

1 **Secular and Multidecadal Warming of the Atlantic Ocean since the mid-**  
2 **20th Century**

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7 Sang-Ki Lee<sup>1,2</sup>, David B. Enfield<sup>1,2</sup>, Wonsun Park<sup>3</sup>, Erik van Sebille<sup>5</sup>, Chunzai Wang<sup>2</sup>, Steven  
8 Yeager<sup>4</sup>, Ben Kirtman<sup>5</sup>, and Molly Baringer<sup>2</sup>

9 <sup>1</sup>Cooperative Institute for Marine and Atmospheric Studies, University of Miami, Miami,  
10 Florida, USA

11 <sup>2</sup>Atlantic Oceanographic and Meteorological Laboratory, NOAA, Miami Florida, USA

12 <sup>3</sup>Leibniz Institute of Marine Sciences (IFM-GEOMAR), Kiel, Germany

13 <sup>4</sup>National Center for Atmospheric Research, Boulder, Colorado, USA

14 <sup>5</sup>Rosenstiel School for Marine and Atmospheric Science, University of Miami, Miami, Florida,  
15 USA

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22 Corresponding author address: Dr. Sang-Ki Lee, NOAA/AOML, 4301 Rickenbacker Causeway,  
23 Miami, FL 33149, USA. E-mail: [Sang-Ki.Lee@noaa.gov](mailto:Sang-Ki.Lee@noaa.gov).

1 As seen in the extended reconstructed sea surface temperature data (ERSST3; Smith et al.  
2 2008), the North Atlantic sea surface temperature (SST) during the instrumental period (Figure  
3 1a) contains both a secular increase at a rate of  $\sim 0.4^{\circ}\text{C}$  per 100yrs during 1901-2010 and a  
4 robust multidecadal signal of similar amplitude, known as the Atlantic multidecadal oscillation  
5 (AMO; Kerr, 2000). Coupled general circulation models forced by external factors such as  
6 greenhouse gases and solar variations do a poor job of reproducing the AMO in the 20th century,  
7 suggesting that it must arise instead from internal interactions of the climate system (e.g., Knight,  
8 2009). It has been shown that in an unforced coupled climate model simulation the Atlantic  
9 meridional overturning circulation (AMOC) exhibits decadal to multidecadal (15 - 70 yrs)  
10 variations and is closely linked both dynamically and statistically to the AMO (e.g., Delworth et  
11 al., 1993). However, neither the long-term variations of the AMOC nor the link between the  
12 AMOC and AMO has been demonstrated with observations in the 20th century because  
13 historical observations of the AMOC are not up to the task.

14 Contrary to the North Atlantic SST, the South Atlantic SST during the instrumental period  
15 (Figure 1b) has increased almost linearly at a rate of  $\sim 0.9^{\circ}\text{C}$  per 100yrs during 1901-2010,  
16 clearly surpassing the global average of  $0.6 \sim 0.7^{\circ}\text{C}$  per 100yrs. Consistent with this surface  
17 trend, recently updated and bias-corrected instrumental records indicate that the heat content of  
18 the Atlantic Ocean in the upper 700m has substantially increased during the 1970s – 2000s at a  
19 rate ( $\sim 8 \times 10^{22}$  J per 40yrs) almost matching that of the Pacific Ocean ( $\sim 6 \times 10^{22}$  J per 40yrs) and  
20 Indian Ocean ( $\sim 2 \times 10^{22}$  J per 40yrs) combined (Levitus et al., 2009), even though the Atlantic  
21 Ocean covers less than 20% of the global ocean in surface area. As a candidate mechanism for  
22 the differential inter-ocean warming, Lee et al. (2011) pointed out the potential role played by  
23 the global overturning circulation. They argued that as the upper ocean warms globally during

1 the 20th century, the inter-ocean heat transport associated with the global overturning circulation  
2 should increase until the deep ocean fully adjusts to the surface warming or the global  
3 overturning circulation slows down. Since the Atlantic Ocean is characterized with advective  
4 heat convergence, they argued that the Atlantic Ocean should therefore gain extra heat from  
5 other oceans. They used a surface-forced global ocean-ice coupled model to test this hypothesis  
6 and to further show that the increased AMOC at 30°S, forced by the increased wind stress curl  
7 over the region from 50° to 30°S, has contributed greatly to the increased inter-ocean heat  
8 transport from the Indian Ocean during the latter half of the 20th century.

9 To sum up, it appears that the AMOC and the associated meridional ocean heat transport  
10 hold the key to our understanding of the observed secular and multidecadal warming of the  
11 North and South Atlantic Oceans since the mid-20th century. Therefore, it is vital to reconstruct  
12 the history of AMOC and associated meridional ocean heat transport in the 20th century,  
13 covering at least one full cycle of the AMO to be able to cleanly distinguish the secular trend  
14 from the multidecadal variations. Here, we attempt an ocean model-based reconstruction, which  
15 is likely to be our best chance for assessing the history of AMOC in the 20th century because the  
16 observed surface flux fields, which constrain ocean-only (or ocean-ice coupled) models, are  
17 available on relatively long time scales.

18

### 19 **20th Century Reanalysis (20CR)**

20 None of the surface-forced ocean model studies so far has been simulated with the surface  
21 forcing prior to the mid-20th century because the surface forcing data, which are typically  
22 derived from atmospheric reanalysis products such as NCEP-NCAR reanalysis, are limited to the  
23 last 50 - 60 years. Recently, the newly developed NOAA-CIRES 20th Century Reanalysis

1 (20CR) has been completed (Compo et al., 2010). The 20CR provides the first estimate of global  
2 surface momentum, heat and freshwater fluxes spanning the late 19th century and the entire 20th  
3 century (1871-2008) at daily temporal and 2° spatial resolutions.

4

## 5 **Model Experiments**

6 The global ocean-ice coupled model of the NCAR Community Climate System Model  
7 version 3 (CCSM3) forced with the 20CR is used as the primary tool in this study. The ocean  
8 model is a level-coordinate model based on the Parallel Ocean Program (POP), divided into 40  
9 vertical levels. The ice model, the NCAR Community Sea Ice Model version 5, is a dynamic-  
10 thermodynamic ice model. Both the ocean and ice models have 320 longitudes and 384 latitudes  
11 on a displaced pole grid with a longitudinal resolution of about 1.0 degrees and a variable  
12 latitudinal resolution of approximately 0.3 degrees near the equator. See Doney et al. (2007) for  
13 more detailed descriptions about the CCSM3 ocean-ice model (CCSM3\_POP hereafter).

14 After the total of 900 years of spin up runs (see Lee et al. (2011) for a detailed description of  
15 the spin up runs), three model experiments are performed. In the control experiment  
16 (EXP\_CTR), the CCSM3\_POP is integrated for 1871-2008 using the real-time daily 20CR  
17 surface flux fields. The next two experiments are idealized experiments designed to understand  
18 the Atlantic Ocean heat content change with and without the influence of the northward heat  
19 transport change at 30°S. The remote ocean warming experiment (EXP\_REM) is identical to  
20 EXP\_CTR except that the surface forcing fields north of 30°S are from the daily 20CR surface  
21 flux fields for the period of 1871-1900 exactly like the spin-up experiment, whereas those south  
22 of 30°S are real time as in EXP\_CTR. The Atlantic Ocean warming experiment (EXP\_ATL) is  
23 also identical to EXP\_CTR except that the surface forcing fields south of 30°S are from the daily

1 20CR surface flux fields for the period of 1871-1900 as in the spin-up experiment, whereas those  
2 north of 30°S are real time as in EXP\_CTRL. Note that the Atlantic Ocean warms only through  
3 anomalous surface warming (i.e., local processes) in EXP\_ATL, and only through anomalous  
4 northward ocean heat transport at 30°S (i.e., remote processes) in EXP\_REM, respectively.

5

## 6 **Model Results**

7 Figure 2a shows the simulated North Atlantic Ocean heat content change in the upper 700m  
8 in reference to the 1871 - 1900 period obtained from the three model experiments, along with the  
9 observed trend of the North Atlantic Ocean heat content during 1969 – 2008. The observed trend  
10 is referenced at 1969 in Figure 2 for a better visual comparison with the simulated trends. The  
11 simulated North Atlantic Ocean heat content in EXP\_CTRL increases moderately during the  
12 1930s - 1940s, and then decreases during the 1940s - 1970s, after which it increases substantially  
13 much like the AMO from observations (Figure 1a). Between the 1970s and 2000s, it increases by  
14  $3 \sim 4 \times 10^{22}$  J. This large increase is reasonably close to the observed North Atlantic Ocean heat  
15 content increase of  $\sim 5.5 \times 10^{22}$  J during the same period (Levitus et al., 2009), suggesting that the  
16 model experiment (EXP\_CTRL) reproduces reasonably well the heat budget trend of the North  
17 Atlantic Ocean after the 1960s.

18 If the northward heat transport in the South Atlantic at 30°S is fixed at its 1871-1900 level by  
19 considering the fully transient surface fluxes only north of 30°S (EXP\_ATL), the simulated  
20 North Atlantic Ocean heat content increases during the 1930s - 1940s, and then decreases during  
21 the 1940s - 1970s, almost perfectly reproducing that of EXP\_CTRL prior to the 1970s. However,  
22 the North Atlantic Ocean heat content, in this case, increases by only  $\sim 1 \times 10^{22}$  J during the 1970s

1 and 2000s, and thus cannot explain the observed North Atlantic Ocean heat content increase  
2 during the same period.

3 On the other hand, if the northward heat transport in the South Atlantic at 30°S is allowed to  
4 vary in real time by considering the fully transient surface fluxes south of 30°S while keeping the  
5 surface fluxes over the Atlantic Ocean at their 1871-1900 levels (EXP\_REM), the North Atlantic  
6 Ocean heat content increases by  $\sim 2 \times 10^{22}$  J during the 1970s – 2000s explaining a moderate  
7 portion of the observed trend. In this case, however, the multidecadal signal during the 1930s –  
8 1970s, which is clearly simulated in both EXP\_CTRL and EXP\_ATL is completely missing. The  
9 absence of this multidecadal signal in EXP\_REM and its presence in EXP\_ATL clearly suggest  
10 that the multidecadal swing in EXP\_CTRL prior to the 1970s is caused by processes internal to the  
11 Atlantic Ocean. During the 1970s – 2000s, on the other hand, remote processes (i.e., increased  
12 inter-ocean heat transport from the Indian Ocean; see Lee et al. (2011)) have contributed more to  
13 the large increase in the North Atlantic heat content, although internal processes have also  
14 contributed.

15 Figure 2b is the same as Figure 2a, except for the South Atlantic Ocean heat content change.  
16 The simulated South Atlantic Ocean heat content in EXP\_CTRL remains unchanged until the  
17 1960s, after which it increases monotonically, consistent with the similar increase in the South  
18 Atlantic SST during the same period from observations (Figure 1b). Between the 1970s and  
19 2000s, it increases by  $\sim 2 \times 10^{22}$  J, slightly less than the observed South Atlantic Ocean heat  
20 content increase of  $\sim 2.5 \times 10^{22}$  J during the same period (Levitus et al., 2009). As in EXP\_CTRL,  
21 the South Atlantic Ocean heat content in EXP\_REM is also characterized with a monotonic  
22 increase after the 1960s, but with a smaller amplitude of  $\sim 1 \times 10^{22}$  J during the 1970s – 2000s.  
23 In the case of EXP\_ATL, however, there is no apparent change in the South Atlantic heat content

1 throughout the 20th century. These results derived from Figure 2b lead to a conclusion that  
2 remote processes mainly drive the South Atlantic Ocean heat content increase during the 1970s -  
3 2000s in EXP\_CTR.

4 Figure 3a shows the time-averaged AMOC during 1979-2008 obtained from EXP\_CTR. The  
5 simulated maximum strength of the AMOC at 35°N is only 11 Sv ( $1\text{Sv} = 10^6 \text{ m}^3\text{s}^{-1}$ ), which is  
6 smaller than the observed range of 14 ~ 20Sv. Despite the smaller maximum strength, the overall  
7 spatial structure of the simulated AMOC is quite close to that derived from observations (e.g.,  
8 Lumpkin and Spear, 2007). Figure 3b shows the time series of the simulated AMOC index  
9 (maximum overturning stream function) at three different latitudes, namely 30°S, the equator,  
10 and 60°N, in reference to the 1871 – 1900 period. It is clear that the AMOC at all three latitudes  
11 increase after the 1950s, with the largest amplitude at 30°S, a lesser amplitude at the equator, and  
12 the smallest amplitude at 60°N. This suggests that both the North (0° – 60°N) and South (30°S –  
13 0°) Atlantic Oceans are subject to advective heat convergence during the 1960s – 2000s,  
14 consistent with the heat content increase in the North and South Atlantic Oceans during the  
15 1970s – 2000s (Figure 2). It appears that more works are needed to understand if and how the  
16 AMOC is linked with the multidecadal variations of the North Atlantic Ocean heat content  
17 especially prior to the 1960s.

18

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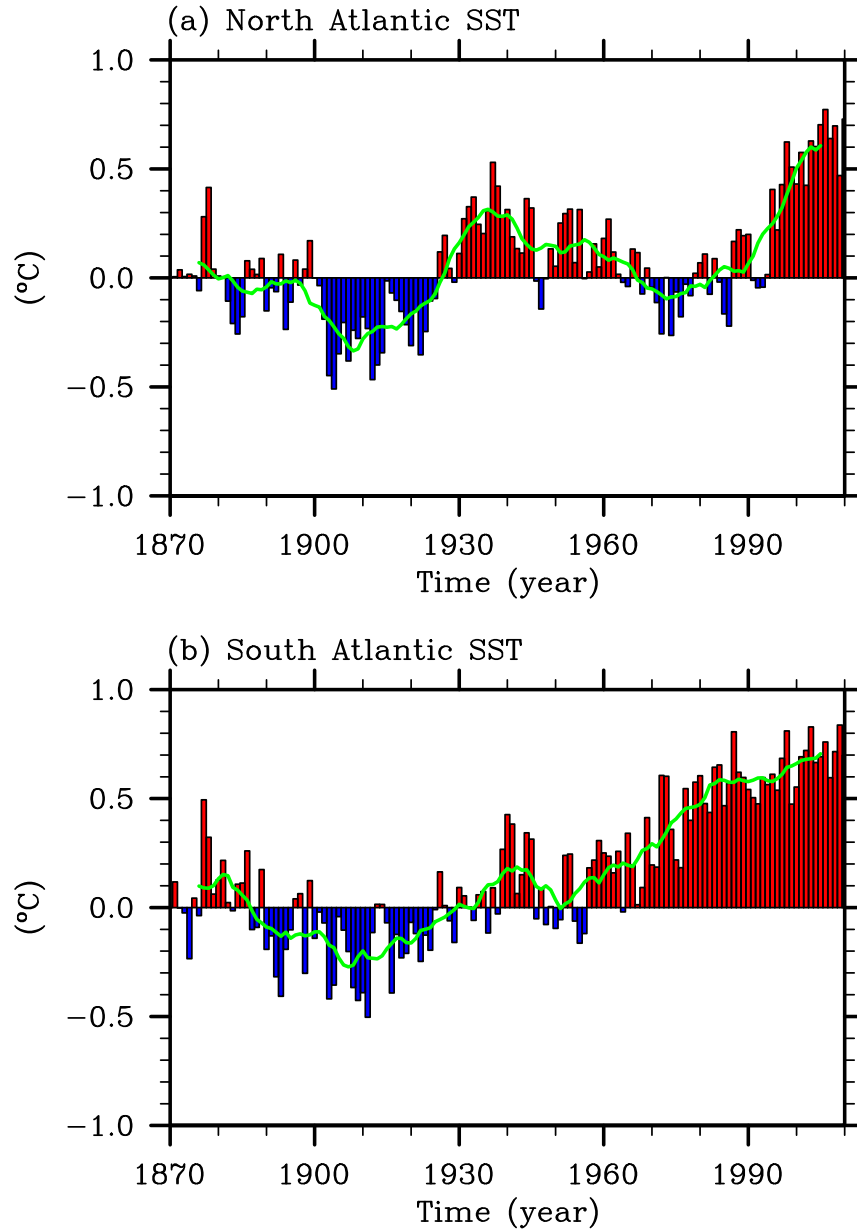
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1 **Figure 1.** Observed (a) North ( $0^{\circ} - 60^{\circ}\text{N}$ ) and (b) South Atlantic ( $30^{\circ}\text{S} - 0^{\circ}$ ) SSTs during the  
2 instrumental period, obtained from ERSST3. The green line in each plot is obtained by  
3 performing a 11-year running average.

4  
5 **Figure 2.** Simulated (a) North and (b) South Atlantic Ocean heat content changes in the upper  
6 700m in reference to 1871-1900 obtained from the three model experiments. The thick black  
7 lines in (a) and (b) are the observed trends of the North and South Atlantic Ocean heat content,  
8 respectively, reproduced from Levitus et al. (2009).

9  
10 **Figure 3.** (a) Time-averaged AMOC during 1979-2008 and (b) 11-year running averaged time  
11 series of the simulated AMOC index (maximum overturning stream function) at  $30^{\circ}\text{S}$ , the  
12 equator, and  $60^{\circ}\text{N}$  in reference to 1871-1900 obtained from EXP\_CTR.

ERSST3: Atlantic SST Anomaly



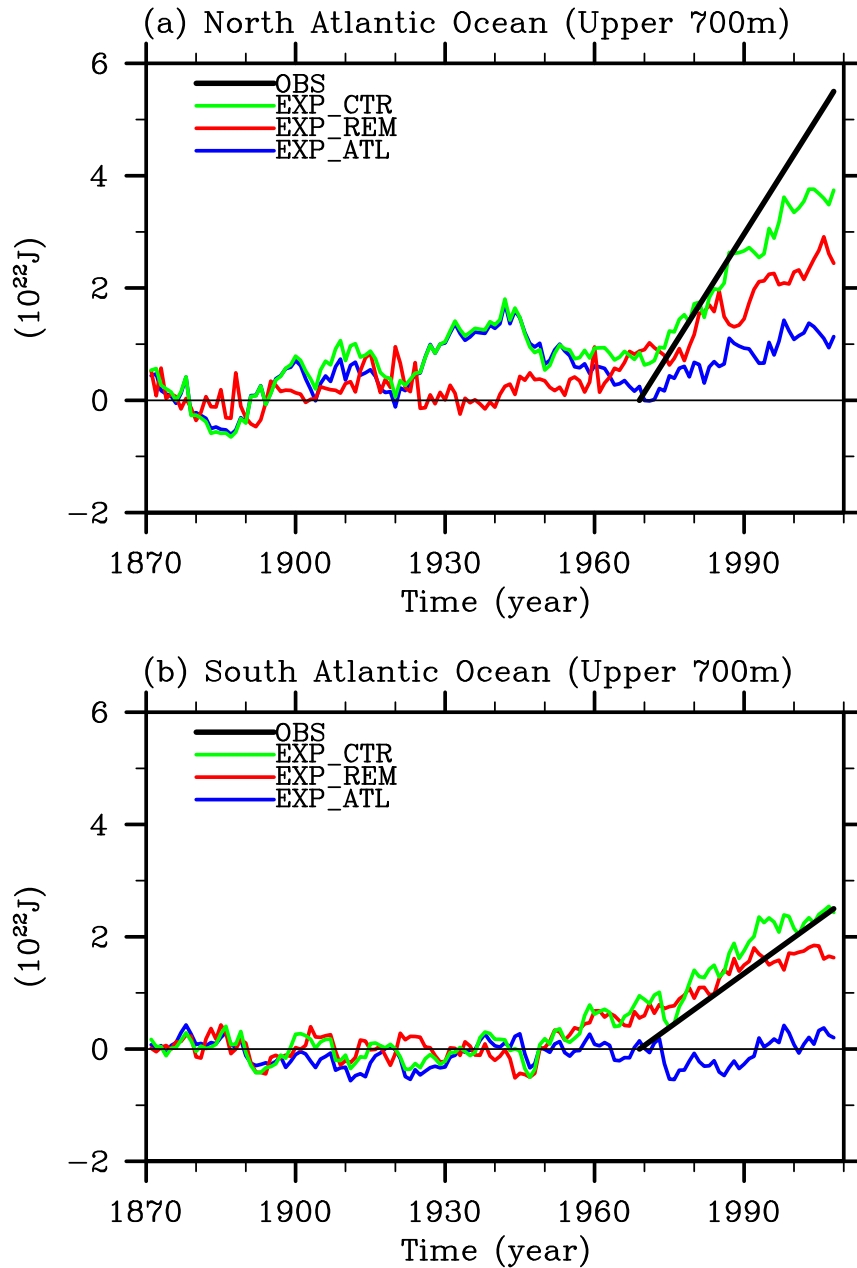
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2 **Figure 1.** Observed (a) North ( $0^{\circ} - 60^{\circ}\text{N}$ ) and (b) South Atlantic ( $30^{\circ}\text{S} - 0^{\circ}$ ) SSTs during the  
3 instrumental period, obtained from ERSST3. The green line in each plot is obtained by  
4 performing a 11-year running average.

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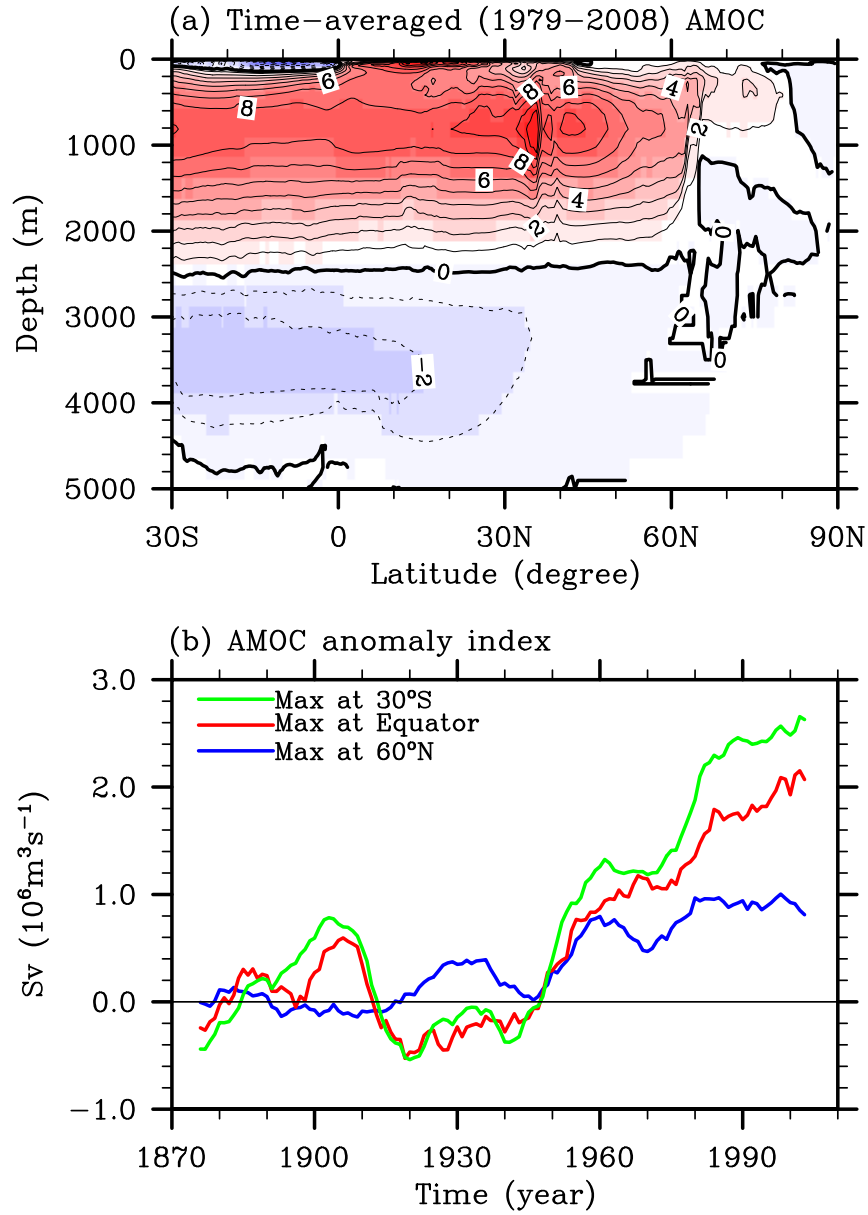
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CCSM3\_POP: Atlantic Ocean Heat Content



1  
2 **Figure 2.** Simulated (a) North and (b) South Atlantic Ocean heat content changes in the upper  
3 700m in reference to 1871-1900 obtained from the three model experiments. The thick black  
4 lines in (a) and (b) are the observed trends of the North and South Atlantic Ocean heat content,  
5 respectively, reproduced from Levitus et al. (2009).

CCSM3\_POP (EXP\_CTR): AMOC



1  
2 **Figure 3.** (a) Time-averaged AMOC during 1979-2008 and (b) 11-year running averaged time  
3 series of the simulated AMOC index (maximum overturning stream function) at 30°S, the  
4 equator, and 60°N in reference to 1871-1900 obtained from EXP\_CTR.