Probability of U.S. regional tornado outbreaks and their links to the springtime ENSO phases and North Atlantic SST variability

By Sang-Ki Lee

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25	This study shows that the four main phases of springtime El Niño-Southern Oscillation
26	(ENSO) evolution (persistent versus early-terminating El Niño, and resurgent versus
27	transitioning La Niña) are linked to distinctive spatial patterns of the probability of U.S. regional
28	tornado outbreaks. In particular, the outbreak probability increases significantly up to 27% over
29	the Ohio Valley, Upper Midwest and Southeast when a La Niña persists into the spring and is
30	followed by another La Niña (i.e., resurgent La Niña). The probability also increases
31	significantly up to 38%, but mainly in the South, when a two-year La Niña transitions to an E
32	Niño (i.e., transitioning La Niña). This study also shows that the North Atlantic sea surface
33	temperature tripole is linked to the tornado outbreak probability over the Southeast and Upper
34	Midwest in early spring, although this link may not be completely distinct from ENSO. These
35	changes in outbreak probability are shown to be largely consistent with remotely forced regional
36	changes in the large-scale tropospheric circulation, low-level wind shear, moisture transports and
37	extratropical storm activity.
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Abstract

1. Introduction

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48 The latest U.S. Natural Hazard Statistics reported that during 2004-2013 tornadoes claimed 49 1,091 lives in the U.S., only trailing behind heat-related fatalities, and caused 21.6 billion dollars 50 in property and crop damages (Table S1). An early prognosis of tornadogenesis combined with 51 an effective warning system will prevent many deaths and serious injuries. A parallel effort to 52 expand the current severe weather outlooks beyond seven days will also help emergency 53 managers, government officials, businesses and the public to better prepare their resources to 54 save lives and to protect critical infrastructure. 55 As summarized in a recent review [Tippet et al., 2014], notable advances have been made 56 since 2011, the year of the record-breaking tornado outbreaks in the U.S., on the potential of 57 expanding the severe weather outlook at the National Oceanic and Atmospheric Administration 58 (NOAA) beyond weather time scales [Tippett et al., 2012; Weaver et al., 2012; Barrett and 59 Gensini, 2013; Lee et al., 2013; Thompson and Roundy, 2013; Elsner and Widen, 2014; Allan et al., 2015]. Among others, Lee et al. [2013] showed that the majority of the extreme U.S. tornado 60 61 outbreaks in the most intense tornado months of April and May during 1950-2010 were linked to 62 a positive Trans-Niño (i.e., a positive zonal gradient of sea surface temperature anomalies (SSTAs) from the central tropical Pacific (CP) to the eastern tropical Pacific (EP)), typically 63 64 occurring during the boreal spring following the peak of La Niña [Trenberth and Stepaniak, 65 2001; Lee et al., 2014a]. They showed using observations and modeling experiments that a positive Trans-Niño could enhance large-scale atmosphere conditions conducive to intense 66 67 tornado outbreaks over the U.S. via extratropical teleconnections. Recently, Allen et al. [2015] 68 showed that La Niña events persisting into the spring could increase U.S. tornado activity, 69 especially over Oklahoma, Arkansas and northern Texas, and vice versa for El Niño events

70 persisting into the spring. They used March-April-May (MAM) SSTAs in the Nino3.4 region 71 (120°W-180°W and 5°S-5°N) to identify El Niño-Southern Oscillation (ENSO) events persisting into the spring. Krishnamurthy et al. [2015] further showed that the seasonal phasing of ENSO is 72 73 critical to its impacts on the North American low-level jets, which influence U.S. tornado 74 activity by controlling low-level vertical wind shear and moisture availability [Muñoz and 75 Enfield, 2011; Weaver et al., 2012]. 76 These recent findings have identified ENSO as a potential source of seasonal predictability 77 for U.S. tornado activity. However, it should be noted that ENSO usually decays rapidly in 78 spring, which is the most active tornado season, shortly after reaching its peak in winter. During 79 this time, the SSTAs in the tropical Pacific are typically much weaker in amplitude and their spatial structure becomes much less coherent [e.g., Lee et al. 2014a]. Additionally, every ENSO 80 81 event is somewhat different from others, which is especially true during the springtime ENSO 82 phase evolution [Trenberth and Stepaniack, 2001; Chiang and Vimont, 2004; Yu and Kim, 2010; 83 Lee et al., 2014a; Yeh et al., 2014; Capotondi et al., 2015]. For example, an ENSO event, while 84 weakening during or after spring, may subsequently evolve into the onset of another ENSO event 85 with either the same or opposite sign in the subsequent months (e.g., 1986-1987 El Niño, 1987-86 1988 El Niño and 1988-1989 La Niña). Hence, it is unlikely that the complexity of springtime 87 ENSO phase evolution can be characterized by using a single ENSO index such as the Nino3.4 88 index or Trans-Niño index. Given the previous finding that ENSO may provide the seasonal predictability of U.S. 89 90 tornado outbreaks in spring, there is a clear need to better characterize the springtime ENSO 91 phase evolution and its link to U.S. tornado activity. On this issue, a new method was recently 92 presented to objectively characterize and explore the differences in the space-time evolution of equatorial Pacific SSTAs observed during El Niño events [Lee et al., 2014b]. An application of this method to the 21 El Niño events during 1949-2013 captured two leading orthogonal modes, which explain more than 60% of the inter-event variance. The first mode distinguishes a strong and persistent El Niño from a weak and early-terminating El Niño (Figures 1a and 1b). A similar analysis applied to the 22 La Niña events during 1949-2013 also revealed two leading orthogonal modes, with its first mode distinguishing a resurgent La Niña and a transitioning La Niña (Figures 1c and 1d).

The main objective of this study is to further advance our understanding of the relationship between the springtime ENSO phase evolution and regional U.S. tornado outbreaks. To achieve this and to move forward with the goal of developing a seasonal outlook for U.S. tornado outbreaks, we first present a new metric to measure the probability of tornado outbreaks within an area centered at a given geographic location (section 2). Then, we use that metric to explore the probability of tornado outbreaks in various regions of the U.S. under the four dominant phases of springtime ENSO evolution identified in *Lee et al.* [2014b] (section 3) and to explain the associated atmospheric processes (section 4). We also report a potential link between the North Atlantic SST tripole and U.S. regional tornado outbreaks in early spring (section 5).

2. Statistical Methods and Data Used

To develop a seasonal outlook for U.S. tornado outbreaks, it is important to understand exactly what a seasonal outlook can and cannot predict. First of all, tornadogenesis is a mesoscale problem that requires overlap of very specific and highly localized atmospheric conditions [e.g., Doswell and Bosart, 2001]. Therefore, it cannot be adequately captured by

large-scale and long-term averaged atmospheric fields. In other words, a seasonal outlook cannot pinpoint exactly when, where and how many tornadoes may strike. Instead, a seasonal outlook may predict in terms of probability which regions are more vulnerable to, or more likely to experience, a widespread outbreak of tornadoes.

To move forward with the goal of developing a seasonal outlook, we propose a tornado outbreak index, which can be used to measure the probability that a tornado outbreak may occur in a predefined region. The following steps describe a method to compute the proposed tornado outbreak index for 1950 - 2014 using the Severe Weather Database (SWD) from NOAA. The Fujita scale-0 (F0) tornadoes are excluded in our analysis to avoid a spurious long-term trend in

the SWD [e.g., Verbout et al., 2006; Lee et al., 2013]. Additionally, the number of F1-F5

tornadoes is weighted in such a way that one Fn tornado is treated as n number of F1 tornadoes.

- 1) The first step is to count the weighted number of F1-F5 tornadoes within a circle of 200 km radius from the center of each 1°×1° grid point for 5 consecutive days. This is referred to as a 5-day overlapping tornado density one value for each day and grid point.
- 130 2) The second step is to define the outbreak threshold as the 99th percentile of the tornado 131 density values for each calendar month and grid point. Note that different threshold values 132 are used for each grid point and calendar month to account for the regionally and seasonally 133 inhomogeneous distribution of tornado statistics.
- The final step is to identify months with one or more *outbreak days* to construct the monthly
 U.S. regional tornado outbreak index for each grid point.

For a subset of data, the numbers of outbreak and non-outbreak years can be counted to compute the probability of U.S. regional tornado outbreaks. Figure S1 shows the outbreak threshold values of the tornado density, and the 90th and 95th percentile probabilities of U.S.

regional tornado outbreaks for March, April and May, obtained by using the steps described above.

In the following sections, we explore the probability of tornado outbreaks in various regions of the U.S. under the four dominant phases of springtime ENSO evolution (Figure 1). We use the Extended Reconstructed Sea Surface Temperature version 3b (ERSST3), an in situ analysis of global monthly SST on a 2°×2° grid [Smith et al., 2008], to compute the leading modes of ENSO variability for the period of 1949-2013 as discussed in Lee et al. [2014b]. The Twentieth Century Reanalysis (20CR) [Compo et al., 2011] and the National Centers for Environmental Prediction-National Center for Atmospheric Research (NCEP-NCAR) reanalysis [Kalnay et al., 1996] are also used to derive atmospheric anomalies associated with the four dominant phases of springtime ENSO evolution.

3. Springtime ENSO Phases and Their Links to U.S. Regional Tornado Outbreaks

Figure 1 shows the time-longitude plots of the tropical Pacific SSTAs, averaged between 5°S and 5°N, for the four leading cases of ENSO variability reproduced from *Lee et al.* [2014b]. The first case exhibits strong and positive SSTAs in the EP during the peak season persisting throughout the boreal spring (+1), and thus is referred to as a strong and persistent El Niño case (e.g., 1982-1983 El Niño); hereafter, any month/season in an ENSO onset year is identified by the suffix (0), whereas any month/season in an ENSO decay year by the suffix (+1). The second case is characterized by relatively weak positive SSTAs in the CP during the peak season and a rapid development of cold SSTAs in the EP shortly after the peak season, and thus is referred to as a weak and early-terminating El Niño case (e.g., 1963-1964 El Niño). The third case describes a La Niña persisting into the spring (+1) and evolving to another La Niña, and thus is referred to

162 as a resurgent La Niña case (e.g., 1998-1999 La Niña). This case is also frequently referred to as 163 a two-year La Niña in the literature [e.g., DiNezio and Deser, 2014]. Finally, the fourth case describes a two-year La Niña transitioning to an El Niño, and thus is referred to as a transitioning 164 165 La Niña case (e.g., 1971-1972 La Niña). Note that these four leading cases of ENSO variability 166 mainly describe ENSO phase evolution in the spring (+1) following the peak of ENSO in boreal 167 winter. For more details on the atmosphere-ocean dynamics linked to the leading modes of 168 ENSO variability, the reader is referred to *Lee et al.* [2014b]. 169 Figure 2 shows the composite SSTAs for the four dominant phases of springtime ENSO 170 evolution in MAM (+1) and the corresponding probability of regional tornado outbreaks in April 171 (+1). See Figures S3-S6 in the supporting information for the composite SSTAs and the 172 probability of outbreaks in March (+1) and May (+1). The gray dots indicate that the SSTAs are statistically significant at 90% based on a Student's t-test. Similarly, the black dots mean that the 173 174 probability of tornado outbreaks is statistically significant at 90% based on the exact binomial test of the null hypothesis, i.e., the springtime ENSO phases have no effect on the probability of 175 176 tornado outbreaks. For each case, we used eight to eleven actual ENSO events for the composite 177 analysis based on the sign and amplitude of the leading principal components of El Niño and La 178 Niña variability (see Table S2 for the list of ENSO events used). 179 As shown in Figures 2a and 2e, the probability of U.S. regional tornado outbreaks in April 180 (+1) is low overall (5% or less) when a strong El Niño persists into the spring (+1) after its peak (i.e., persistent El Niño). The probability of outbreaks may reach as high as 25% along the 181 Mississippi river and over New York and central Florida. However, they are statistically 182 183 insignificant at the 90% confidence level. The overall low probability of outbreaks shown in 184 Figure 2e implies that the outbreak frequency is largely suppressed by a strong El Niño that

persists throughout the boreal spring (+1), as suggested by Allen et al. [2015]. However, the 185 186 statistical significance of the reduction cannot be established since the frequency distribution of 187 the outbreak chance is highly skewed to the right. 188 Similar to the persistent El Niño case, when a weak El Niño terminates early and cold SSTAs develop over the EP in the spring (+1) (i.e., early-terminating El Niño), the probability of 189 190 outbreaks is also low overall (5% or less) although it increases somewhat in limited regions over 191 Missouri and northern Texas in April (+1) (Figures 2b and 2f), and over Iowa, Wisconsin and 192 Illinois in May (+1) (Figure S4f). 193 When a La Niña persists into the spring (+1) and evolves to another La Niña (i.e., resurgent La Niña), the probability of outbreaks increases significantly up to 27% in April (+1) in wide 194 195 regions in the Ohio Valley, Southeast, Upper Midwest and Great Plains (see Figure S2 for the 196 U.S. climate regions as defined by the National Climate Data Center), particularly Nebraska, 197 Illinois, Indiana, Kentucky, Tennessee, Mississippi, Alabama, Georgia, Virginia, North Carolina 198 and Wisconsin (Figures 2c and 2g). It is interesting to note that the record-breaking 2011 tornado 199 outbreaks occurred during a resurgent La Niña. Similarly, the super tornado outbreaks in 1974 occurred during a resurgent La Niña. 200 201 As shown in Figure 2d, when a two-year La Niña transitions to an El Niño (i.e., transitioning 202 La Niña), the cold SSTAs in the CP are nearly dissipated away while the warm SSTAs in EP 203 become strong and statistically significant in MAM (+1). In this case, the probability of tornado outbreaks increases strongly and significantly up to 38% in the South U.S., particularly over 204 205 Kansas, eastern Colorado and Texas in April (+1) (Figure 2h), and over southern Texas in March (+1) (Figure S6f). 206

Two important questions arise among others as to why the probability of U.S. regional tornado outbreaks increases in the spring following the peak of La Niña, and why the regions affected during the resurgent La Niña case are quite different from the regions affected during the transitioning La Niña case. We attempt to address these questions in the next section.

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4. Springtime Atmospheric Variability over the U.S. Linked to ENSO Phases

It is well known that La Niña causes the winter atmospheric jet stream to take an unusually wavy southeastward path into the U.S. from southwestern Canada, thus bringing colder and drier upper level air to the U.S. Hence, the winter storm activity over the U.S. tends to increase, causing more frequent wet conditions particularly over the Ohio Valley [e.g., Ropelewski and Halpert, 1986; Eichler and Higgins, 2006; Mo, 2010]. As shown in Figure 3a, in the spring (+1) of the resurgent La Niña case, an anomalous cyclone develops over the U.S. bringing colder and drier upper-level air to the U.S., and thus the extratropical storm activity increases, suggesting that the mechanism through which La Niña affects U.S. weather in winter months still prevails in the spring (+1) of the resurgent La Niña case [Lee et al., 2014a]. The anomalous cyclone and the associated increase in equivalent barotropic winds in turn enhance the low-level vertical wind shear (850 - 1000 hPa) east of the Rockies, and shift the moisture stream originating from the Gulf of Mexico more toward the east (Figure 3b), thus producing a set of favorable atmospheric environments for tornado outbreaks in the Southeast and Ohio Valley, consistent with Figure 2g. In the spring (+1) of the transitioning La Niña case, on the other hand, the extratropical storm activity decreases slightly over the South and Southeast. Consistent with this feature, an anomalous anticyclone forms east of the Rockies inducing anomalous southerly winds over the South (Figure 3c). Therefore, the low-level vertical wind shear increases over the South and the

230	moisture stream from the Gulf of Mexico is enhanced more toward the South (Figure 3d). These
231	changes in the atmospheric environments are largely consistent with the increased probability of
232	tornado outbreaks in the South (Figure 2h).
233	The large-scale atmospheric patterns in the spring (+1) following the peak of El Niño are
234	largely opposite to those following the peak of La Niña [e.g., Lee et al., 2014a]. Thus, as shown
235	in Figure S7, the atmospheric environments over the U.S. linked to the persistent and early-
236	terminating El Niño cases are largely unfavorable for tornado outbreaks, in line with Figures 2e
237	and 2f.
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239	5. North Atlantic SST Tripole and U.S. Regional Tornado Outbreaks
240	As shown in Figure S8, there is a coherent and statistically significant pattern of springtime
241	SSTAs in the North Atlantic linked to the extreme U.S. tornado outbreaks (see Tables S3 for the
242	list of extreme U.S. tornado outbreak years used). This pattern is very similar to the North
243	Atlantic SST tripole, which is the dominant mode of interannual SST variability in the Atlantic
244	in boreal winter and spring and is known to be linked to multiple forcing mechanisms including
245	the North Atlantic Oscillation, tropical Atlantic variability and extratropical teleconnections from
246	the tropical Pacific [e.g., Xie and Tanimoto, 1998; Okumura et al., 2001; Peng et al., 2002; Wu et
247	al., 2007; Schneider and Fan, 2012]. Hence, we further explore the potential link between the
248	North Atlantic SST tripole and the probability of U.S. regional tornado outbreaks. First, we
249	performed an empirical orthogonal function (EOF) analysis of the North Atlantic SSTAs in
250	MAM to sort the past 66 years (i.e., 1949-2014) based on the amplitude of the North Atlantic
251	SST tripole mode (i.e., the leading EOF mode). Then, we selected the top (i.e., positive) 10 cases

and the bottom (i.e., negative) 10 cases from the sorted years to perform composite analysis (see Table S4 for the list of the positive and negative North Atlantic SST tripole years used).

As summarized in Figures 4 and S9, the North Atlantic SST tripole is indeed linked to the probability of U.S. regional tornado outbreaks in March and April. During its negative phase (i.e., cold in the tropical North Atlantic, warm in the subtropical North Atlantic and cold in the subpolar North Atlantic), a robust anomalous anticyclone straddles the subtropical North Atlantic extending westward over the U.S. In March, the increase in the equivalent barotropic winds along the poleward half of the anomalous anticyclone enhances the low-level vertical wind shear over the Upper Midwest and Ohio Valley (Figure 4b). It appears that the moisture transport increases somewhat along the gulf coast mainly toward the Southeast, which is likely due to the increased SSTAs in the Gulf of Mexico (Figure 4c). These changes in the low-level vertical wind shear and moisture transport are fairly consistent with the significantly increased probability of tornado outbreaks in the Southeast and Upper Midwest (Figure 4a). In April, the anomalous anticyclone over the U.S. retreats to the east enhancing the moisture supply and the low-level vertical wind shear along the South and Great Plains (Figures S9a-S9c). These relationships are almost exactly the opposite during the positive phase of the North Atlantic SST tripole (Figures 4d-4f and S9d-S9f).

Although the results summarized in Figures 4 and S9 are promising, it must be noted that a negative phase of the North Atlantic SST tripole forms more frequently in the early spring following the peak of La Niña (Table S3). Therefore, it is unclear whether the North Atlantic SST tripole adds much to the predictability of U.S. regional tornado outbreaks. Further studies using model experiments and advanced statistical methods are needed to better understand the impact of North Atlantic SST variability on the probability of U.S. regional tornado outbreaks.

6. Discussion

This study illustrates the links between the leading modes of springtime ENSO variability 10 and the probability of U.S. regional tornado outbreaks. However, the leading modes of El Niño and La Niña variability represent only about 35% of the total variance [Lee et al., 2014b]. Thus, it is important to take into account the second leading modes, which represent an additional approximately 28% of the total variance [Lee et al., 2014b]. The two leading modes of ENSO variability and the North Atlantic SST tripole mode could be used as predictors of the monthly U.S. regional tornado outbreak index, using a logistic regression analysis [Cox, 1958]. The regression coefficients obtained from the logistic regression analysis could be used to estimate the probability of U.S. regional tornado outbreaks given the amplitudes of the three predictors. Combining this statistical tool with a dynamic seasonal forecast model, which could be used to obtain the three predictors with 1-3 months lead time, it might be possible to build a seasonal outlook for U.S. regional tornado outbreaks. To this end, it is quite promising that high-resolution climate models are now beginning to demonstrate skill in simulating and predicting seasonal variations in some of the elements critical to U.S. tornado outbreak risk [e.g., Krishmamurthy et al., 2015; Yang et al., 2015; Jia et al., 2015].

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387 Figure captions: 388 Figure 1. Time-longitude plots of the leading orthogonal modes of the tropical Pacific SSTAs 389 averaged between 5°S and 5°N for 21 El Niño and 22 La Niña events during 1949-2013, 390 reproduced from Lee et al. [2014b], namely (a) the persistent El Niño, (b) early-terminating El 391 Niño, (c) resurgent La Niña and (d) transitioning La Niña. The unit is in °C. 392 393 Figure 2. Composite (a-d) SSTAs for the four dominant phases of springtime ENSO evolution in 394 MAM (+1) and (e-h) the corresponding probability of U.S. regional tornado outbreaks in April (+1). The gray dots in panels a-d indicate that the SSTAs are statistically significant at 90% 395 396 based on a student-t test. The black dots in panels e-h indicate that the probability of tornado 397 outbreaks is statistically significant at 90% based on a binomial test. The unit is in °C for the 398 SSTAs and in % for the probability of tornado outbreaks. 399 400 Figure 3. (upper row) Anomalous geopotential height at 500 hPa (color shades) and variance of 401 5day high-pass filtered meridional winds at 300 hPa (contours), and (lower row) anomalous 402 moisture transport (vectors) and low-level vertical wind shear (850 - 1000 hPa; color shades) in 403 MAM (+1) for (a,b) the resurgent La Niña and (c,d) transitioning La Niña cases. The units are in gpm for geopotential height, in m² s⁻² for variance of meridional winds, in kg m⁻¹ s⁻¹ for moisture 404 transport, and in m s⁻¹ for vertical wind shear. 405 406 407 Figure 4. (top row) Probability of U.S. regional tornado, (middle row) composite SSTAs (color 408 shades) and geopotential height anomalies at 500 hPa (contours), and (bottom row) low-level 409 vertical wind shear anomalies (color shades) and moisture transport anomalies (vectors) in

410	March for (a-c) the negative and (d-f) positive North Atlantic SST tripole. The unit is in % for				
411	the probability of tornado outbreaks, in °C for the SSTAs, in gpm for geopotential height, in kg				
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