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14        **Abstract**

15        Since 1994 the US Global Drifter Program (GDP) and its international partners cooperating within the  
16        Data Buoy Cooperation Panel (DBCP) of WMO-UNESCO have been deploying drifters equipped with  
17        barometers primarily in the extra-tropical regions of the world's oceans in support of operational weather  
18        forecasting. To date, the impact of the drifter data isolated from other sources has never been studied.  
19        This essay quantifies and discusses the effect and the impact of in situ sea-level atmospheric pressure  
20        (SLP) data from the global drifter array on numerical weather prediction using observing system  
21        experiments and forecast sensitivity observation impact studies. The in situ drifter SLP observations are  
22        extremely valuable to anchor the global surface pressure field and significantly contribute to accurate  
23        marine weather forecasts, especially in regions where no other in situ observations are available, like, for  
24        example, in the Southern Ocean. Furthermore, the forecast sensitivity observation impact analysis  
25        indicates that The SLP drifter data is the most valuable per-observation contributor from the Global  
26        Observing System (GOS). All these results give evidence that surface pressure observations of drifting  
27        buoys are essential ingredients of the GOS and their quantity, quality and distribution should be preserved  
28        as much as possible in order to avoid any analysis and forecast degradations. The barometer upgrade  
29        program offered by the GDP, under which GDP funded drifters can be equipped with partner-funded  
30        accurate air pressure sensors, is a practical example of how the DBCP collaboration is executed. Interested  
31        parties are encouraged to contact the GDP to discuss upgrade opportunities.

32        **Capsule Summary**

33        *In-situ*, sea-level air pressure data from the global array of surface drifters significantly contribute to  
34        accurate marine weather forecasting

35 **A Global Array of Drifting Barometers**

36 Since 1994, the NOAA-funded Global Drifter Program (GDP; [Maximenko et al., 2013; Niiler, 2001]), in  
37 collaboration with the international partners of the Data Buoy Cooperation Panel (DBCP), a joint body of  
38 the World Meteorological Organization (WMO) and the Intergovernmental Oceanographic Commission  
39 (IOC) of UNESCO, has been deploying Surface Velocity Program (SVP<sup>1</sup>) Lagrangian drifters drogued at 15m  
40 depth and equipped with barometers (SVPB hereafter) in the world’s oceans with focus in the extra-  
41 tropical regions. The SVPB drifters are designed to make accurate measurements of Sea-Level  
42 Atmospheric Pressure (SLP) and to report the data in real-time through the Global Telecommunication  
43 System (GTS) of the WMO Information System (WIS) in order to contribute to the World Weather Watch  
44 (WWW).

45 The synergy between the oceanographers and the meteorologists, particularly under the WMO and  
46 IOC umbrella, has fostered a very successful and ongoing collaboration in which the oceanographers that  
47 manage the GDP offer the opportunity to upgrade standard drifters with barometers for a very low cost  
48 and the meteorologists provide hardware and valuable deployment infrastructures for the drifters.

49 In general, oceanographers are mostly concerned with studying the circulation and the dynamics of  
50 the ocean currents at global and regional scales as well as gathering accurate in-situ sea surface  
51 temperature (SST) data whilst meteorologists are mainly interested in global in-situ SLP data with  
52 particular attention in regions where observations are sparse.

53 The SLP data from drifters are generally regarded as important for operational weather forecasting  
54 and for other oceanographic and severe weather forecast applications [Healy, 2013; Maximenko et al.,

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<sup>1</sup> The name SVP (and SVPB, etc.) for designating Lagrangian drifters was kept as a legacy of the former Surface Velocity Programme (SVP) of the World Ocean Circulation Experiment (WOCE, 1990-2002)

55 2013; *WIGOS*, 2012]. However, until now, a formal assessment of the effect and impact of SLP data from  
56 SVPB drifters on numerical weather prediction (NWP) was never conducted. This essay reports on the  
57 main findings of a study sponsored by the DBCP under the framework of the pilot project called  
58 “Evaluation of the Impact of Sea Level Atmospheric Pressure Data Over the Ocean from Drifting Buoys on  
59 Numerical Weather Prediction Models” (PP-SLP) and it is meant to raise awareness among the  
60 oceanographic and atmospheric communities of the role of the global drifter array in supporting NWP and  
61 climate services. This essay is also intended to promote the drifter barometer upgrade program sponsored  
62 by the GDP and discussed every year at the DBCP plenary sessions.

### 63 **The Global Drifter Program, the SVPB Drifter Technology and the Data Denial Study**

64 The SVP drifter design emerged from “holey-sock” drogue drifters deployed as early as 1979 in the  
65 Tropical Pacific and standardized in 1987 as part of the former Tropical Ocean and Global Atmosphere  
66 (TOGA) program [*Lumpkin and Pazos*, 2007; *Niiler*, 2001]. This regional drifter program quickly grew into  
67 a global array and to date more than 22,000 drifters have been deployed to fulfill the GDP objective of  
68 maintaining a global array of 1,250 drifters. This size of the array is sufficient to keep the potential satellite  
69 SST bias error (PSBE) below 0.5°C because the number and distribution of drifters, as well as the accuracy  
70 of their SST data, that ranges between 0.05°C and 0.1°C, is directly proportional to the PSBE [*Zhang et al.*,  
71 2009]. The GDP was the first component of the Global Ocean Observing System (GOOS) to be fully  
72 implemented when the array reached 1250 drifters for the first time on September 18, 2015.

73 The SVPB drifter (Figure 1) has the same drogue (sea anchor) and surface buoy of the SVP drifter  
74 [*Niiler*, 2001]. The drogue is a cylindrical tube of Cordura® nylon connected to the surface buoy with a  
75 tether. The center of the drogue is located at a depth of 15m. The much larger drag of the drogue than  
76 that of tether and surface buoy combined ensures that the drifter behaves as a Lagrangian instrument,  
77 i.e. that it moves with the same velocity of the surrounding water at the target depth of 15m. The error

78 of the Lagrangian velocity is essentially due to the slip of the drogue through the water due to the action  
79 of wind and waves on the surface buoy, and it is less than  $1 \cdot 10^{-2} \text{ ms}^{-1}$  for winds up to  $10 \text{ ms}^{-1}$  [Niiler *et al.*,  
80 1995]. A more complete description of the SVP drifter technology can be found in Niiler [2001].

81 Since the drag of the drogue is much larger than that of the surface buoy, the latter is often pulled  
82 underwater by surface gravity waves. Therefore the tube that connects the barometer sensor with the  
83 atmosphere is protected from water intrusions by a self-draining air pressure port waterproofed by two  
84 Gore-Tex® screens. The invalid air pressure readings taken when the drifter is submerged are removed  
85 onboard before being transmitted. The SLP filtering algorithm is very robust and has been successfully  
86 tested in a variety of environments including hurricane conditions. Drifters are fitted with either a High  
87 Precision Barometer (HPB) by Honeywell, stable over the two years long nominal lifespan of the drifters  
88 and with an accuracy of  $\pm 0.4 \text{ hPa}$ , or with an Integrated Pressure Transducer, also by Honeywell that has  
89 similar specifications and accuracy.

90 The SLP is measured every hour and two satellite data telecommunication systems, Argos or Iridium,  
91 are used to telemeter the data. The data latency depends mostly on the satellite system of choice. The  
92 Argos satellite network adds an average of about one to two hours to the data latency, but the average  
93 delay drops to a few minutes if the Iridium satellite system is used. The GDP is targeting an optimum mix  
94 of data telemetry communications to minimize the data latency. Additional delays typically of less than  
95 15 minutes, and inherent to the way the data are processed, quality controlled, encoded, and distributed  
96 over the GTS, are also introduced.

97 Since 2007 about 50% of drifters in the global array are of the SVPB type and the growth of the  
98 barometer array has been primarily limited by funding. Most of the SVPB hardware is supported by the  
99 GDP but significant contributions also come from DBCP members by way of SVPB hardware purchase or  
100 upgrades of GDP's SVP drifters with barometers. The operational service for Surface Marine Observations

101 (E-SURFMAR) of the Economic Interest Group (EIG) of European National Meteorological Services  
102 (EUMETNET) also provides an important contribution to the SVPB drifter array, mainly in the North  
103 Atlantic Ocean.

104 Besides the SVPB drifters, other sources of *in situ* SLP data over the ocean exist, and these include but  
105 are not limited to Automatic Weather Stations (AWS) installed on Voluntary Observing Ship (VOS) and  
106 moored buoys. The AWS coverage is limited to standardized shipping lanes and leave substantially under-  
107 sampled areas in the southern hemisphere and in the southern ocean. The SLP data from moorings are  
108 mainly coming from coastal areas with few exceptions such as the deep-water tropical moored arrays. Ice  
109 buoys are also used to make observations, including SLP over sea-ice in the Polar Regions.

110 With regard to satellite observations, while SLP pressure field gradients can be estimated relatively  
111 well from the satellite derived Surface Vector wind observations over the oceans (mainly from  
112 scatterometer data), it is not possible to anchor adequately the surface pressure field with these satellite  
113 data alone. Global Navigation Satellite Systems radio occultation (GNSSRO) provides useful information  
114 on the SLP field. However, such data are also sensitive to atmospheric temperature and humidity profiles,  
115 and small biases in prior knowledge of these variables lead to biases in retrieved surface pressure  
116 estimates. Because of this problem, GNSSRO measurements cannot fully compensate for the lack of *in*  
117 situ observations of surface pressure [Healy, 2013].

118 The data denial study, or observing system experiment (OSE hereafter), discussed in this essay was  
119 performed at ECMWF and was designed to quantify the effect of the SVPB drifter data only. The principle  
120 of the OSE is that a data assimilation and forecast models, the ECMWF Integrated Forecast System (IFS)  
121 four-dimensional variational (4D-Var) system [Janisková and Lopez, 2013; Rabier et al., 2000] in this case,  
122 is used to produce a control run, in which all of the available data are assimilated, and also a data denial  
123 run in which the SLP observations from the SVPB drifters are withheld. The experiments were run at

124 forecast model resolution around 40km horizontally and 91 model levels vertically. The error of the two  
125 runs is computed for selected variables by comparing them with the higher resolution operational ECMWF  
126 analyses. Two OSEs were performed, one for November-December 2010 and one for July-August 2012. In  
127 the former, 879,107 SLP-SVPB data were denied, corresponding to 96.3% of the available SLP data from  
128 all buoys and, in the latter, 720,257 SLP-SVPB data were denied, corresponding to 94.8% of the available  
129 SLP data from all buoys (Figure 2). The two periods were chosen because the amount of SLP data was  
130 largest and to contrast two different seasons. The full details of the OSE experiment are discussed in  
131 Horányi et al. (2016) and in this essay the main results are highlighted.

### 132 **The Effect of the SLP Data from Drifters on Weather Forecast is Significant**

133 In the following discussion, and for the sake of brevity, only the results from one of the two seasons  
134 are shown since they are very similar. A first proof of the significant influence of the SVPB data is given by  
135 the mean difference, up to 0.7 hPa, of the SLP analyses between the control and the denial experiments  
136 (Figure 3). The NWP analysis is represented in the model grid (around 40km horizontal resolution) and the  
137 differences between the sea level pressure fields of the control and denial experiments are computed and  
138 then averaged over the 2 months. Therefore, the averaged differences shown in Figure 3 are smooth and  
139 relatively small due to the time averaging. It should be noted that the differences between the control  
140 and denial experiments computed for each assimilation/denial run are significantly larger than the  
141 accuracy of the barometers.

142 The largest SLP differences, which show also a seasonal dependence, are found where the majority of  
143 the SVPB data were collected, particularly in the Arctic, in the Southern Ocean and in the North Atlantic.

144 Interestingly, a comparison between Figure 2 and Figure 3 for winter 2010 shows that even the denial  
145 of few SVPB drifters in the tropical eastern Pacific (at about 20°N, 140°W), in the western tropical Atlantic  
146 (at about 15°N, 45°W) and in the equatorial Indian Ocean (at about 0°N, 95°E) can have a large effect on

147 the initial conditions, suggesting the importance of *in situ* SLP data at low latitudes where SVPB drifters  
148 are not normally deployed apart from targeted small arrays in the paths of some tropical cyclones.

149 The normalized (by the control) SLP root mean-squared (RMS) forecast error differences between the  
150 control and denial experiment (Figure 4) clearly indicates a substantial forecast degradation up to 72 hours  
151 ahead when the drifter data are denied. The beneficial effect of the drifter data is most pronounced in  
152 the southern hemisphere and in the Arctic Ocean. However, substantial drifter positive impact can be  
153 seen in all ocean basins as well as in the tropical and equatorial region of the eastern Pacific Ocean.

154 Substantial degradations of the denial run are also found for the near-surface (1,000 hPa) wind  
155 forecast (not shown). One remarkable result of this study is that the beneficial effect of the drifter  
156 observations is not just limited to the surface, but extends high into the troposphere. For example, the  
157 normalized geopotential height RMS error differences (Figure 5, left panel) show that the strong beneficial  
158 effect of the drifter data is located in the subtropical region of the southern hemisphere as well as in the  
159 equatorial and tropical regions up to 250 hPa (Horányi et al. 2016). The beneficial effect of the SVPB data  
160 (up to 5 days ahead), for the geopotential height, lasts longer in the equatorial and tropical regions.  
161 Similarly, the drifter data improve the wind forecast in the subtropical and high latitude regions and up to  
162 400-200 hPa (Figure 5, right panels). Full details of the OSE study can be found in Horányi et al. (2016).

### 163 **Other Techniques Indicate Positive Impact of the SLP Data from Drifters on Numerical Weather** 164 **Prediction**

165 OSEs are expensive to run but have the benefit of quantifying the effect of a single perturbation, in  
166 this case the denial of the SVPB drifter data, on all forecast metrics (root mean-squared error for instance)  
167 at all forecast ranges. In contrast, the adjoint-based Forecast Sensitivity Observation Impact (FSOI)  
168 quantifies the value of any or all components of the observing system on a specific measure of forecast  
169 impact when the entire observational dataset is present in the assimilation system [Cardinali, 2009; Gelaro

170 *et al.*, 2007; Langland and Baker, 2004; Zhu and Gelaro, 2008]. While much less expensive than OSEs, the  
171 reliance of the FSOI technique on the adjoint model, and the inherent assumption of linearity, restricts its  
172 use to forecast ranges of less than two days for most global applications.

173 Here we use the FSOI to quantify the contribution of each assimilated observation in the ECMWF IFS  
174 to the reduction of a measure of 24-h global forecast error combining wind, temperature and surface  
175 pressure in terms of a dry energy norm. A comparative diagram of these contributions (Figure 6) shows  
176 that the SVPB drifters have the largest impact on a per-observation basis of all the data types assimilated  
177 in the ECMWF system during this period. It should be noted, however, that the combined impact of all  
178 SVBP observations is still relatively small (not shown) since they are far fewer in number than many other  
179 data types, especially compared to satellite observations. More details of the FSOI evaluation can be found  
180 in Horányi et al. (2016). Additional information comparing the impact of various observing systems on  
181 NWP, including the high impact on a per observation basis of SLP from drifters can be found in WIGOS  
182 [2012]

### 183 **Climate Applications of *in-situ* SLP Data over the Ocean and Final Remarks**

184 Global and accurate SLP observations are important because they allow the description, with a good  
185 approximation, of the geostrophic, barotropic global atmospheric circulation [Blunden and Arndt, 2013],  
186 which accounts for the largest part of the total atmospheric circulation.

187 Climate changes are often felt through changes in ocean temperature, ocean circulation, sea-level rise  
188 and, perhaps even more dramatically, through changes in air temperature and atmospheric circulations  
189 (e.g. increase/shift of storminess and of extreme events). Changes in the atmospheric circulations will also  
190 impact ocean surface waves and wind regimes such as the monsoons, the hydrological cycle via  
191 modifications of the precipitation and evaporation patterns, and can potentially induce localized climate  
192 changes that will likely have high impact on society.

193 The latest International Panel for Climate Change (IPCC) Report [2013] states that humanity has likely  
194 altered global SLP patterns. Notable examples include the likely link between stratospheric ozone  
195 depletion and the positive trend of the Southern Annular Mode (SAM) in the Austral Summer, and a  
196 poleward shift of the southern Hadley Cell during Austral summer [IPCC, 2013]. The former is directly  
197 correlated to sea level pressure changes over the high latitudes/subtropics [IPCC, 2013].

198 SLP data, including those collected by the SVPB drifters, are used in multiple ways by climate scientists,  
199 including for trend computations, for climate model diagnostics, and for constructing climate indexes.  
200 Long-term mean-SLP changes also affect the mean sea-level due to the inverse barometer effect (a  
201 variation of 1 mbar corresponds, approximately, to a change of 1 cm in sea level), thus providing another  
202 strong rationale for the deployment of a global array of SVPB drifters.

203 The OSE and FSOI studies described in this essay further highlight in a quantitative fashion the crucial  
204 role of the SVPB drifter data, collected by the US Global Drifter Program and its international partners that  
205 operate under the DBCP umbrella, in improving short- to medium-range NWP. The study periods have  
206 sufficiently large number of drifter-borne observations and are of sufficient length to detect statistically  
207 significant beneficial effects and positive impacts of the data.

208 The beneficial effect of the SVPB drifter data in the forecast is detectable not only near the surface,  
209 but also higher in the troposphere, up to 250 hPa (Horányi et al. 2016). The largest beneficial effect is  
210 observed in the mean sea level pressure field forecast, but also the predicted wind field is significantly  
211 affected.

212 The reduced improvements of the SLP forecast in the OSE simulation in the tropical and equatorial  
213 regions can partly be attributed to the relatively small variability of the signal at low latitudes, but it should  
214 also be noted that very few drifter SLP observations were available there for the denial. Furthermore,  
215 when even a few *in situ* data points are available in the tropics or at the equator, the local beneficial effect

216 is large (see e.g. Figure 4 and Figure 5). This suggests that an attempt to extend the SVPB drifter array to  
217 the tropical region should be made and the impact of the data should be monitored and quantified with  
218 FSOI or similar diagnostics.

219 The FSOI analysis indicates that The SLP drifter data is the most valuable per-observation contributor  
220 from the Global Observing System (see Horányi et al. 2016 for a more complete discussion details). The *in*  
221 *situ* drifter SLP observations are extremely valuable to anchor the global surface pressure field and  
222 significantly contribute to accurate marine weather forecasts, especially in regions where no other *in situ*  
223 observations are available. All these results give evidence that surface pressure observations of drifting  
224 buoys are essential ingredients of the Global Observing System and their quantity, quality and distribution  
225 should be preserved as much as possible in order to avoid any analysis and forecast degradations.

226 The global drifter barometer array is mainly implemented through international collaboration under  
227 the WMO/UNESCO umbrella. The barometer upgrade program offered by the US GDP, under which GDP-  
228 funded drifters can be equipped with partner-funded accurate air pressure sensors, is a practical example  
229 of how such collaboration is executed. Entities interested in this upgrade program can contact the GDP  
230 offices located at the Scripps Institution of Oceanography or at the Atlantic Oceanographic and  
231 Meteorological Laboratory. The participation in the activities of the DBCP, where the implementation of  
232 the SVPB array is discussed every year, is open to all United Nations member states  
233 (<http://www.jcommops.org/dbcp/>). The tasks involved in maintaining the Global Observing System are  
234 demanding and wide and proactive participation of national and international entities is the key for the  
235 success of this program.

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280 **Figure Captions**

281 Figure 1: Schematic of the SVPB drifter manufactured at the Scripps Institution of Oceanography .. 15

282 Figure 2: Total distribution of buoy data in the control experiments. Top: November-December 2010;  
 283 bottom: July-August 2012. The colors indicate the sea level pressure values, in hPa, measured by the  
 284 drifting buoys. In the top panel, 879,107 SLP data locations are shown, and 720,257 are shown in the  
 285 bottom panel. Since the drifters collect and report the data hourly, many point nearly overlap on the  
 286 global scale maps. .... 16

287 Figure 3: Average sea level pressure analyses differences, in hPa, between the control and denial  
 288 experiments. Top: November-December 2010. Bottom: July-August 2012. .... 17

289 Figure 4: Differences of mean sea level pressure RMS errors, normalized by the RMS error of the  
 290 control, between the control and denial experiment for November-December 2010. Red (blue) colors  
 291 indicate degradations (improvements) in the denial experiment. Forecast ranges: 12h, 24h, 48h, 72h, 96h  
 292 and 120h. .... 18

293 Figure 5: Normalized root mean-squared error difference latitude-pressure cross sections between  
 294 the control and denial experiments. November-December 2010. Left: geopotential height, right: vector  
 295 wind normalized root mean-squared error difference latitude-pressure cross sections between the control  
 296 and denial experiments. Significant differences are denoted by hashes. Red (blue) colors indicate forecast  
 297 degradations (improvements) for the denial experiment. .... 19

298 Figure 6: ECMWF operational mean FSOI for the different observing systems for July-August, 2012.  
 299 The FSOI values are also normalized by the total forecast error for easier comparison. The observing  
 300 systems displayed are SYNOP surface observations (surface pressure, moisture and wind), aircraft  
 301 measurements (wind and temperature), drifters and moored buoys (surface pressure and wind from  
 302 drifters and moored buoys), radiosondes (wind, temperature, and moisture), pilot/profiler (wind),

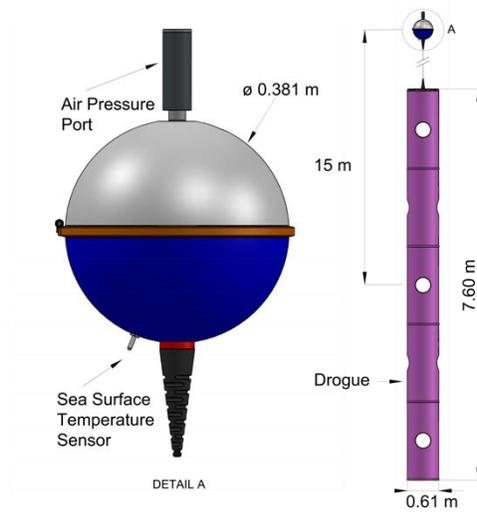
303 geostationary atmospheric motion vectors (wind), scatterometer (surface wind), microwave sounder  
304 radiances (MHS, AMSU-B and AMSU-A), infrared sounder radiances (HIRS, AIRS and IASI), satellite radio  
305 occultation (GPS-RO), geostationary satellite radiances (GEOS-RAD), microwave imager (SSMIS, TMI,  
306 AMSR-E), multi-spectral radiometer (MERIS) and radar precipitation (GBRAD)..... 20

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309 **Figures**

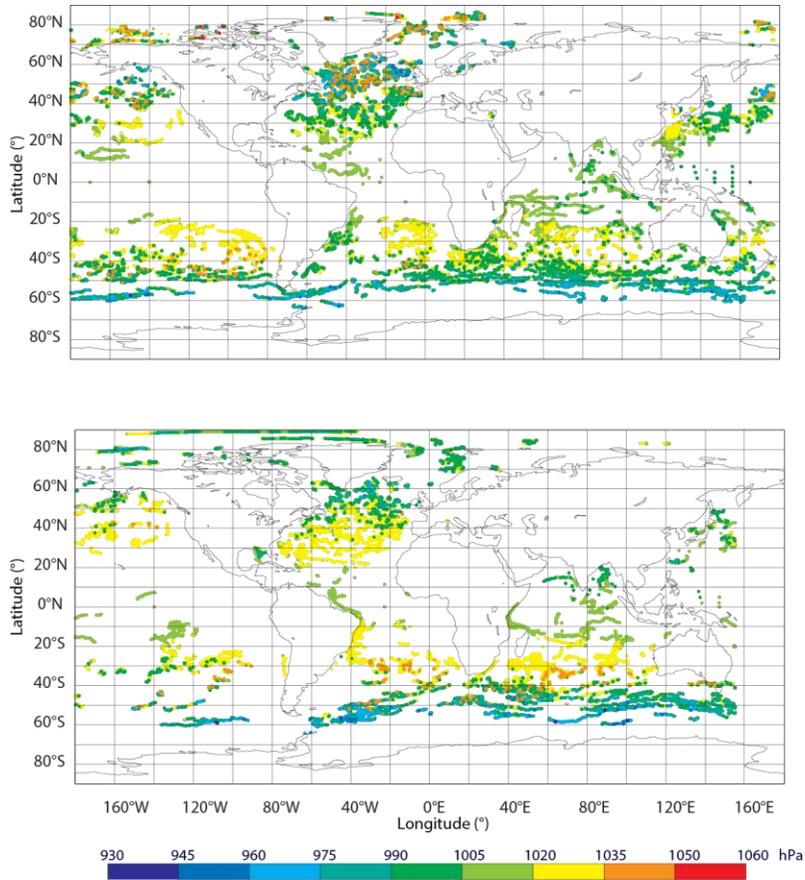
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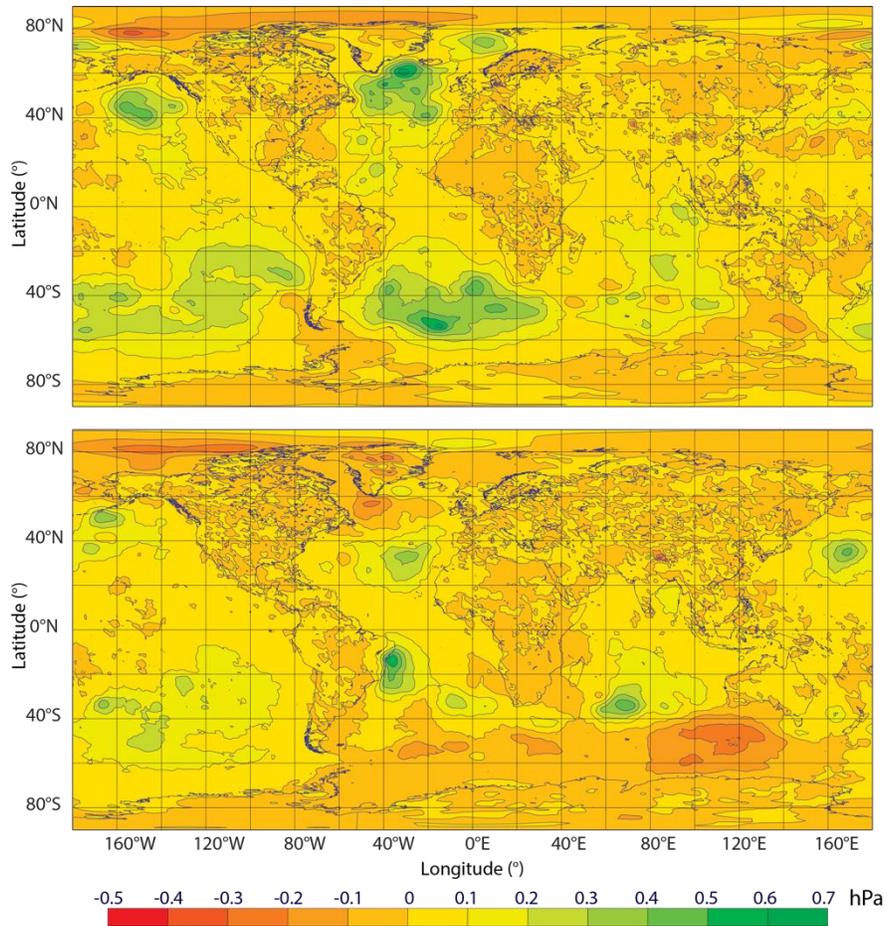
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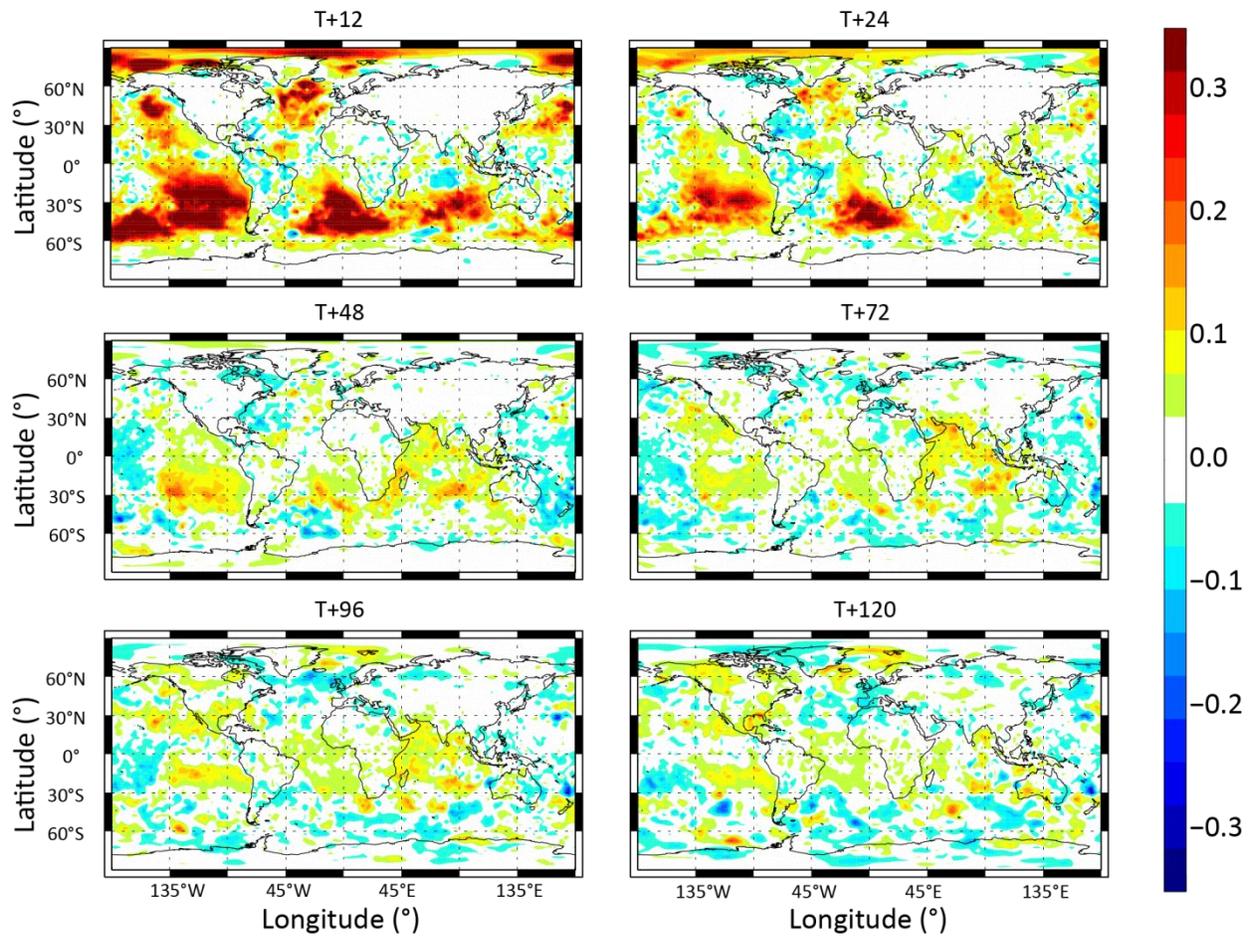
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322 Figure 3: Average sea level pressure analyses differences, in hPa, between the control and denial  
 323 experiments. Top: November-December 2010. Bottom: July-August 2012.

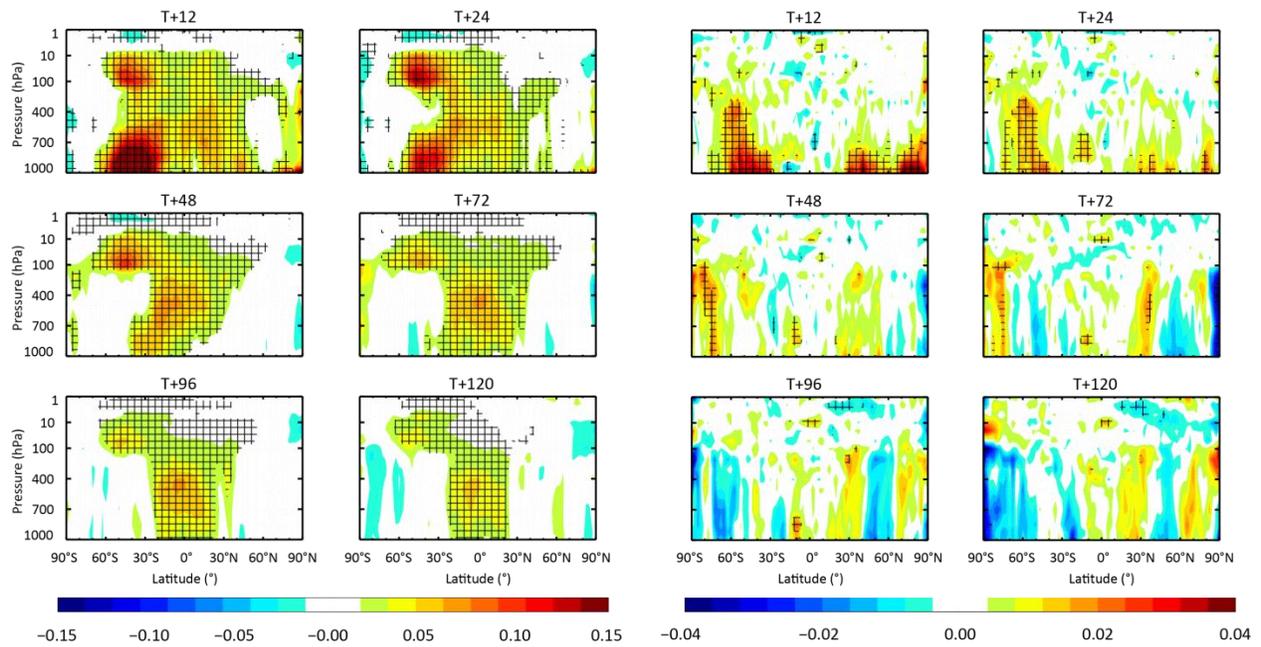
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326 Figure 4: Differences of mean sea level pressure RMS errors, normalized by the RMS error of the  
 327 control, between the control and denial experiment for November-December 2010. Red (blue) colors  
 328 indicate degradations (improvements) in the denial experiment. Forecast ranges: 12h, 24h, 48h, 72h, 96h  
 329 and 120h.

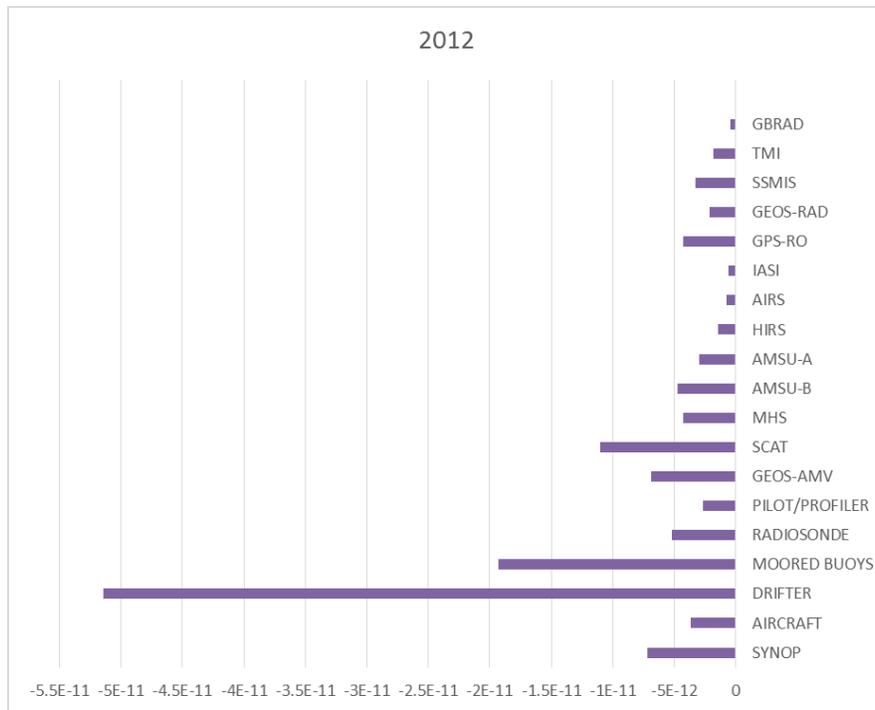
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332 Figure 5: Normalized root mean-squared error difference latitude-pressure cross sections between  
 333 the control and denial experiments. November-December 2010. Left: geopotential height, right: vector  
 334 wind normalized root mean-squared error difference latitude-pressure cross sections between the control  
 335 and denial experiments. Significant differences are denoted by hashes. Red (blue) colors indicate forecast  
 336 degradations (improvements) for the denial experiment.

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339 Figure 6: ECMWF operational mean FSOI for the different observing systems for July-August, 2012.

340 The FSOI values are also normalized by the total forecast error for easier comparison. The observing

341 systems displayed are SYNOP surface observations (surface pressure, moisture and wind), aircraft

342 measurements (wind and temperature), drifters and moored buoys (surface pressure and wind from

343 drifters and moored buoys), radiosondes (wind, temperature, and moisture), pilot/profiler (wind),

344 geostationary atmospheric motion vectors (wind), scatterometer (surface wind), microwave sounder

345 radiances (MHS, AMSU-B and AMSU-A), infrared sounder radiances (HIRS, AIRS and IASI), satellite radio

346 occultation (GPS-RO), geostationary satellite radiances (GEOS-RAD), microwave imager (SSMIS, TMI,

347 AMSR-E), multi-spectral radiometer (MERIS) and radar precipitation (GBRAD).

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