely that the 944 and 942 mb values were central pressure readings and these are included as such into HURDAT. 944 mb and 942 mb suggest winds of 118 and 116 kt, respectively, from the Gulf of Mexico pressure-wind relationship. The new Brown et al. (2006) north of 25N pressure-wind relationship gives winds of 111 and 113 kt, respectively. 110 kt is chosen for HURDAT late on the 12th and early on the 13th based upon these observations. However, an eye reading of 931 mb was measured by ship on 04 UTC of the next day on the 13th. This

value suggests winds of 128 kt from the Gulf of Mexico pressure-wind relationship. The new Brown et al. (2006) pressure-wind relationship for north of 25N suggests winds of 123 kt. 125 kt is chosen for HURDAT, up from 115 kt originally at 06 UTC on the 13th.

The hurricane weakened significantly before landfall in Texas. A likely central pressure reading of 950 mb on 15 UTC on the 14th suggests winds of 110 kt from the Gulf of Mexico pressure-wind relationship. Both Ho et al. and Jarvinen et al. accepted this value as a likely landfall pressure along with an RMW of about 35 nmi. It is to be noted that the 950 mb central pressure and 35 nmi RMW values provide a good match in SLOSH model runs against observed storm surge measurements (Jarvinen et al.) Climatological RMW for this latitude of landfall and central pressure is substantially smaller - 18 nmi (Vickery et al.). This would suggest that the maximum sustained winds were about 100 kt both at 15 UTC at the ship report and at about 21 UTC at landfall in Texas. The new Brown et al. (2006) pressure-wind relationship for north of 25N filling cyclones also analyzes about 101 kt. 100 kt at landfall represents a reduction in the analyzed Saffir-Simpson Category assigned to south Texas from a 4 down to a 3 (ATX3). However, the wind speed in HURDAT at 18 UTC on the 14th right before landfall is adjusted upward sharply from 75 to 100 kt in the reanalysis. Application of the Schwerdt et al. idealized hurricane wind profile suggests that central Texas (BTX) should also be considered a Category 3 impact (BTX3), which is reasonable given

the landfall position was very close to the boundary between south and central Texas coast. Peak observed winds after landfall (within plus/minus two hours of synoptic times) were 34 kt at San Antonio at 00 UTC on the 15th, 44 kt at San Antonio at 06 UTC, and 49 kt at Del Rio at 12 UTC. (These convert to 29, 37, and 41 kt, respectively, after accounting for the high bias of the anemometer used and adjusting to a peak 1 min wind from these peak 5 min values [Fergusson and Covert 1924 and Powell et al. 1996]). However, with the landfall between Corpus Christi and Brownsville and with the anemometer at Corpus Christi becoming inoperable after 17 UTC, higher winds were quite likely present at 00 and 06 UTC on the 15th. A run of the Kaplan and DeMaria (1995) inland wind decay model suggests winds of 71, 49, and 35 kt, for the same synoptic periods. Given the low bias of the Kaplan and DeMaria model for the 12 UTC time, winds after landfall are chosen to be somewhat higher than the model: 75, 55 and 40 kt, respectively.

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Figure 1: Ship location accuracy example from COADS database. The red line with arrows misleadingly suggests a zig-zagged ship track according to COADS. Times of observations are given in parenthesis. Sea level pressure (in mb) and wind barbs are provided. The resolution of ship observations in COADS during early in the 20<sup>th</sup> Century is typically given in 1.0 to 0.5° latitude-longitude increments, which contributes towards uncertainty in the location of TCs. Figure 2: The revised (top) and original (bottom) Atlantic basin TC track map for 1911, 1912, 1913, 1914, 1915, 1916, 1917, 1918, 1919, and 1920.

Figure A1: Synoptic analysis for Storm #2, 1919 (the Key West Hurricane) at 12 UTC, September 10<sup>th</sup>, 1919. Observations of wind (full barb is 10 kt) and sea level pressure from ship and weather stations are provided. The track of the hurricane is given in blue with revised positions and maximum winds every six hours.

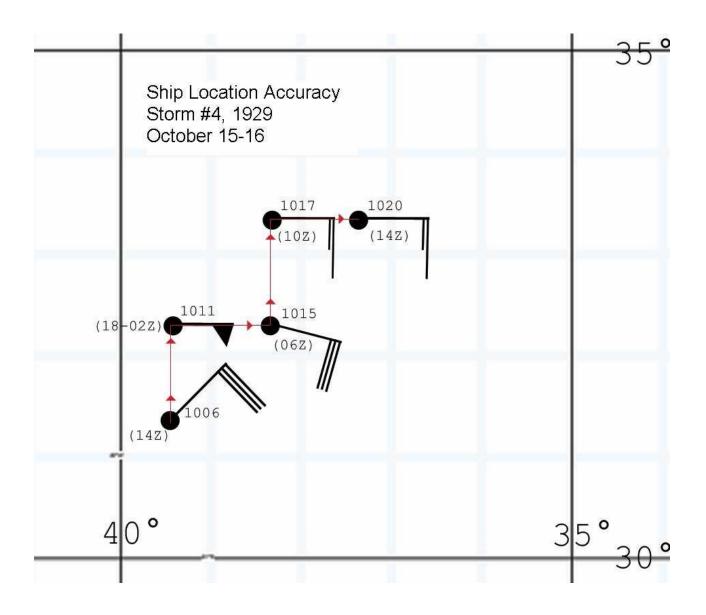
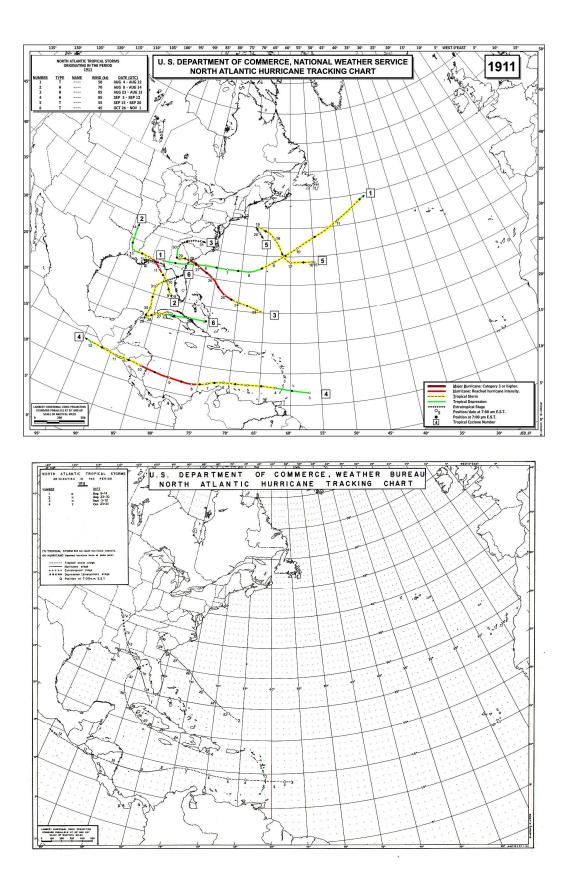
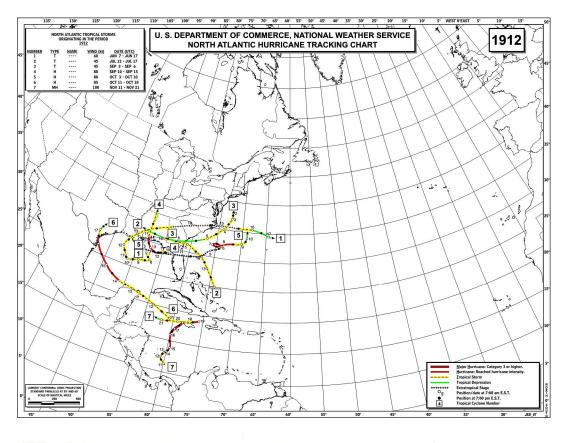
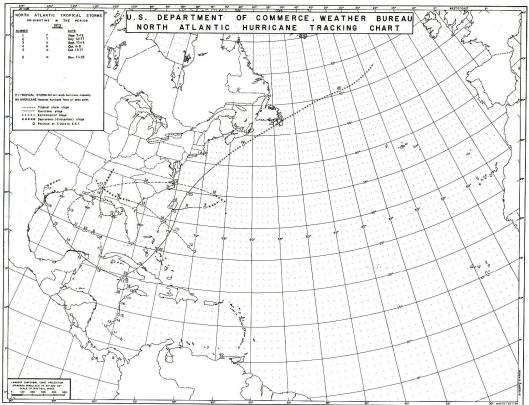
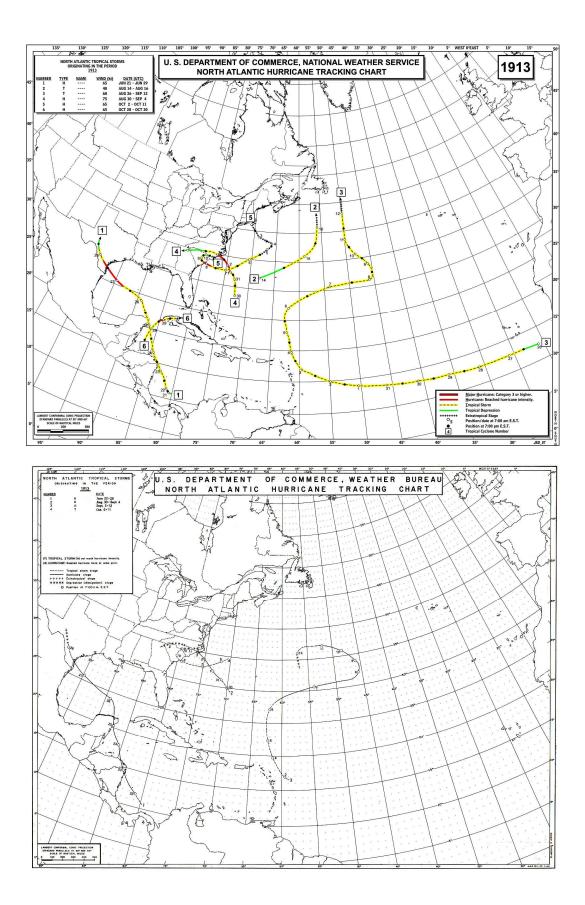


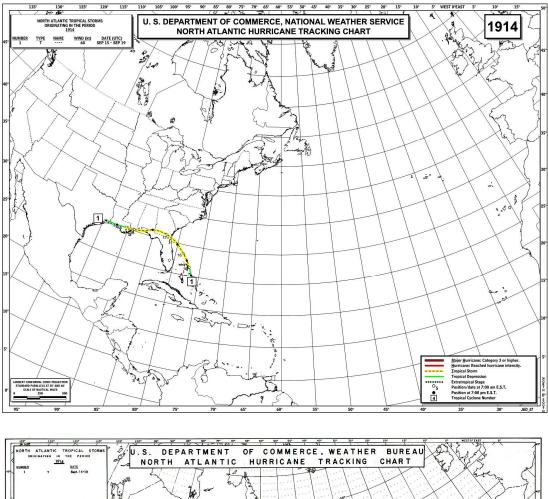
Figure 1: Ship location accuracy example from COADS database. The red line with arrows misleadingly suggests a zig-zagged ship track according to COADS. Times of observations are given in parenthesis. Sea level pressure (in mb) and wind barbs are provided. The resolution of ship observations in COADS during early in the 20<sup>th</sup> Century is typically given in 0.5 to 1.0<sup>o</sup> latitude-longitude increments, which contributes towards uncertainty in the location of TCs.

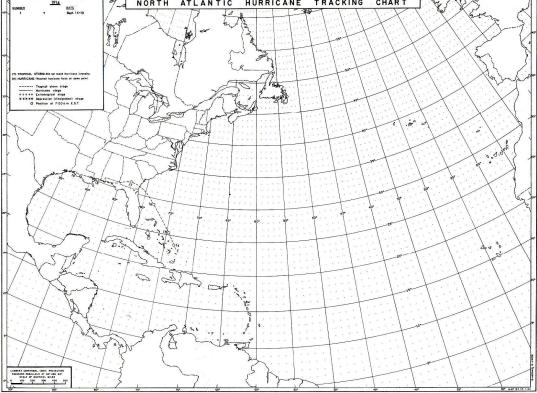


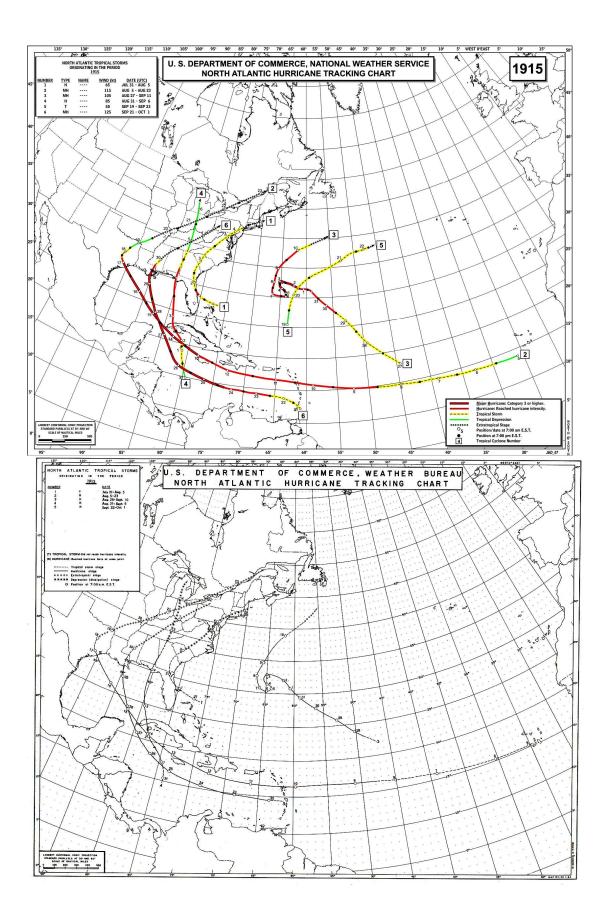


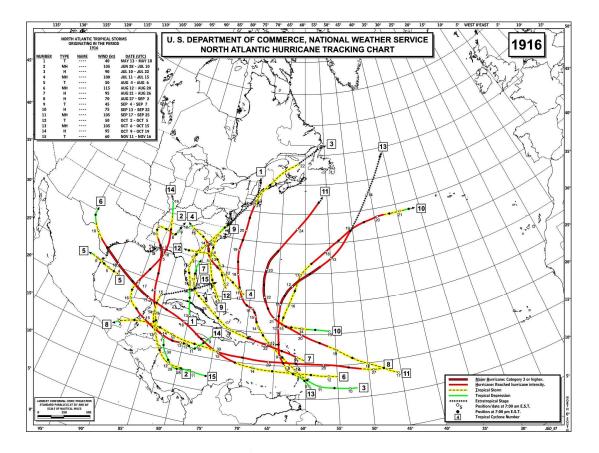


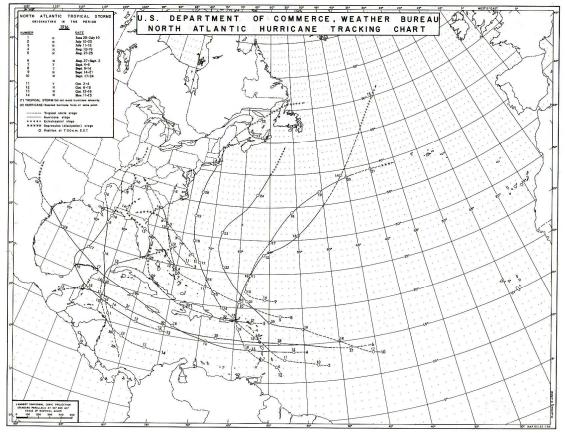


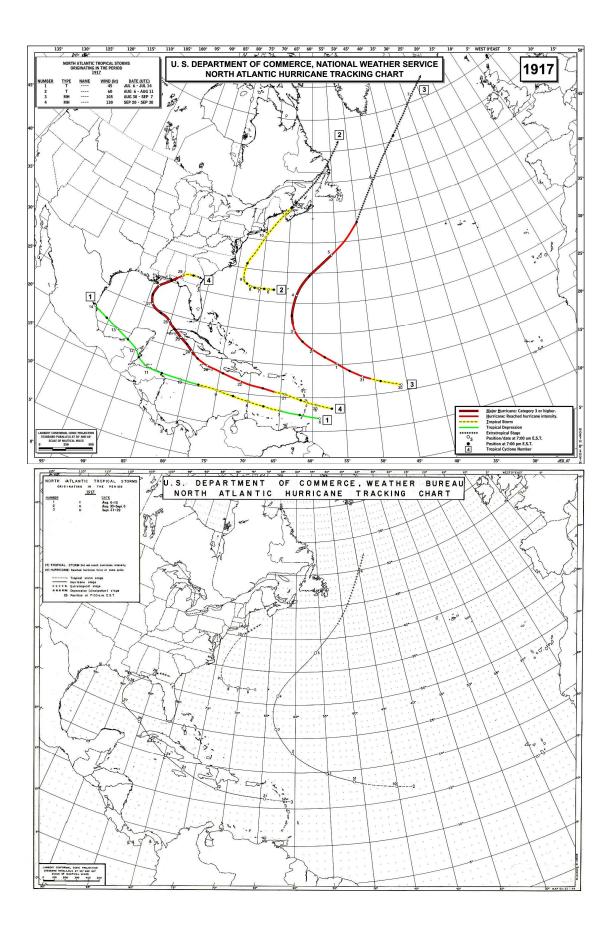


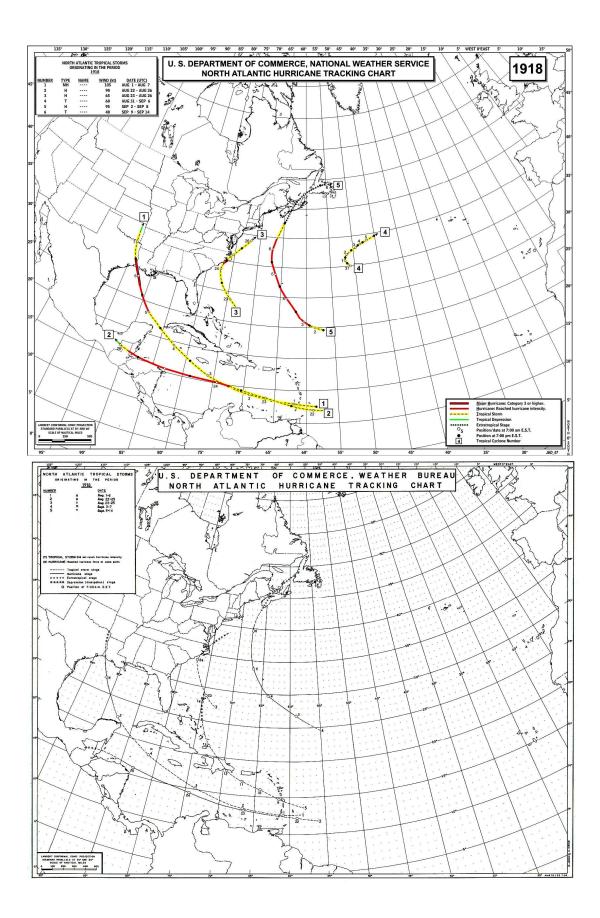


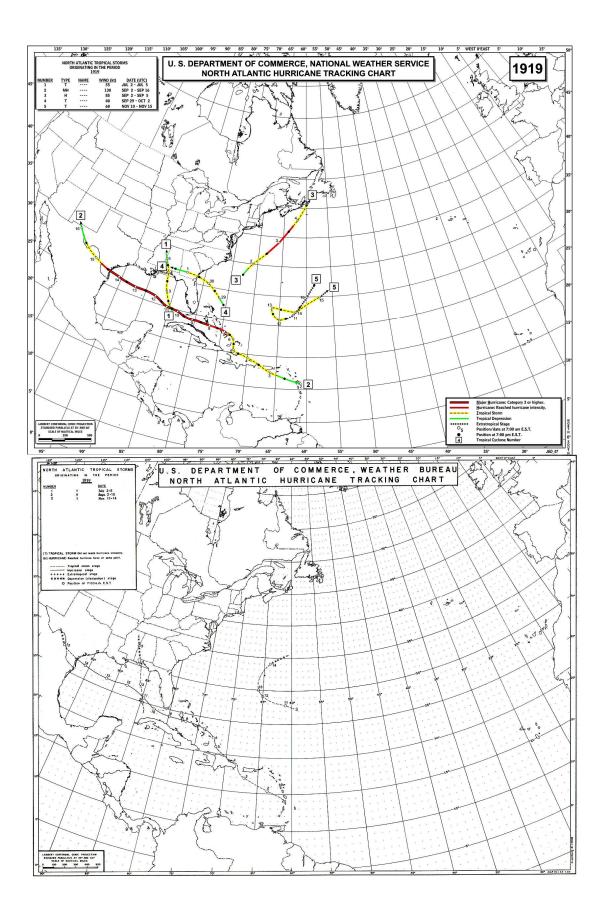












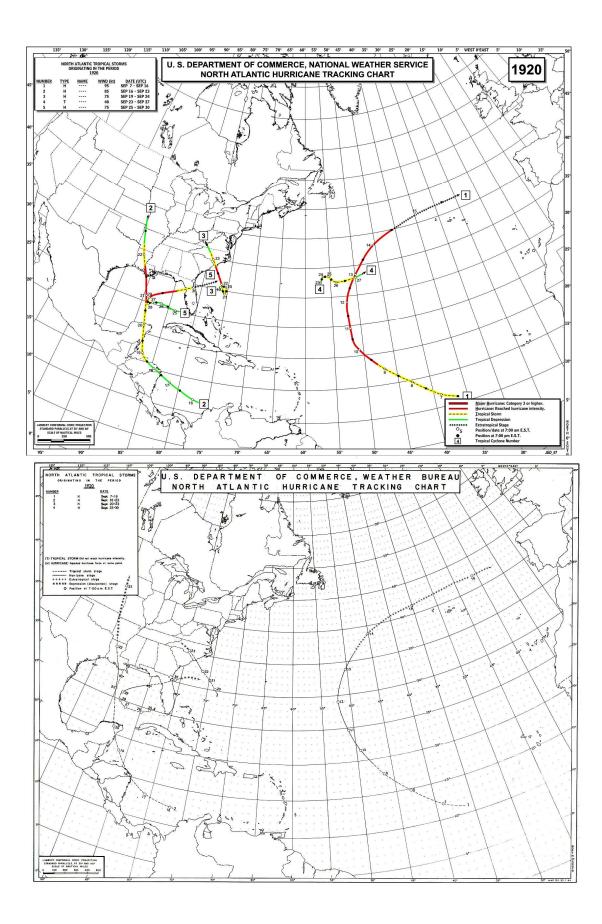


Figure 2: The revised (top) and original (bottom) Atlantic basin TC track map for 1911, 1912, 1913, 1914, 1915, 1916, 1917, 1918, 1919, and 1920.

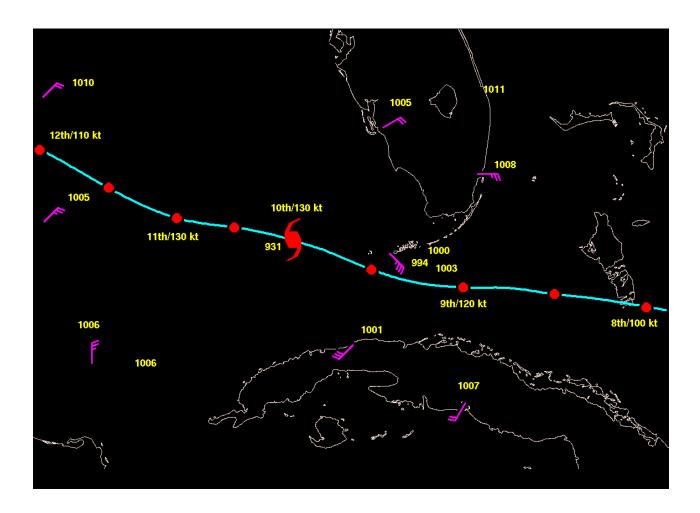


Figure A1: Synoptic analysis for Storm #2, 1919 (the Key West Hurricane) at 12 UTC, September 10<sup>th</sup>, 1919. Observations of wind (full barb is 10 kt) and sea level pressure from ship and weather stations are provided. The track of the hurricane is given in blue with revised positions and maximum winds every six hours.

Table 1: Estimated average position and intensity errors and frequency undercounts in the revised best track for the years 1851-1920. Negative bias errors indicate an underestimation of the true intensity. By 1920, only a few coastal areas in the Atlantic basin remained sparsely populated (i.e., less than two people per square mile), though some coastal regions (such as in Mexico due to the ongoing Mexican Revolution) were not well monitored. The Tropical Storm and Hurricane Undercount refer to annual numbers of systems that likely were not observed based upon density of ship traffic across the Atlantic basin.

			Intensity Error	Intensity Error	Tropical Storm
Situation	Dates	Position Error	-	-	& Hurricane
			(absolute)	(bias)	Undercount
Open ocean	1851-1885	120 nmi	25 kt	-15 kt	4 to 6
	1886-1920	100	20	-10	3 to 4
Landfall at sparsely	1851-1885	120	25	-15	1 to 2
populated area	1886-1920	100	20	-10	0 to 1
Landfall at settled	1851-1885	60	15	0	0
Area	1886-1920	60	12	0	0

	Tropical		Major		
Year	Storms &	Hurricanes	Hurricanes	ACE	
	Hurricanes		Trufficulies		
1911	4/6	3/3	0/0	36/35	
1912	6/7	4/4	1/1	74/56	
1913	4/6	3/4	0/0	43/36	
1914	1/1	0/0	0/0	3/3	
1915	5/6	4/5	3/3	118/127	
1916	14/15	11/10	6/5	177/144	
1917	3/4	2/2	2/2	52/61	
1918	5/6	3/4	0/1	29/40	
1919	3/5	1/2	1/1	48/55	
1920	4/5	4/4	0/0	31/30	
Average 1911-1920	4.9/6.1	3.5/3.8	1.3/1.3	61.1/58.7	
Average 1966-2006	11.1	6.2	2.3	91.0	

Table 2: Original/revised tropical storm and hurricane, hurricane, major hurricane and Accumulated Cyclone Energy (ACE) counts. ACE is expressed in units of  $10^4$  kt<sup>2</sup>.

Table 3: Continental U.S. Hurricanes: 1911 to 1920. Date/Time is day and time when the circulation center crosses the U.S. coastline (including barrier islands). Time is estimated to the nearest one hour. Lat/Lon is location estimated to the nearest 0.1 degrees latitude and longitude. Max Winds are the estimated maximum sustained (1 min) surface (10 m) winds to occur along the U. S. coast. Saffir-Simpson is the estimated Saffir-Simpson Hurricane Scale at landfall based upon maximum sustained surface winds. RMW is the radius of maximum winds (primarily for the right front quadrant of the hurricane), if available. Central Pressure is the minimum central pressure of the hurricane at landfall. Central pressure values in parentheses indicate that the value is a simple estimation (based upon a pressure-wind relationship), not directly measured or calculated. OCI is the sea level pressure at the outer limits of the hurricane circulation as determined by analysis of the outer closed isobar (in increments of 1 mb). Size is the quadrant-averaged radius of the OCI (in increments of 25 nmi). States Affected is the impact of the hurricane upon individual U.S. states by Saffir-Simpson Scale (again through the estimate of the maximum sustained surface winds at each state)<sup>5</sup>. Original Assessment is the Saffir-Simpson categorization by states originally provided in HURDAT.

#/Date	Time	Lat	Lon	Max S	Saffi	r- RMW	Central	OCI	Size	States	Original
			L	Winds S	Simps	on	Pressure			Affected	Assessment
2-8/11/1911	2200Z	30.3N	87.5W	70kt	1		(985mb)	1013mb	250nmi	AFL1,AL1	AFL,AL1
3-8/28/1911	0900Z	32.2N	80.7W	85kt	2	27nmi	972mb	1014mb	225nmi	SC2,GA1	SC2,GA2
4-9/14/1912	0800Z	30.3N	88.4W	65kt	1	50nmi	(988mb)	1007mb	150nmi	AL1,AFL1	AL1

<sup>&</sup>lt;sup>5</sup> ATX-South Texas, BTX-Central Texas, CTX-North Texas, LA-Louisiana, MS-Mississippi, AL-Alabama, AFL-Northwest Florida, BFL-Southwest Florida, CFL-Southeast Florida, DFL-Northeast Florida, GA-Georgia, SC-South Carolina, NC-North Carolina, VA-Virginia, MD-Maryland, DE-Delware, NJ-New Jersey, NY-New York, PA-Pennsylvania, CT-Connecticut, RI-Rhode Island, MA-Massachusetts, NH-New Hampshire, ME-Maine. In Texas, south is roughly from the Mexico border to Corpus Christi; central is from north of Corpus Christi to Matagorda Bay and north is from Matagorda Bay to the Louisiana border. In Florida, the north-south dividing line is from Cape Canaveral [28.45N] to Tarpon Springs [28.17N]. The dividing line between west-east Florida goes from 82.69W at the north Florida border with Georgia, to Lake Okeechobee and due south along longitude 80.85W.

6-10/16/1912	1800Z 27.1N	97.4W	85kt	2		(973mb)	1012mb	250nmi	ATX2	ATX1
1-6/28/1913	0100Z 27.1N	97.4W	65kt	1		(988mb)	1009mb	200nmi	ATX1	ATX1
4-9/3/1913	0800Z 34.7N	76.6W	75kt	1	38nmi	976mb	1016mb	200nmi	NC1	NC1
5-10/8/1913	1400Z 33.1N	79.4W	65kt	1		(989mb)	1012mb	150nmi	SC1	TS
1914 - None										
1-8/1/1915	1800Z 28.7N	80.8W	65kt	1	15nmi	990mb	1015mb	175nmi	DFL1	TS
2-8/17/1915	0700Z 29.2N	95.1W	115kt	4	25nmi	940mb	1009mb	325nmi	CTX4,BTX1,LA1	CTX4
4-9/4/1915	1100Z 30.0N	85.4W	80kt	1	25nmi	982mb	1012mb	225nmi	AFL1	AFL1
6-9/29/1915	1800Z 29.1N	90.3W	110kt	3	20nmi	944mb	1009mb	300nmi	LA3,MS2	LA4
2-7/5/1916	2100Z 30.4N	88.4W	105kt	3	26nmi	950mb	1008mb	250nmi	MS3,AL2,AFL2	MS3,AL3
4-7/14/1916	0800Z 32.9N	79.5W	95kt	2	20nmi	960mb	1013mb	175nmi	SC2	SC1
6-8/18/1916	2200Z 27.0N	97.4W	105kt	4	40nmi	932mb	1012mb	250nmi	ATX3	ATX3
14-10/18/1916	1400Z 30.4N	87.4W	95kt	2	19nmi	970mb	1010mb	325nmi	AL2,AFL2	AL2,AFL2
4-9/29/1917	0200Z 30.4N	86.6W	100kt	3	40nmi	949mb	1011mb	250nmi	AFL3,LA2,AL1	AFL3
1-8/6/1918	1800Z 29.8N	93.2W	110kt	3	12nmi	(955mb)	1012mb	150nmi	LA3,CTX1	LA3
2-9/10/1919	0700Z 24.6N	82.9W	130kt	4	15nmi	927mb	1009mb	275nmi	BFL4,CFL2	BFL4
2-9/14/1919	2100Z 27.2N	97.3W	100kt	3	35nmi	950mb	1006mb	250nmi	ATX3,BTX3	ATX4
2-9/22/1920	0100Z 29.1N	90.8W	85kt	2	28nmi	975mb	1009mb	250nmi	LA2	LA2

Table 4: Major Landfalling (non-continental U.S) Hurricanes: 1911 to 1920. The names listed are unofficial ones that the hurricanes are known by at these locations. Max Winds are the estimated maximum sustained (1 min) surface (10 m) winds to occur at along the coast at landfall/closest approach. Saffir-Simpson is the estimated Saffir-Simpson Hurricane Scale at landfall based upon maximum sustained surface winds. Central Pressure is the minimum central pressure of the hurricane at landfall/closest approach. Central pressure values in parentheses indicate that the value is a simple estimation (based upon a pressure-wind relationship), not directly measured or calculated. Original Winds are the winds in HURDAT that were originally provided at landfall/closest approach.

#/Date	Name	Location	Max	Saffir	Central	Original
			Winds	Simpson	Pressure	e Winds
1911 - None						
7-11/18/1912		Jamaica	100kt	3	965mb	130kt
1913 - None						
1914 - None						
2-8/13/1915		Cayman Islands	100kt	3	(960mb)	100kt
2-8/14/1915		Cuba	105kt	3	(955mb)	105kt
1916 - None						
4-9/25/1917	Neuva Gerona	Cuba	130kt	4	928mb	100kt
1918 - None						
1919 - None						
1920 - None						

Table A1: Significant (near hurricane force and greater) reports collected in the database for Storm #2, 1919 (the Key West Hurricane). Note that the complete database includes all reports of gales force (34 kt) or stronger and 1005 mb pressures or lower. Sources shown here are Monthly Weather Review (MWR) and Comprehensive Ocean Atmosphere Dataset (COADS). Notes

include the ship name, minimum pressure and maximum winds, if known.

Day	OBS	PRES	WIND	DIR	LOCATION	LAT	LON	SOURCE	SHIP/COMMENTS
1919 STO		pt 960			SHIP	040	700	MWR	Considere
9-Sep	3Z		70	Ν		240 246	790	MWR	Corydon Winona
9-Sep	20Z	938	70	IN	SHIP		829		winona
10-Sep		937			REBECCA SHOALS	245	825	MWR	Г.v.e
10-Sep		932			DRY TORTUGAS	246	829	MWR	Eye
10-Sep	07	927			SHIP	246	829	MWR	Fred W. Weller - Eye
10-Sep	0Z	933	0.4		SHIP	246	829	MWR	Winona
10-Sep	0Z	984	61	NE	KEY WEST	245	818	MWR	Max-W (no further obs)
10-Sep	0048Z	000	59	NE	SAND KEY	245	819	MWR	Max-W (no further obs)
10-Sep	1Z	982		NE	KEY WEST	245	818	MWR	
10-Sep	2Z	981		NE	KEY WEST	245	818	MWR	
10-Sep	3Z	980		NE	KEY WEST	245	818	MWR	
10-Sep	4Z	930			SHIP	246	829	MWR	Winona - Eye
10-Sep	4Z	976		NE	KEY WEST	245	818	MWR	Min-P
10-Sep	0410Z	960		SE	SAND KEY	245	819	MWR	Min-P
10-Sep	5Z	935		NW	SHIP	246	829	MWR	Winona
10-Sep	5Z	979		Е	KEY WEST	245	818	MWR	
10-Sep	6Z	981		Е	KEY WEST	245	818	MWR	
10-Sep	7Z	983		Е	KEY WEST	245	818	MWR	
10-Sep	8Z	933		NW	SHIP	246	829	MWR	Winona
10-Sep	8Z	984		Е	KEY WEST	245	818	MWR	
10-Sep	12Z	931		SE	SHIP	246	829	MWR	Winona
10-Sep	14Z	933		SE	SHIP	246	829	MWR	Winona
10-Sep	18Z	941		NW	SHIP	246	829	MWR	Winona
11-Sep	1Z	941			SHIP	246	829	MWR	Winona
11-Sep	7Z	945		SE	SHIP	246	829	MWR	Winona
11-Sep	14Z	947		SE	SHIP	246	829	MWR	Winona
11-Sep	19Z	962		SE	SHIP	246	829	MWR	Winona
12-Sep	3Z	963		SE	SHIP	246	829	MWR	Winona
12-Sep	13Z	944			SHIP	262	878	MWR	Lake Deval - Eye?
12-Sep	20Z	948			SHIP	270	890	MWR	Lake Grandon
12-Sep	21Z	942			SHIP	270	885	MWR	Tegulcigalpa - Eye?
13-Sep	4Z	931			SHIP	265	905	MWR	Berwyn - Eye
14-Sep	13Z	950			SHIP	270	950	MWR	F.R. Kellogg - Eye
14-Sep	16Z	982	59	Ν	CORPUS CHRISTI	278	975	OMR	
14-Sep	17Z		61		CORPUS CHRISTI	278	975	OMR	Max-W (no further obs)
14-Sep	20Z	970			CORPUS CHRISTI	278	975	MWR	Min-P
15-Sep	0Z	985		Е	CORPUS CHRISTI	278	975	OMR	
1-									